



Forward modelling with application to A-Train observations

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Thanks: Mark Webb, Mark Ringer, Yoko Tsushima, Jonny Williams



Outline

- Motivation
- Radiative transfer
- Examples of model evaluation using satellite data

- Retrievals and forward modelling
- Initial applications

- Applications to A-Train
- COSP
- Concluding remarks

Motivation

- Improve models (parameterizations) to produce better predictions
- Model evaluation: How well do models represent the real world?
- How can we answer this question?
Using observations (satellite data)

“the most scientifically valuable thing that can come out of a comparison of measurements with model results is to show that the model has failed”

(Randall and Wielicki, *BAMS*, 78(3), 1997)

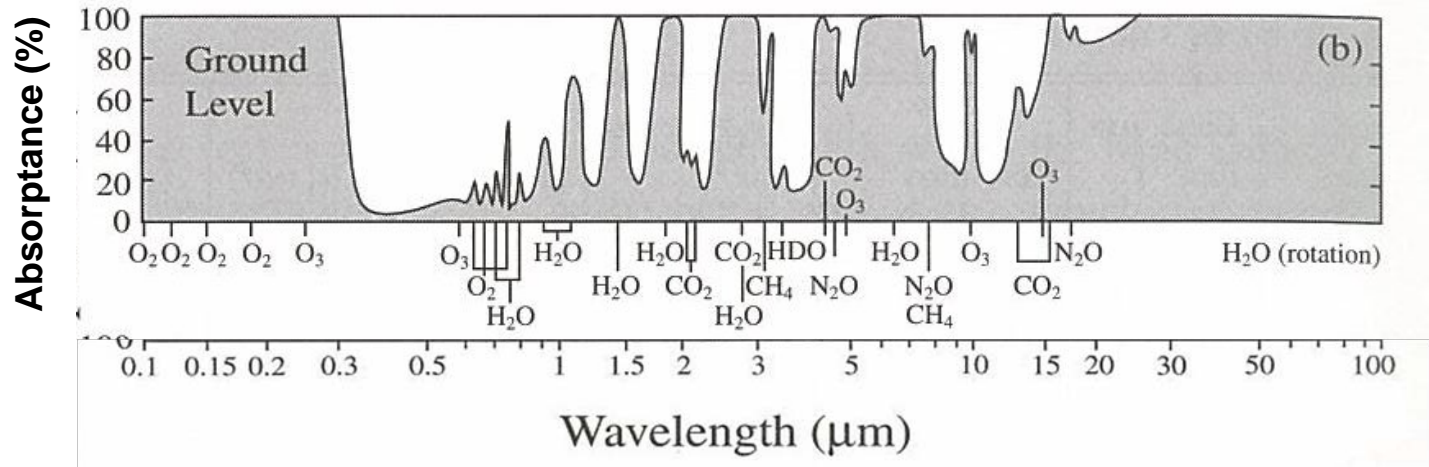
How can we best exploit observations in this context?

- Forward modelling: useful tool to make optimal use of observations
- A-Train: new perspective of clouds/precipitation and interactions with radiation

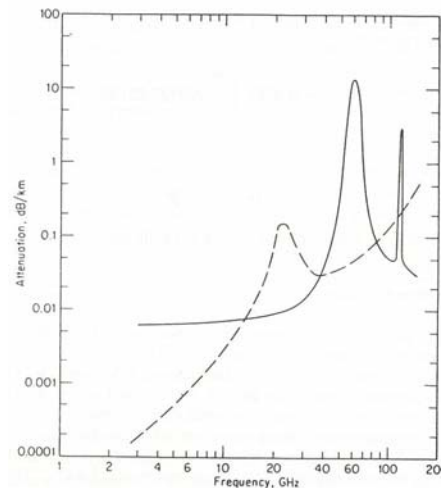


Radiative transfer

Absorption



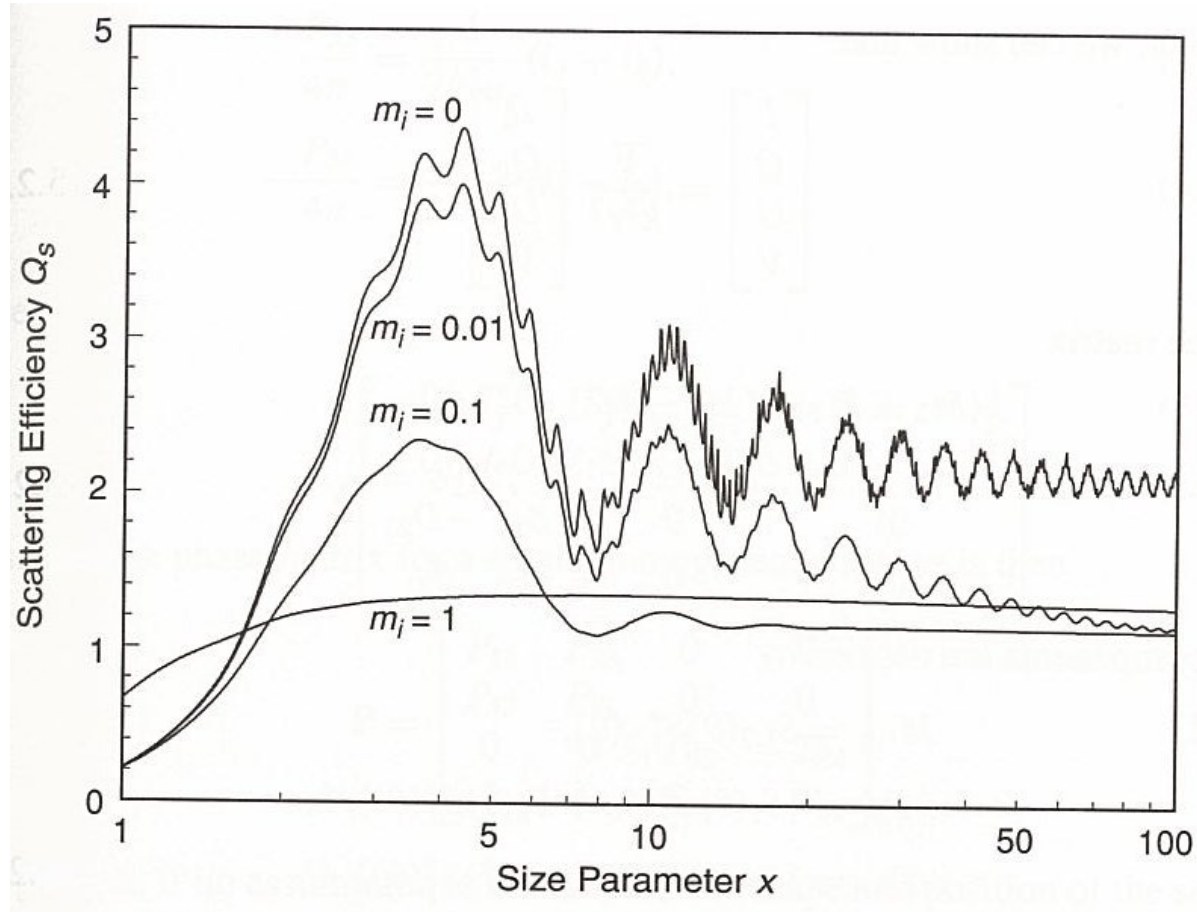
(Thomas and Stamnes, *Radiative transfer in the atmosphere and ocean*, 2002. Fig 11.2)



(Sauvageot, *Radar meteorology*, 1992. Fig 2.13)

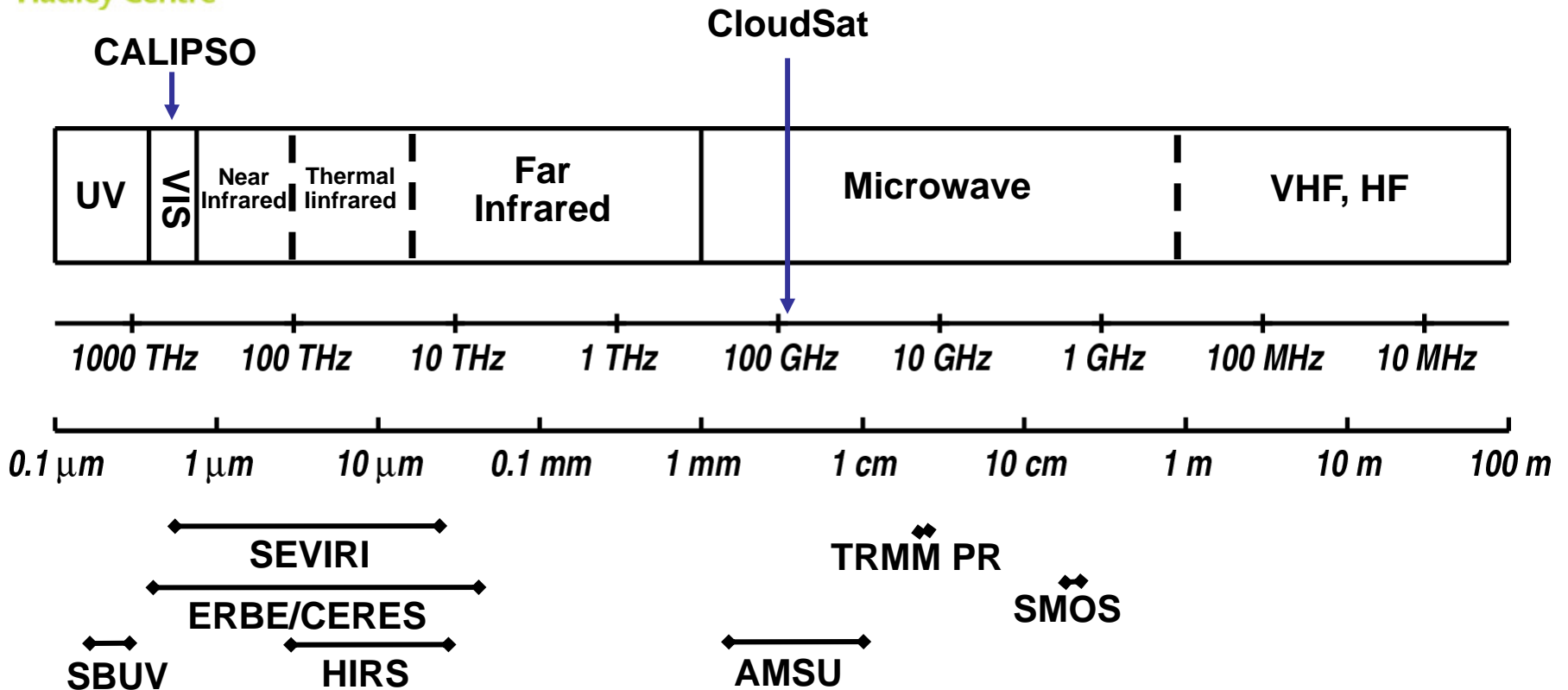
Scattering by particles

$$x = \frac{2\pi r}{\lambda}$$



Dependency on the *ratio* size/wavelength

Electromagnetic spectrum

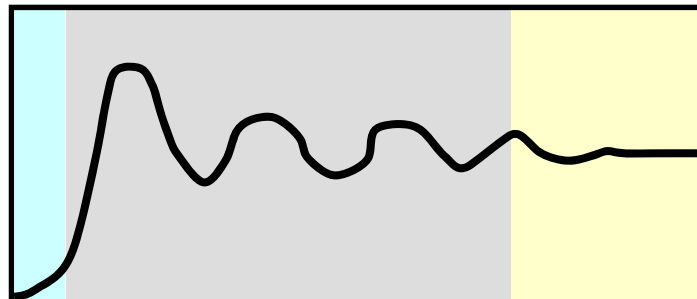


Size parameter

	Size (a)	Size parameter		
		$\lambda=0.5 \mu\text{m}$	$\lambda=10 \mu\text{m}$	$\lambda=1 \text{ cm}$
Aerosol	1 μm	1.26×10^1	6.3×10^{-1}	6.3×10^{-4}
Water droplet	10 μm	1.26×10^2	6.3×10^0	6.3×10^{-3}
Ice crystal	100 μm	1.26×10^3	6.3×10^1	6.3×10^{-2}
Raindrop	1 mm	1.26×10^4	6.3×10^2	6.3×10^{-1}
Snowflake	1 cm	1.26×10^5	6.3×10^3	6.3×10^0

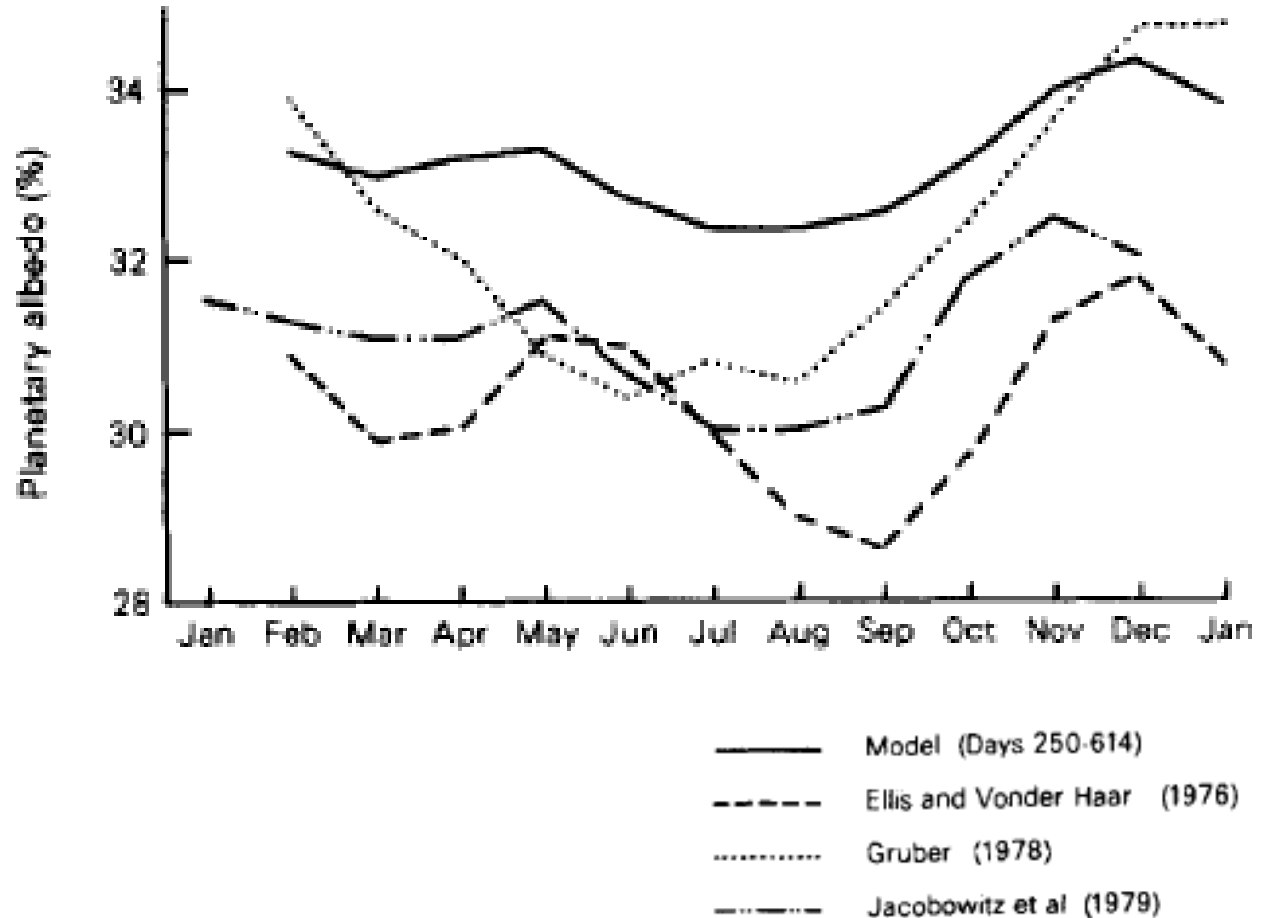
Rayleigh	Mie	G. Optics
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(Adapted from Liou, *An introduction to atmospheric radiation*, 2002. Table 5.1)



Examples of model evaluation using satellite retrievals

A global view



A more detailed view

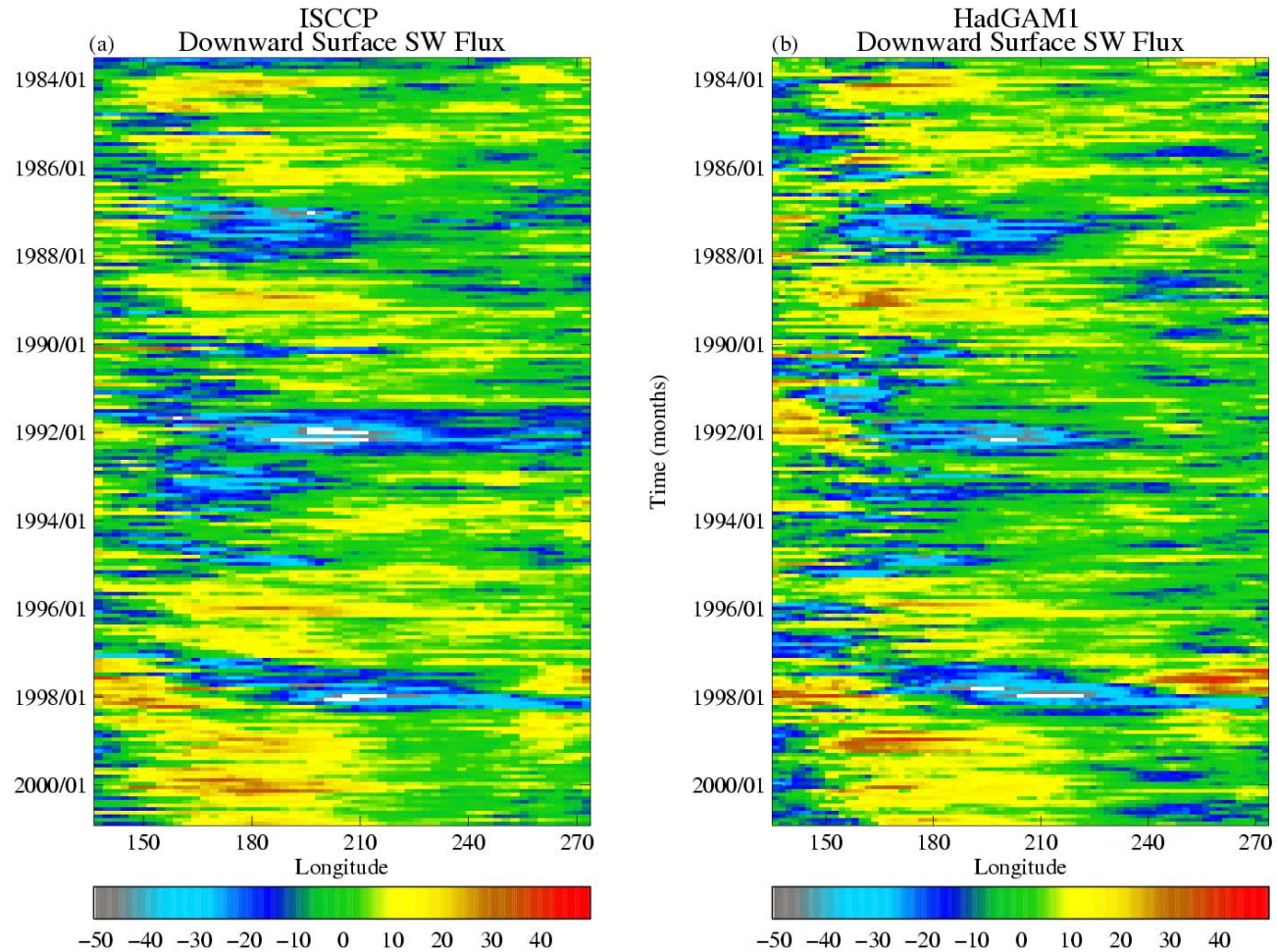


FIG. 10. Hovmoeller plots of the tropical Pacific (10°S , 10°N) as derived from satellite data and simulated by HadGAM1. (a) $S_{s,d}$ from ISCCP-FD, (b) $S_{s,d}$ from HadGAM1, (c) $L_{s,d}$ from ISCCP-FD, and (d) $L_{s,d}$ from HadGAM1.



Large differences in models

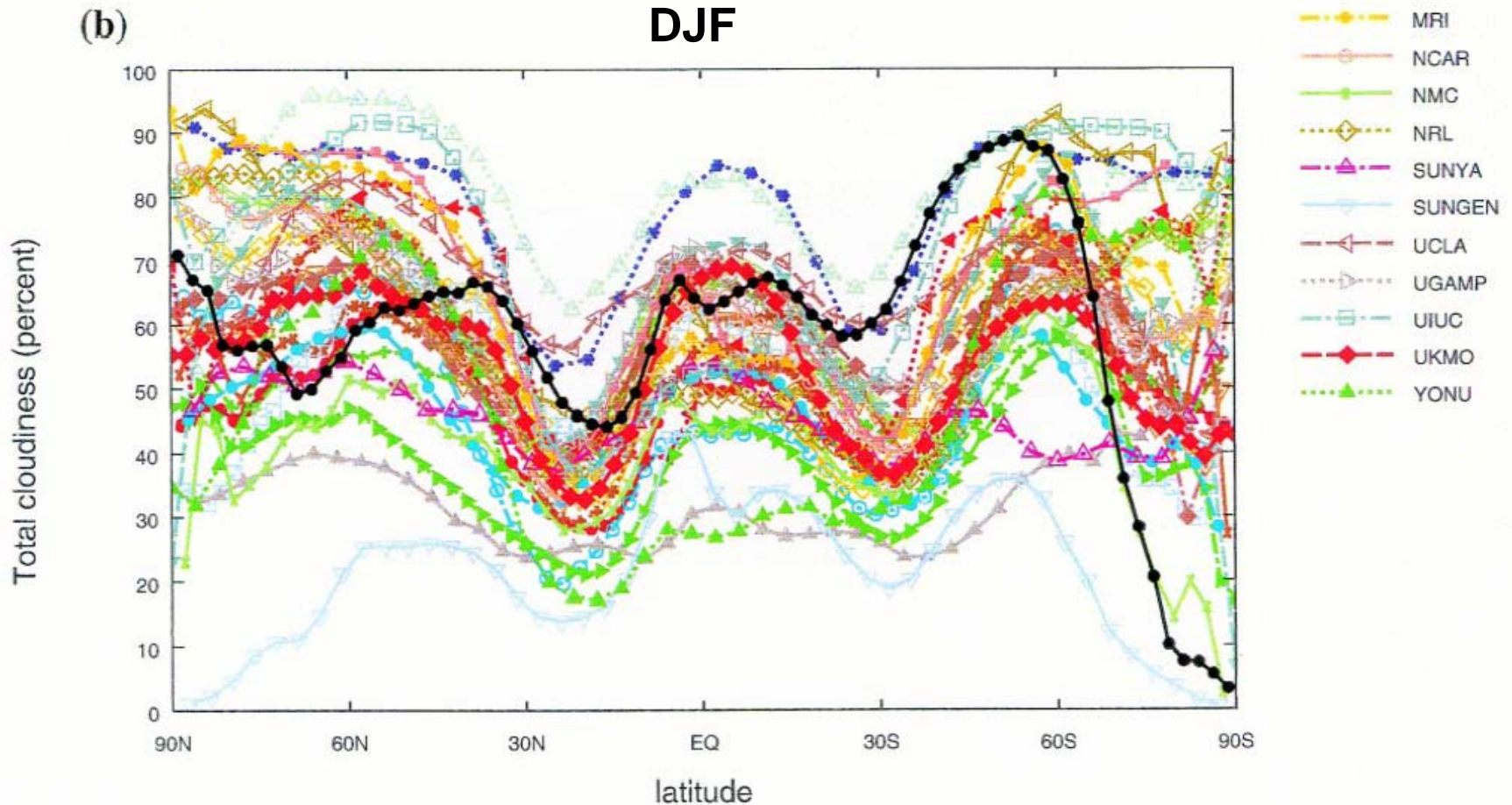
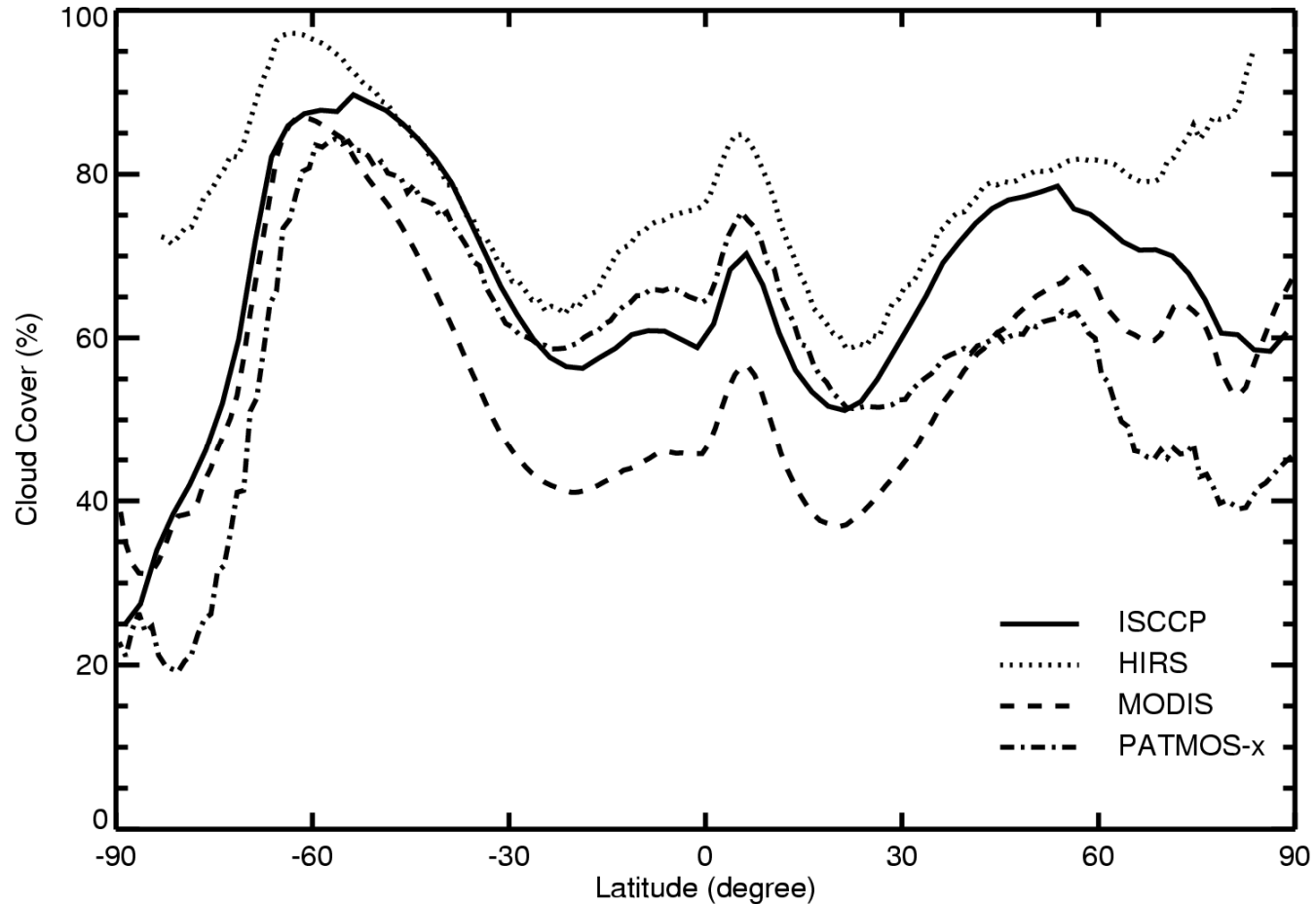


FIG. 7. As in Fig. 6 except for the (a) the outgoing longwave radiation, with observations from the NCEP database (Gruber and Krueger 1984); (b) total cloudiness with observations from ISCCP for 1983–90 (Rossow et al. 1991).

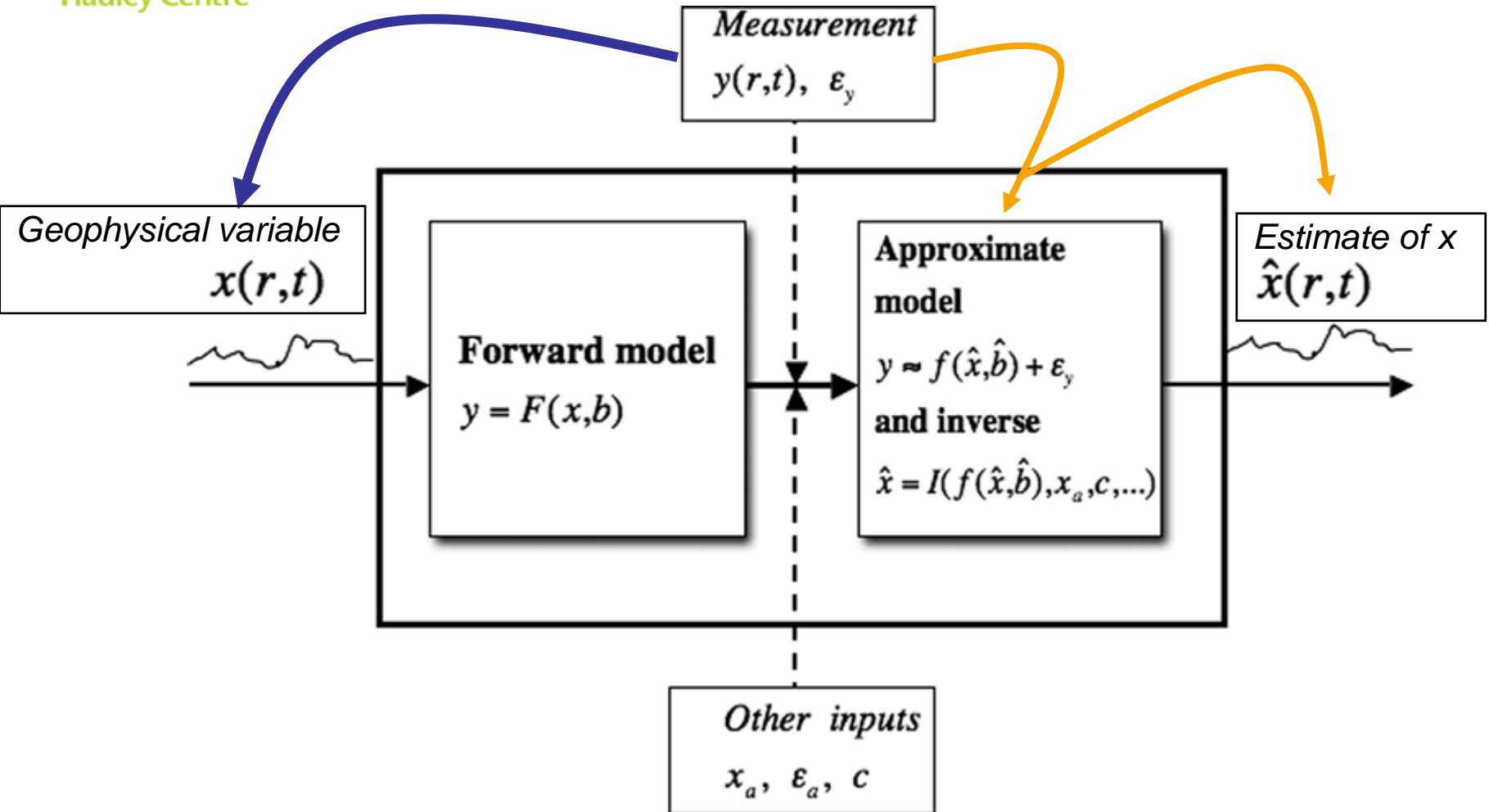
And the retrievals?





Retrievals and forward modelling

Forward modelling at the core of retrievals



The importance of a priori information

$$\hat{\mathbf{x}} - \mathbf{x}_0 = \mathbf{W} \cdot (\mathbf{y}_m - \mathbf{y}_c[\mathbf{x}_0])$$

$\hat{\mathbf{x}}$: vector of retrieved parameters

\mathbf{x}_0 : first guess

\mathbf{y}_m : vector of measurements

\mathbf{y}_c : corresponding vector for first guess

\mathbf{W} : inverse matrix

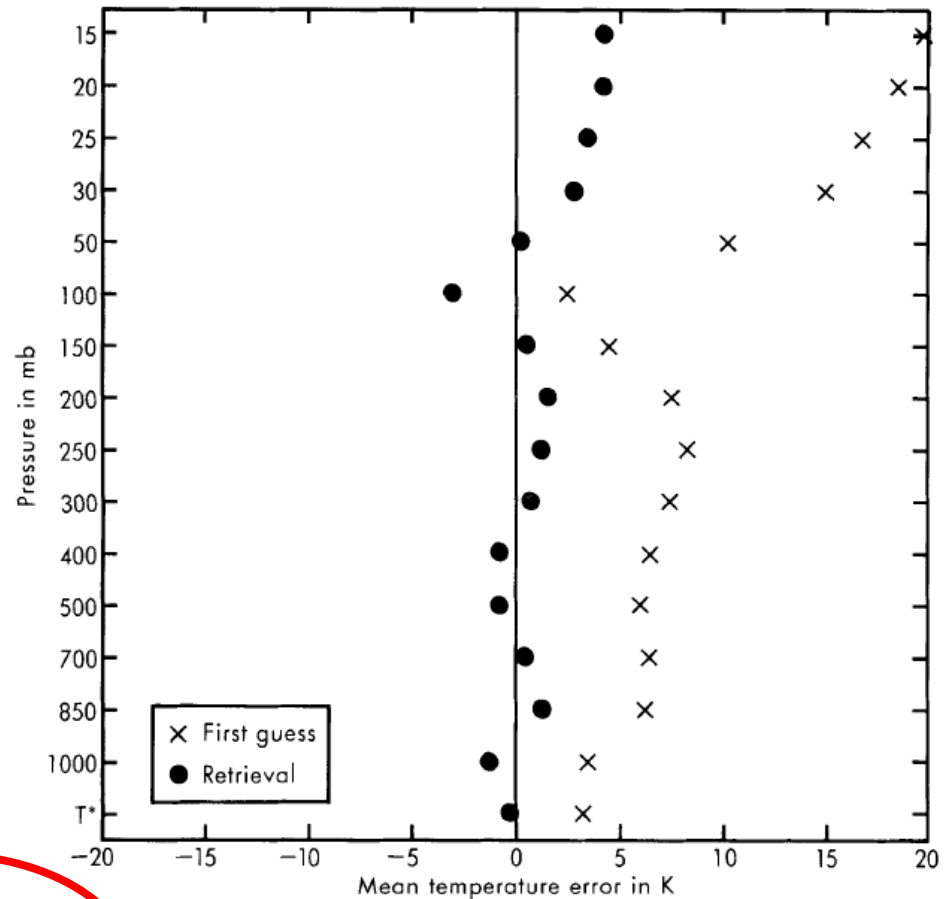
\mathbf{K} : forward matrix

$\mathbf{R} = \mathbf{W} \cdot \mathbf{K}$: averaging kernel

\mathbf{I} : identity matrix

$\mathbf{R} \neq \mathbf{I}$

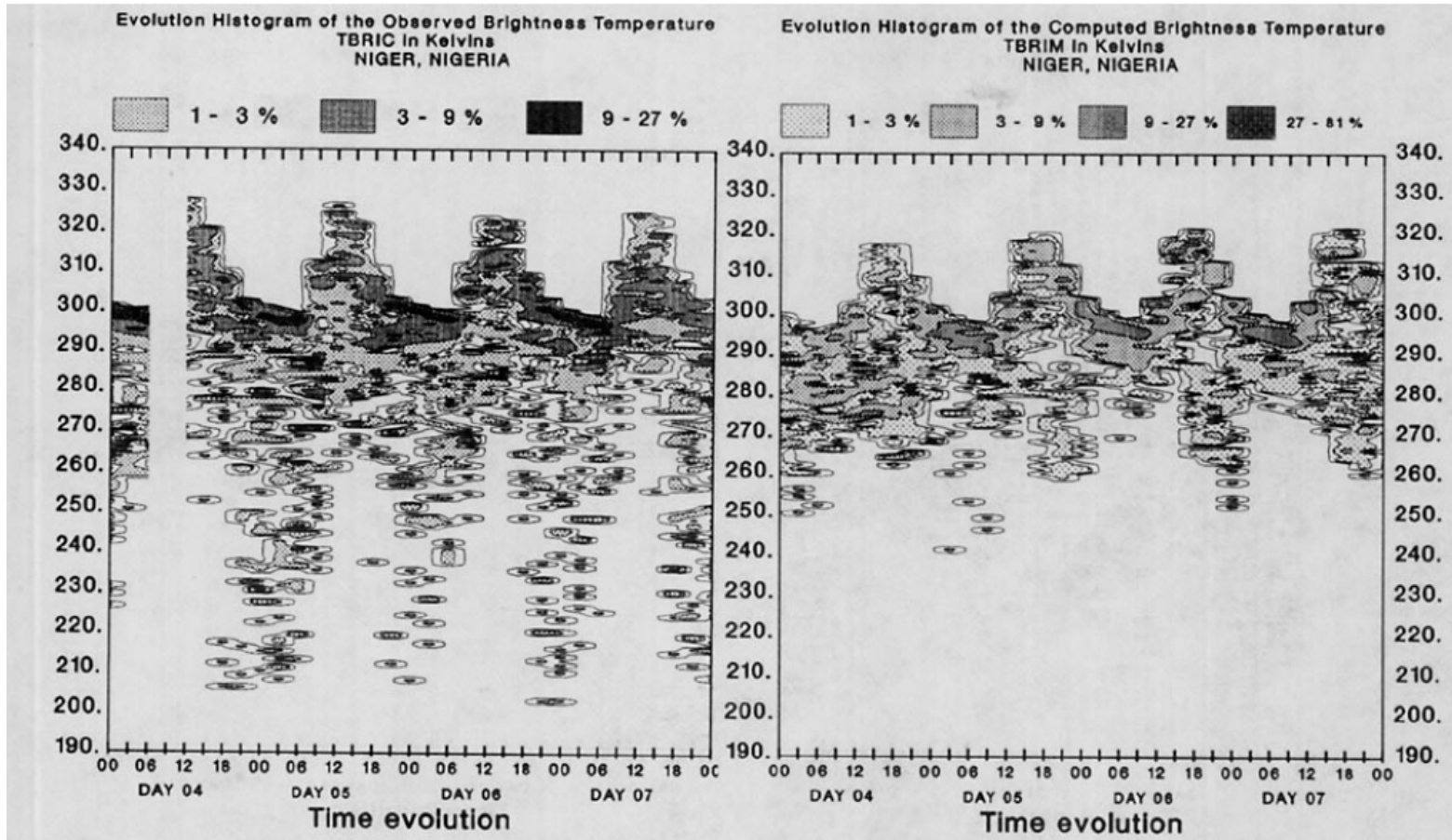
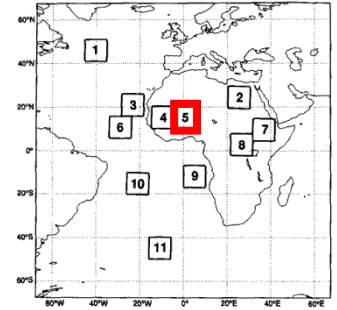
$$\overline{(\hat{\mathbf{x}} - \mathbf{x}_T)} = (\mathbf{I} - \mathbf{R}) \cdot \overline{(\mathbf{x}_0 - \mathbf{x}_T)} + \mathbf{W} \cdot \overline{\epsilon_m}$$





Initial applications of forward modelling and simulators

Diurnal cycle of surface temperature and cloudiness



Upper tropospheric humidity and circulation

July 1987

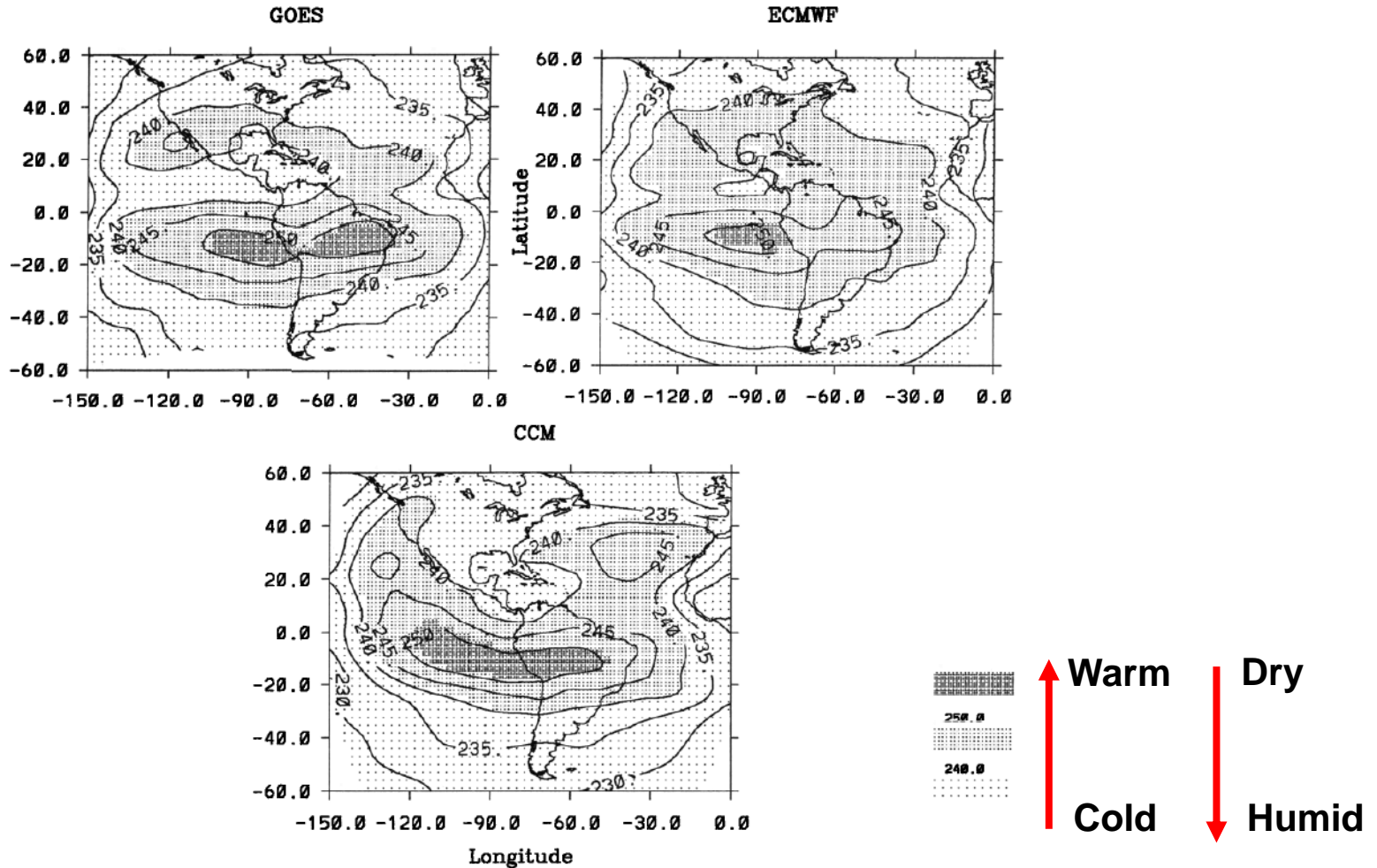
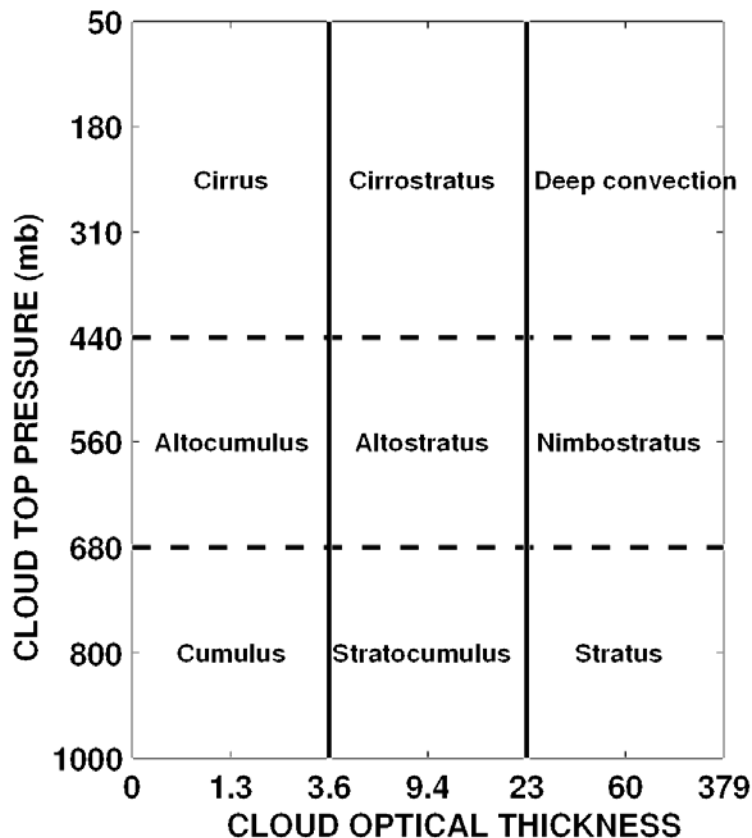


Fig. 12. Monthly mean $T_{6.7}$ from GOES observations, ECMWF analyses, and CCM simulations for (a) July 1987 and (b) January 1988. Units are Kelvin.

The International Satellite Cloud Climatology Project (ISCCP) simulator

ISCCP-D cloud classification

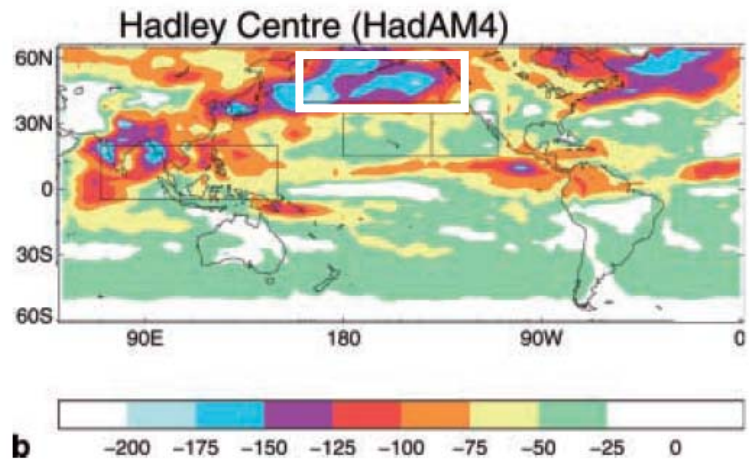
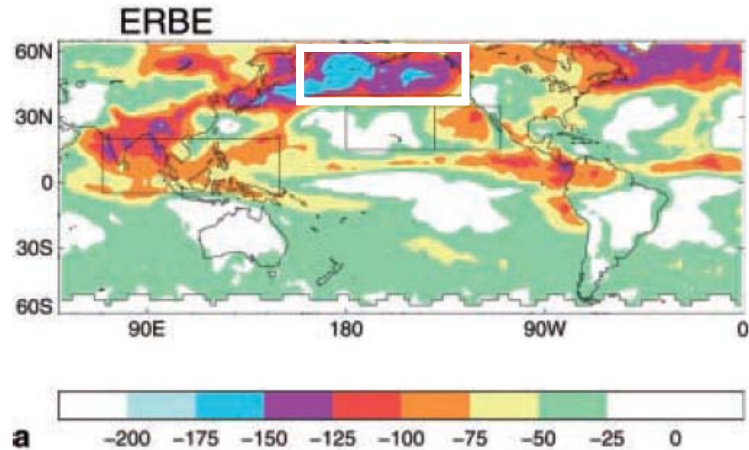


(Rossow et al., *BAMS*, 80, 2261-2287, 1999)

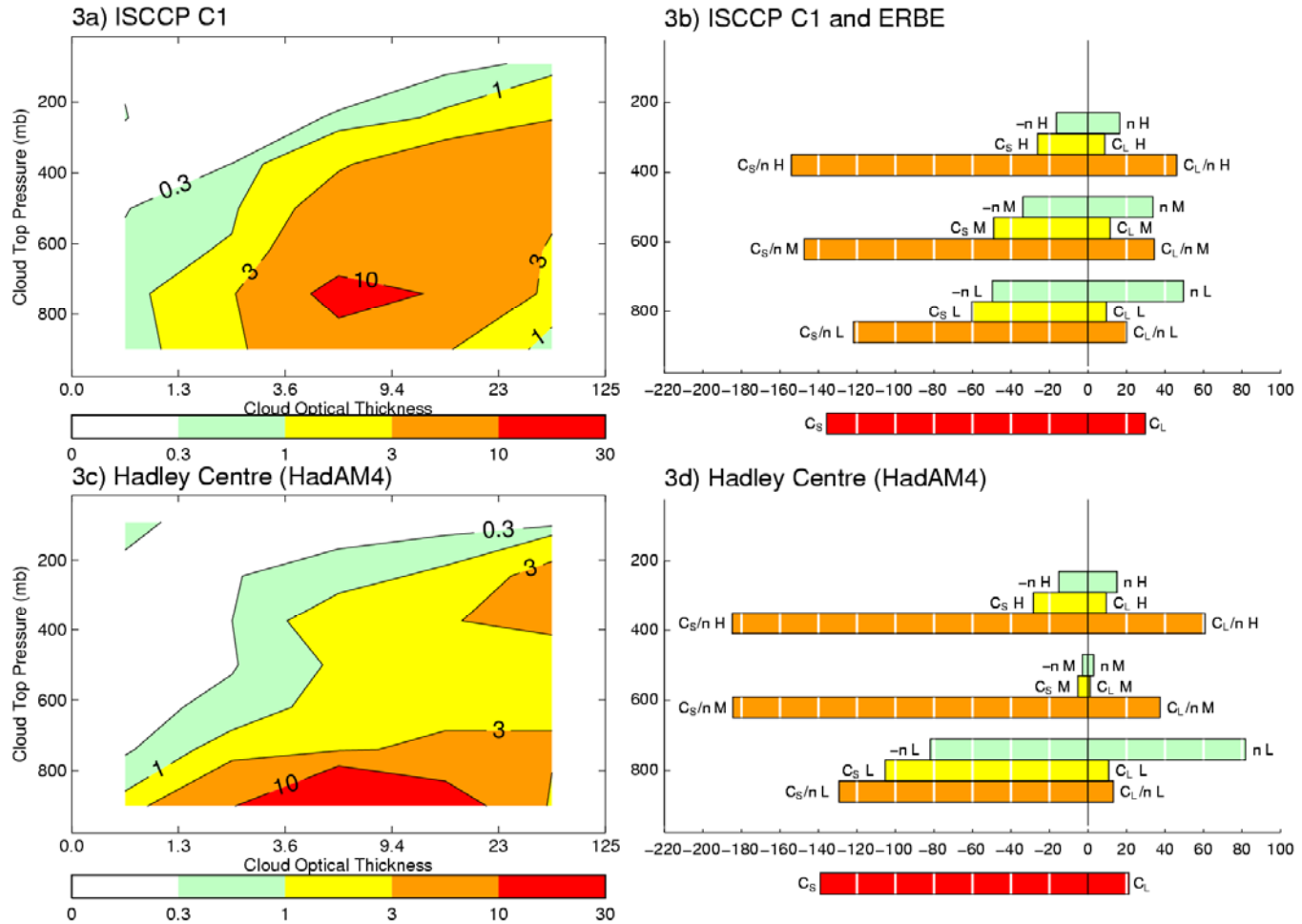
•The ISCCP simulator was developed by S. Klein¹ and M. Webb² to simulate ISCCP cloud data from model data

1: Klein and Jakob, *MWR*, 127, 2514-2531, 1999.
2: Webb et al., *Clim. Dyn.*, 17(12), 2001.

ISCCP simulator reveals compensation of errors



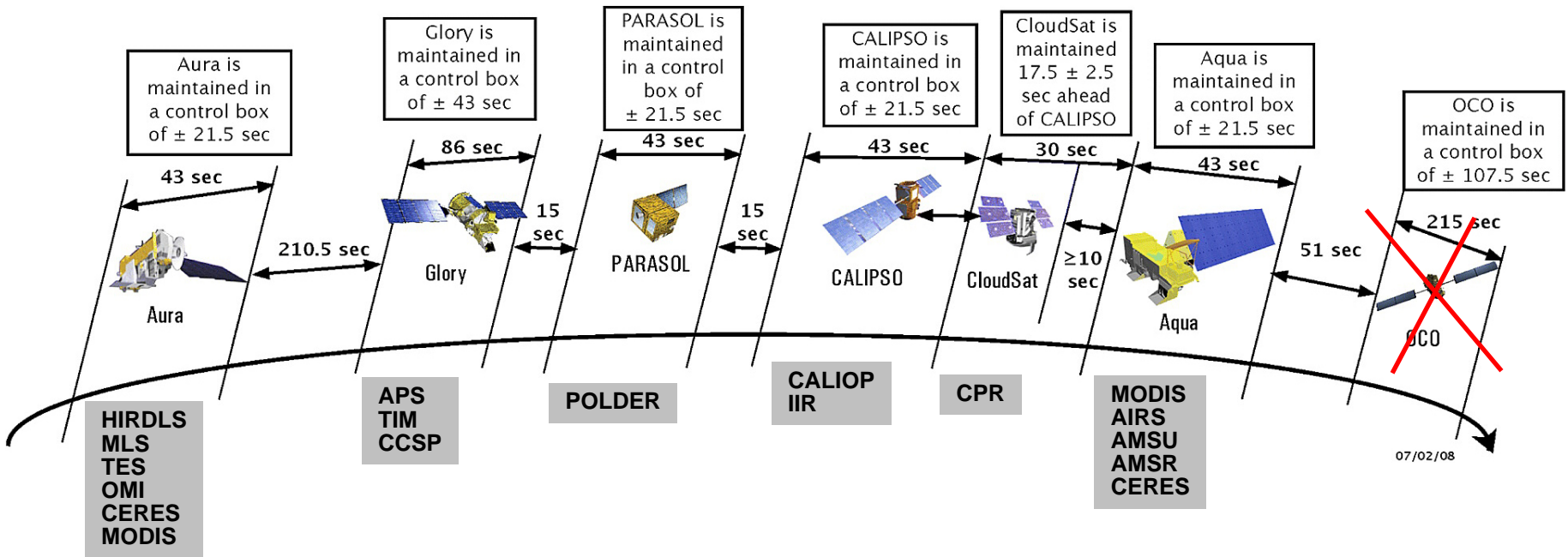
ISCCP simulator reveals compensation of errors





Applications to A-Train

The A-Train



-Synergy: many different instruments observing the same atmospheric column

-CALIPSO and CloudSat: active instruments

CloudSat and CALIPSO



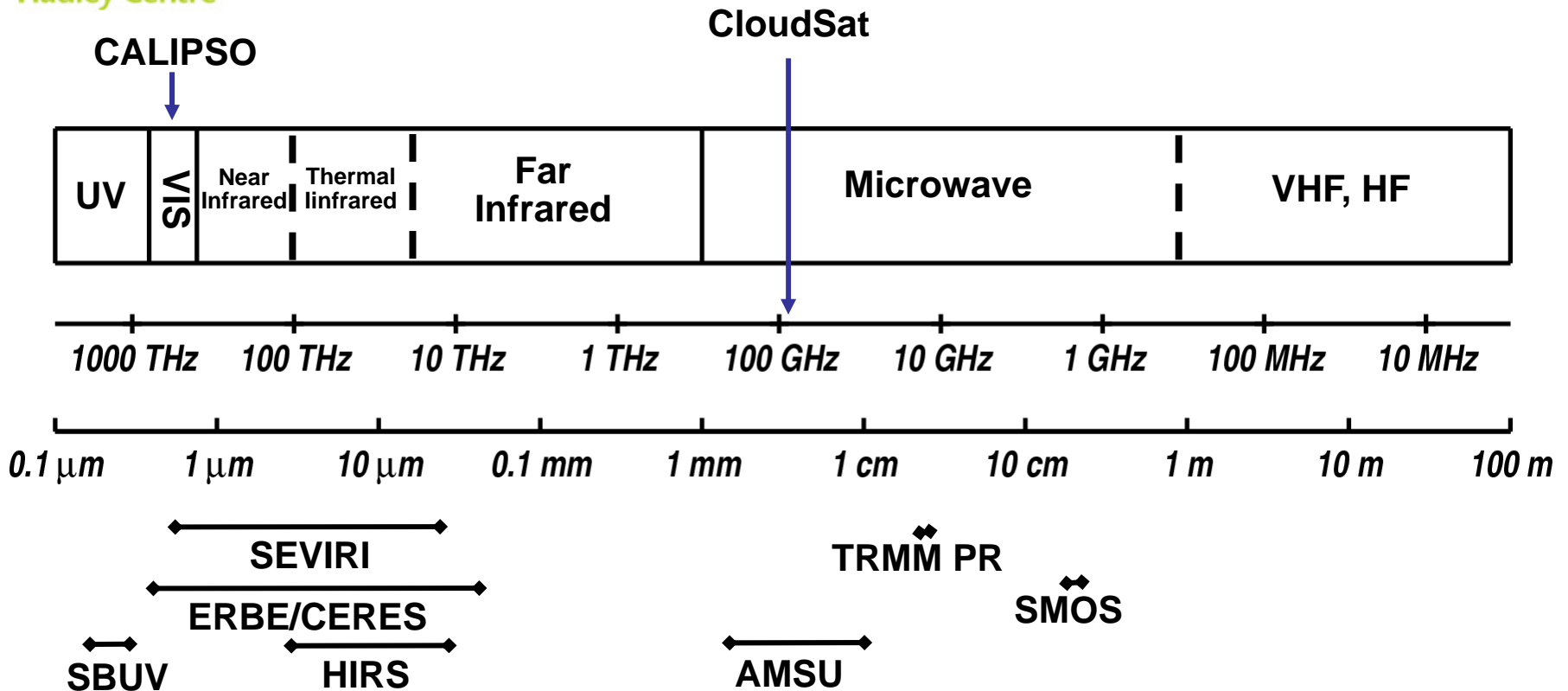
CloudSat:

- Near-nadir pointing 94GHz radar.
- 480m vertical resolution, oversampled at 240m.
- 1.4km x 2.5 km horizontal resolution

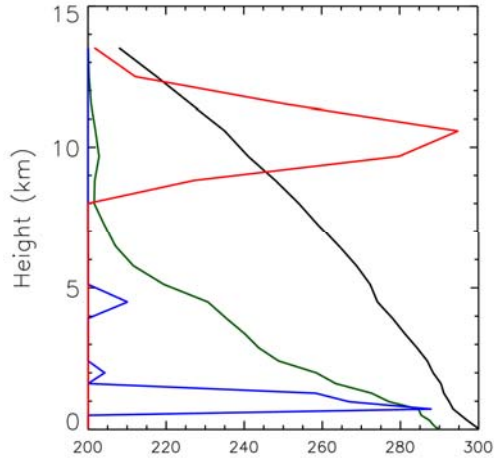
CALIPSO

- Near-nadir pointing lidar (532 and 1064 nm).
- 30m vertical resolution, 60m above 8km
- Horizontal resolution: 333m x 75m, 1km x 75m above 8km

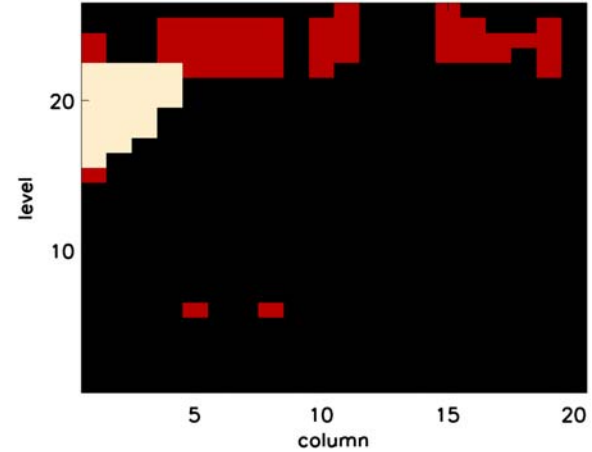
Electromagnetic spectrum



Methodology

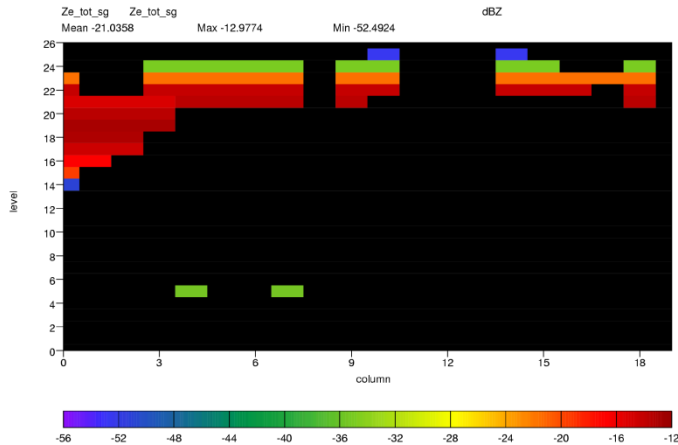


MAX/RANDOM

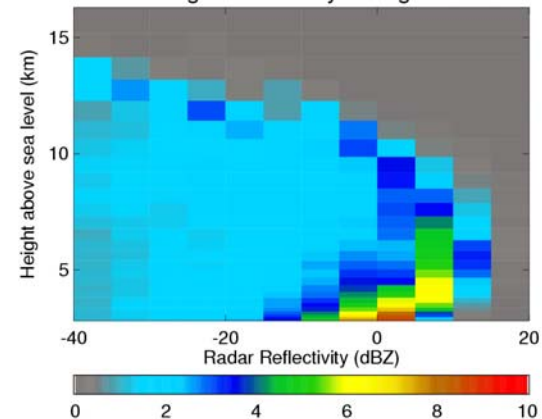


+ Forward model

2.5°x2.5°



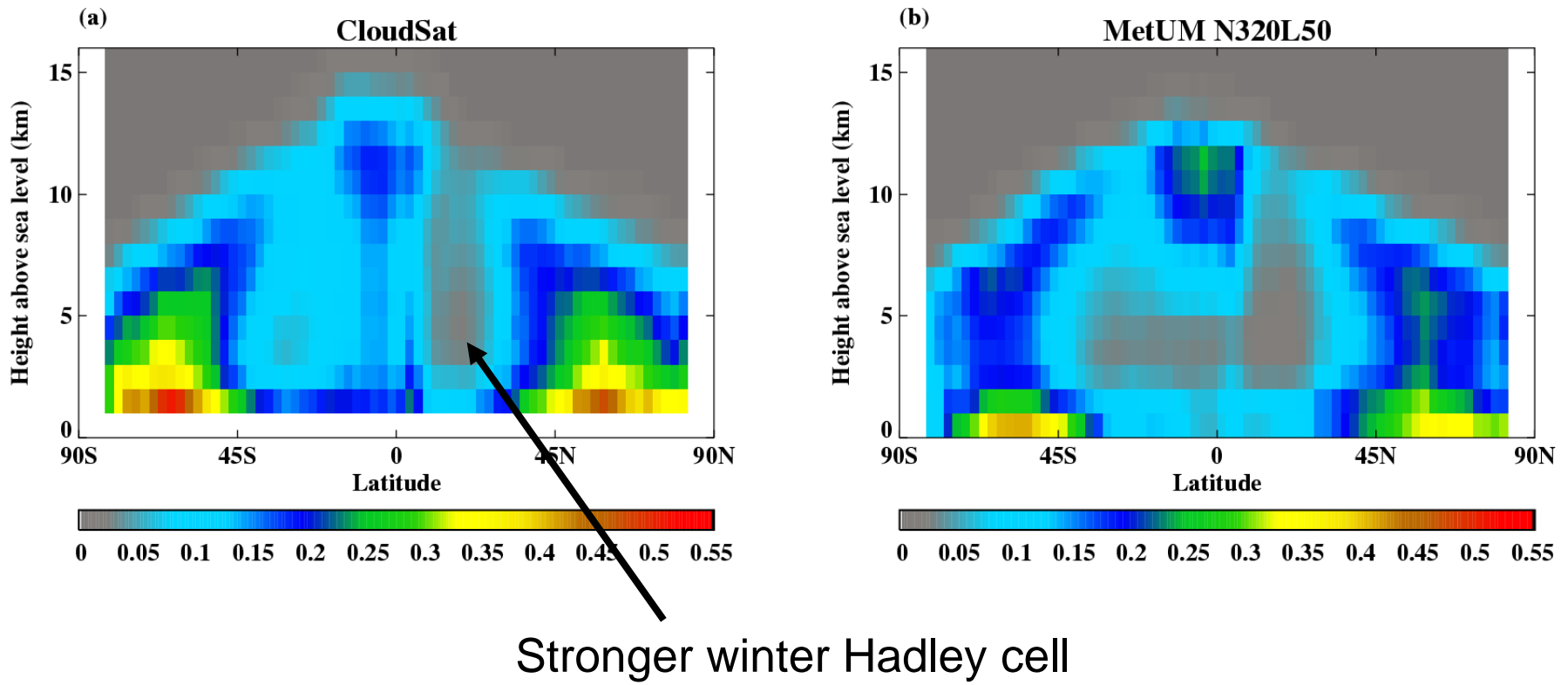
Height-Reflectivity Histogram



CloudSat: a new dimension

Vertical distribution of hydrometeors

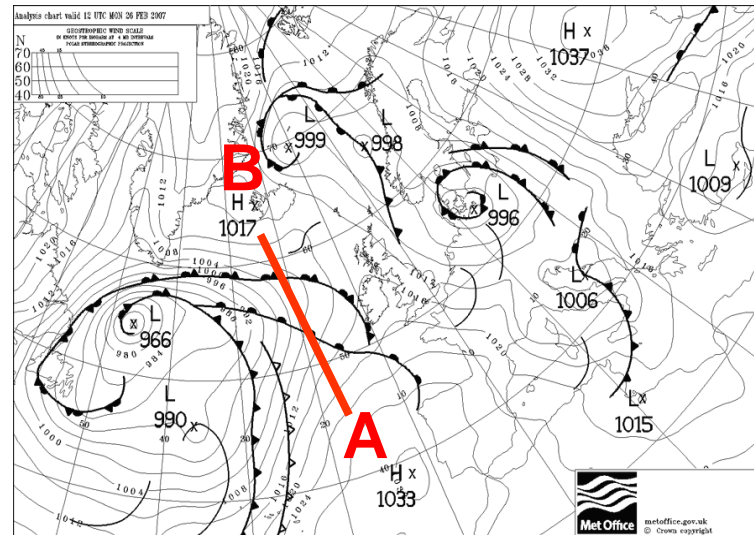
Global - DJF 2006



From 2D...

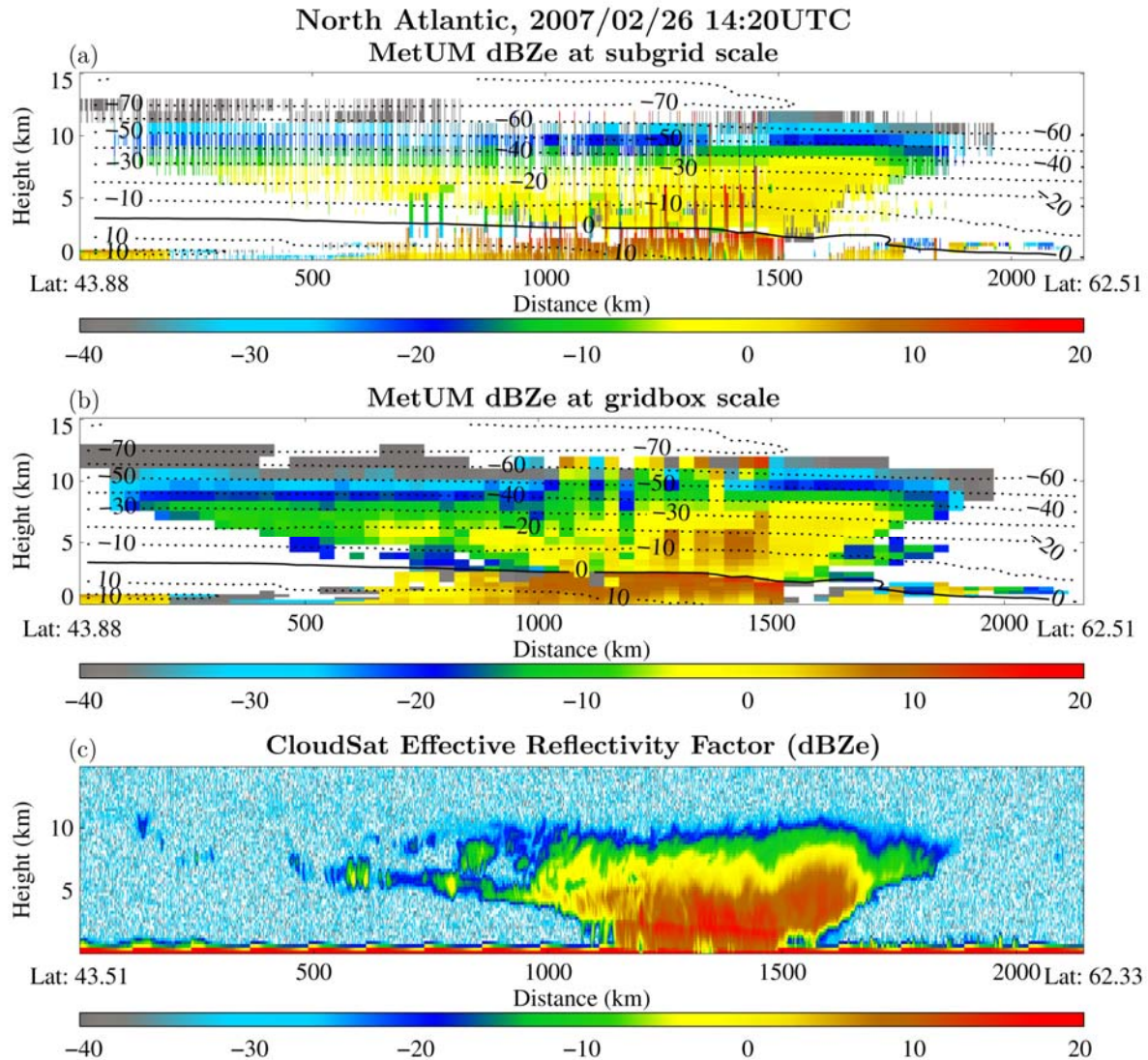
Case study 26/02/2007

- Analysis chart valid at 12 UTC
- CloudSat overpass at ~14:15 UTC



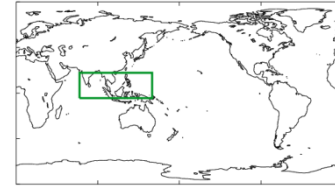
...to “3D”

Case study 26/02/2007

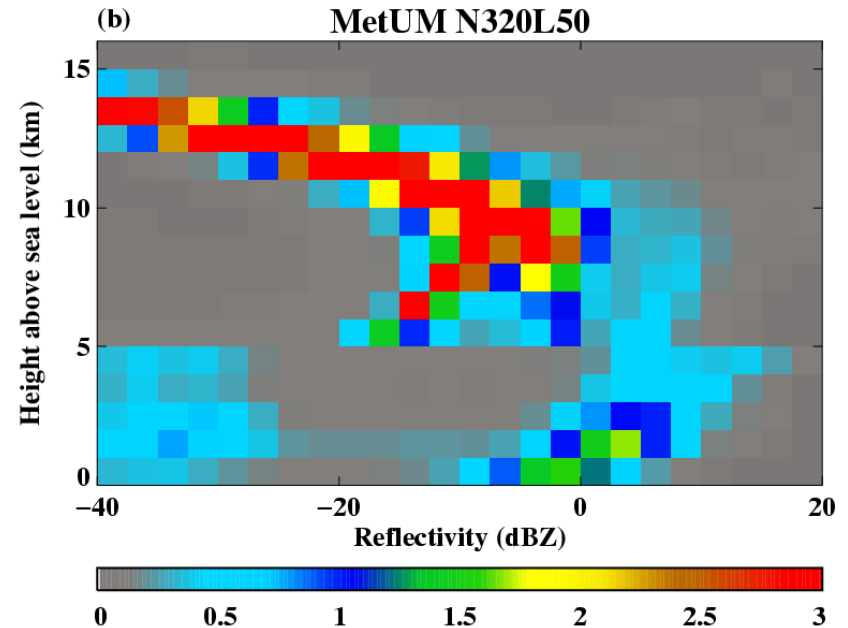
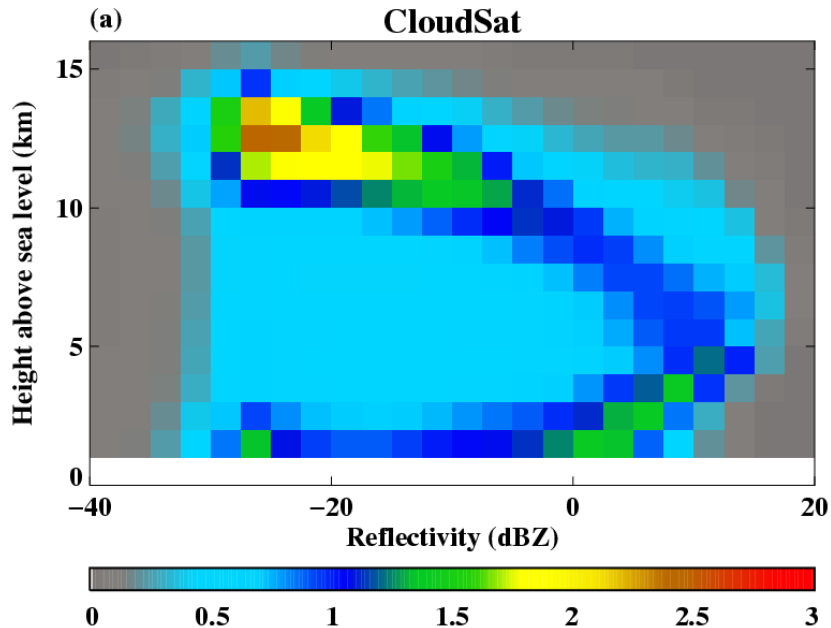


A more statistical approach

PDF(Z) as function of height



Tropical Warm Pool - DJF 2006-2007

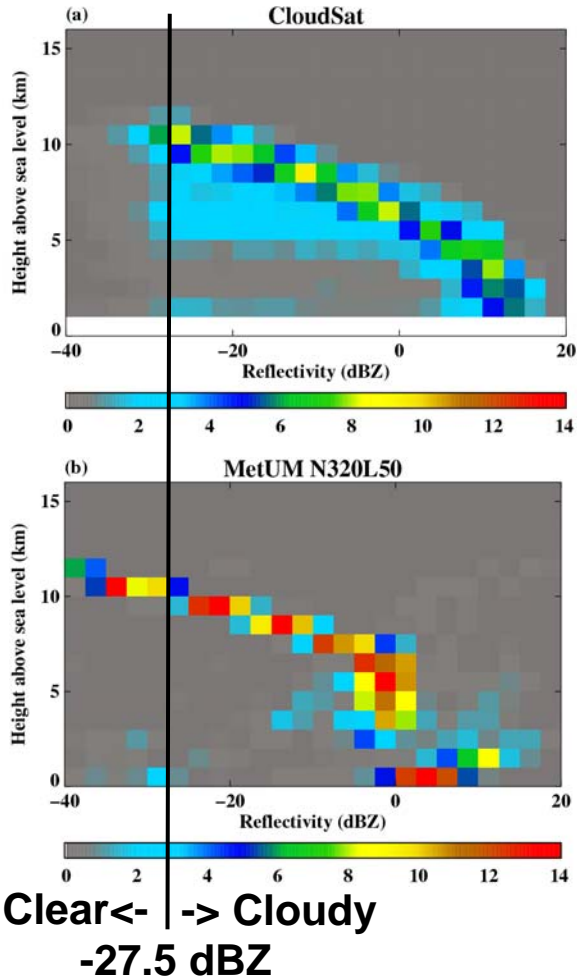


$$Z = \int D^6 n(D) dD$$

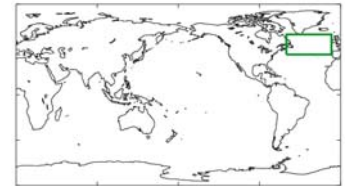
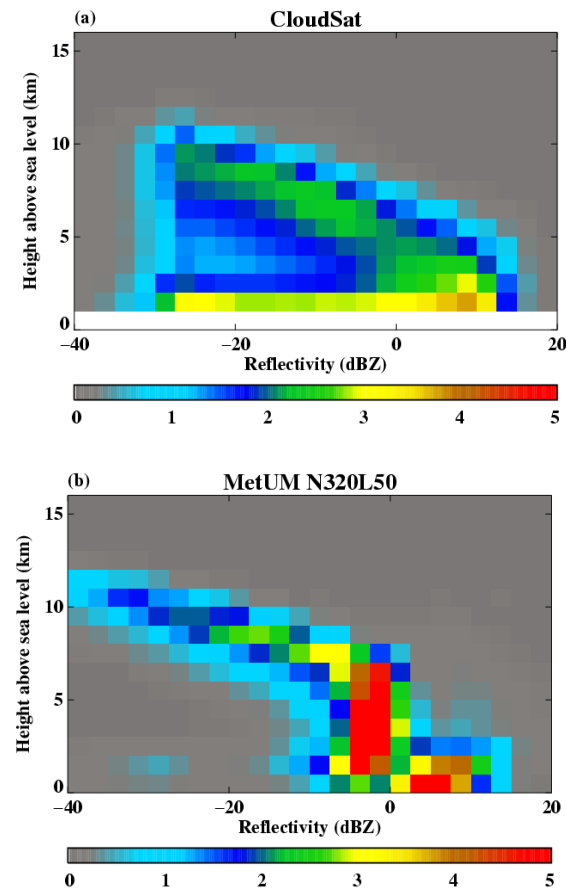
$$Z_e \propto \frac{q_h^a}{N^b} f_{Mie}(D_e) e^{-2 \int_0^z \sigma_e(z) dz}$$

Is the case study representative of the North Atlantic?

Case Study

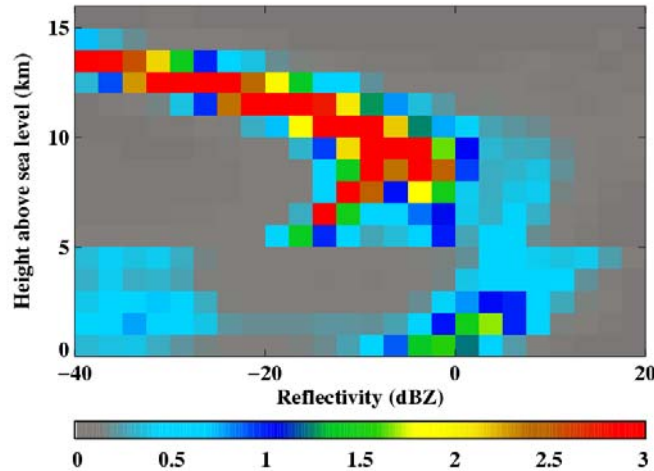


North Atlantic

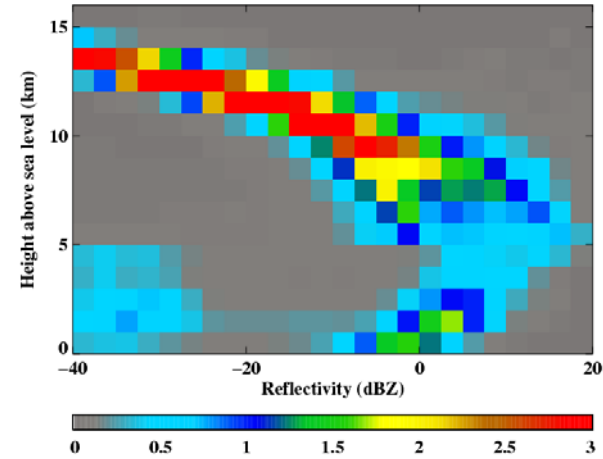


Flexibility of forward modelling: sensitivity tests

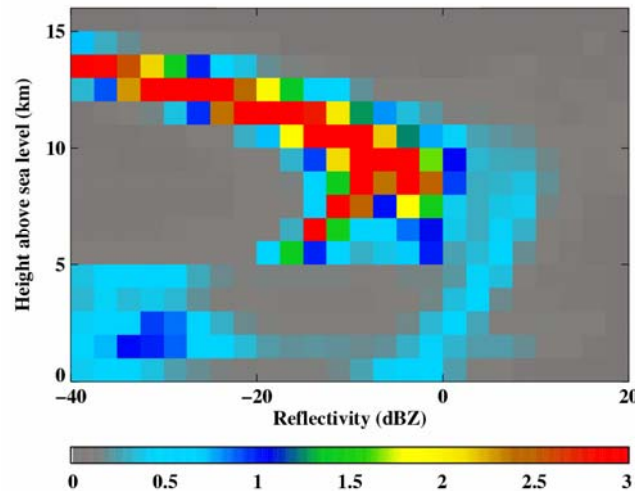
Control



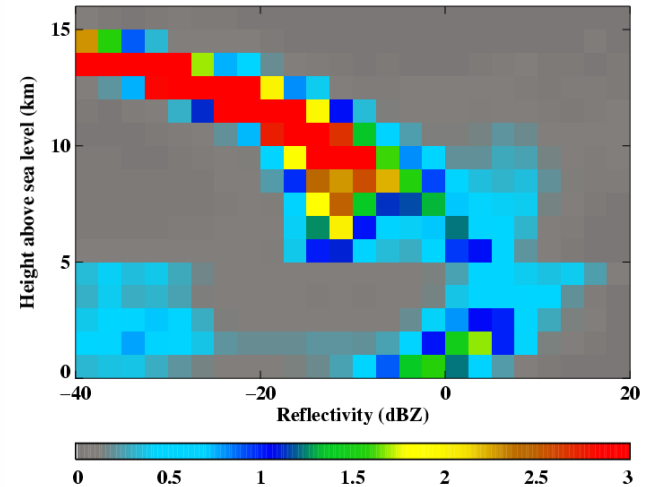
Rayleigh



No large-scale rainfall

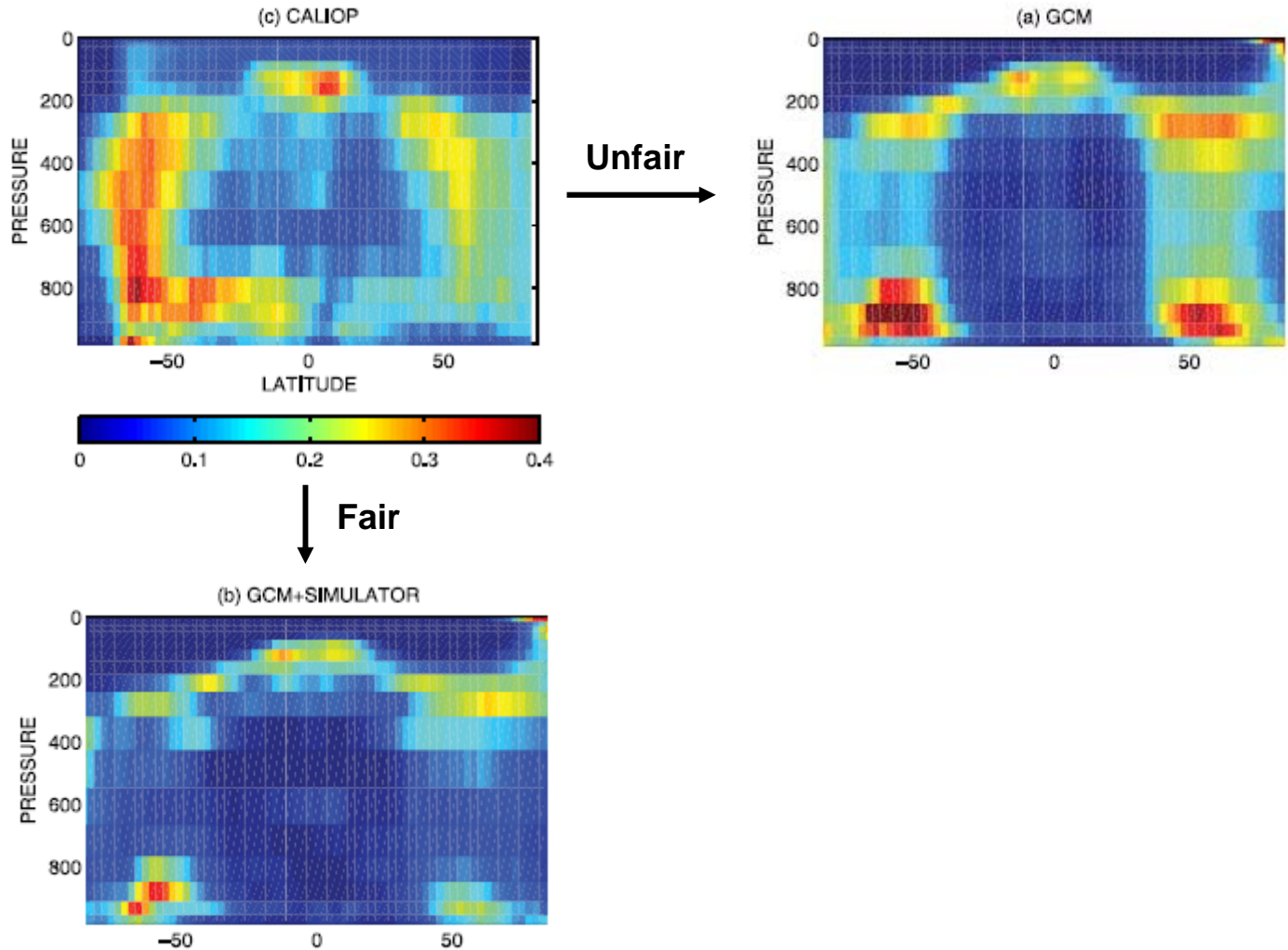


Hogan et al., 2006



Allowing for fair comparisons

LMDZ4 v CALIOP/CALIPSO





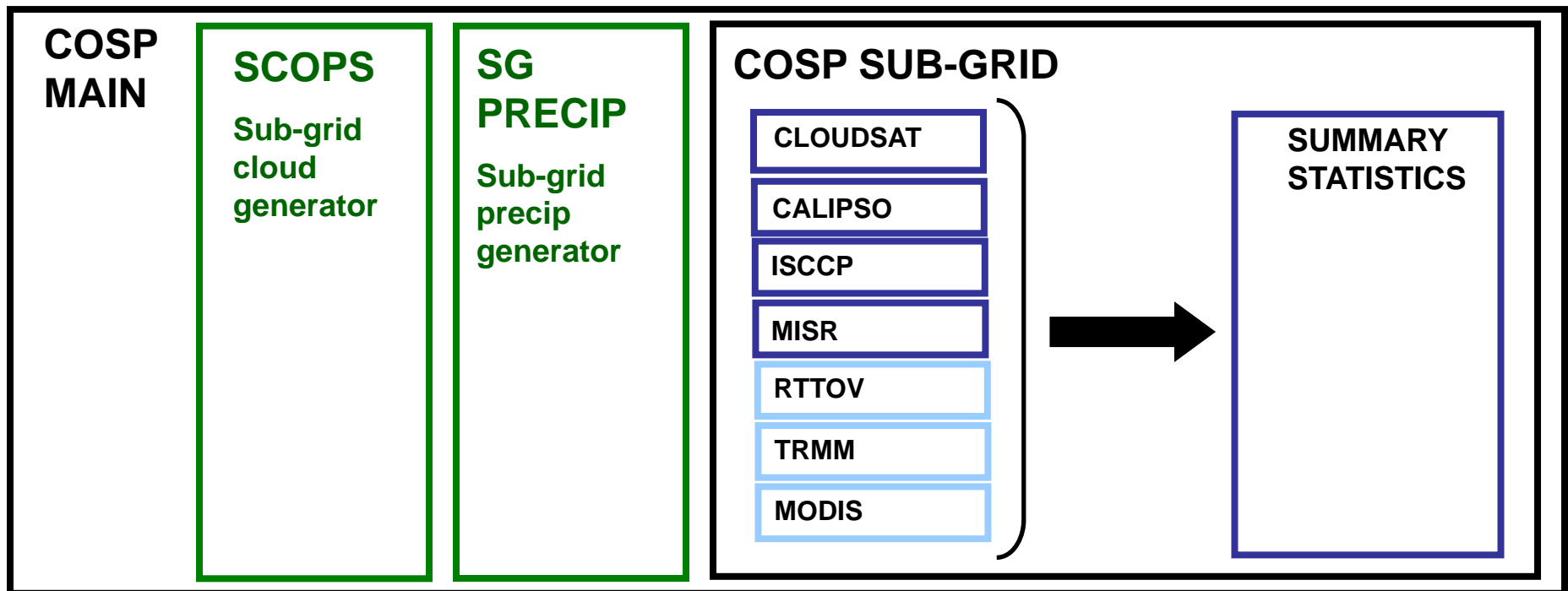
An integrated approach: the CFMIP Observational Simulator Package (COSP)



COSP collaborators

- **Met Office:** M. J. Webb
- **LMD/IPSL:** S. Bony, H. Chepfer, J.-L. Dufresne
- **PCMDI/LLNL:** S. Klein, Y. Zhang
- **U. Washington:** R. Marchand
- **Monash U.:** J. Haynes

- New modules can be plugged in without the need of interfacing
- Used in the Couple Model Intercomparison Project (experiments for next IPCC report)
- Designed to be used by climate, NWP and cloud-resolving models





Output diagnostics

CloudSat

- Radar reflectivity in each subcolumn
- Height-reflectivity histograms

CALIPSO

- Lidar total backscatter (532 nm)
- Lidar molecular backscatter
- Height-scattering ratio histograms
- Low-level cloud fraction (CTP > 680 hPa)
- Mid-level cloud fraction (440 < CTP < 680 hPa)
- High-level cloud fraction (CTP < 440 hPa)
- 3D Cloud fraction
- Total cloud fraction

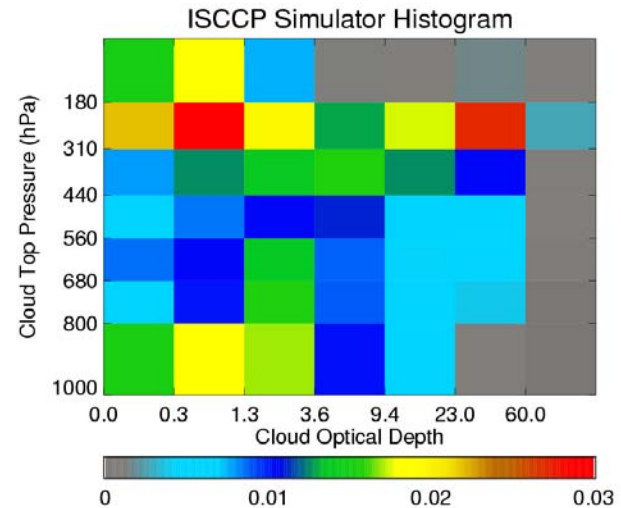
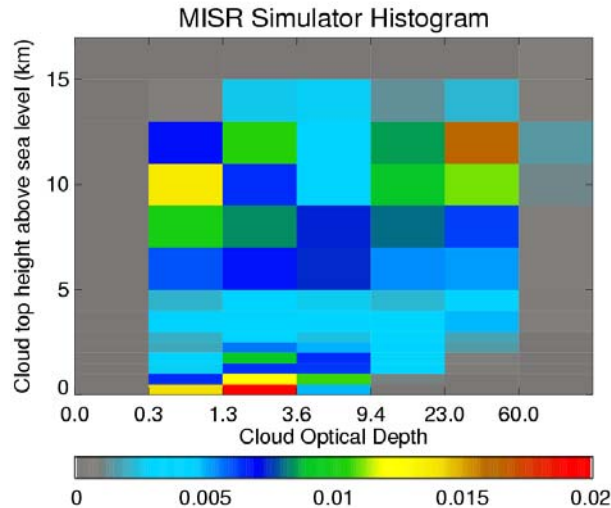
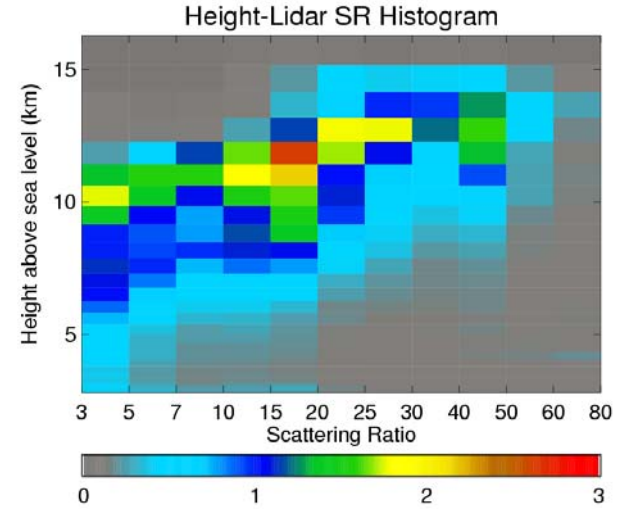
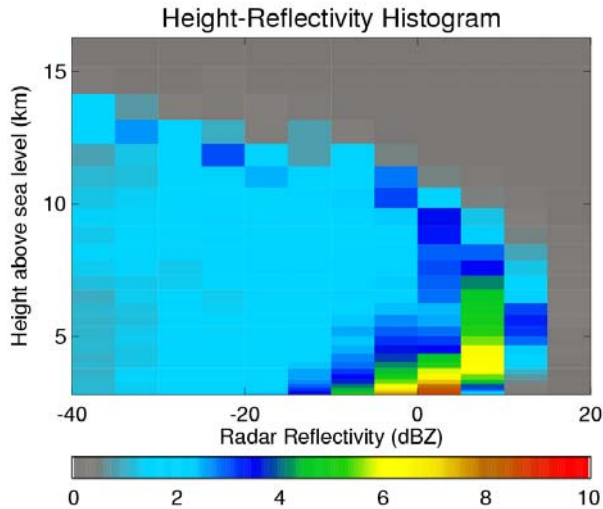
MISR, PARASOL and combined

- PARASOL mono-directional reflectance
- MISR CTH-Tau histograms
- Total cloud fraction from CALIPSO & CloudSat
- 3D cloud fraction as seen from CALIPSO but not CloudSat

ISCCP

- Mean cloud albedo
- Cloud optical depth in each subcolumn
- Mean cloud top pressure
- Mean 10.5 micron brightness temperature
- Mean clear-sky 10.5 micron brightness temperature
- Mean cloud optical depth
- Cloud top pressure in each subcolumn
- CTP-tau histograms
- Total cloud fraction

Examples



Key messages

- **Use of satellite retrievals introduce uncertainties in model evaluation**
- **Forward modelling avoids some of these ambiguities**
 - **Real observations, bring models and observations onto the same grounds**
 - **Sensitivity tests can be easily carried out**
 - **But: Interpretation in observations space**
- **The A-train provides a new dimension in clouds and precipitation**
- **COSP as an integrated approach to exploit forward modelling in model evaluation**

Future

- **Understand the uncertainties in forward modelling**
- **Exploit this approach in model intercomparisons**
- **Implement new forward models in COSP**
- **New synergistic diagnostics**



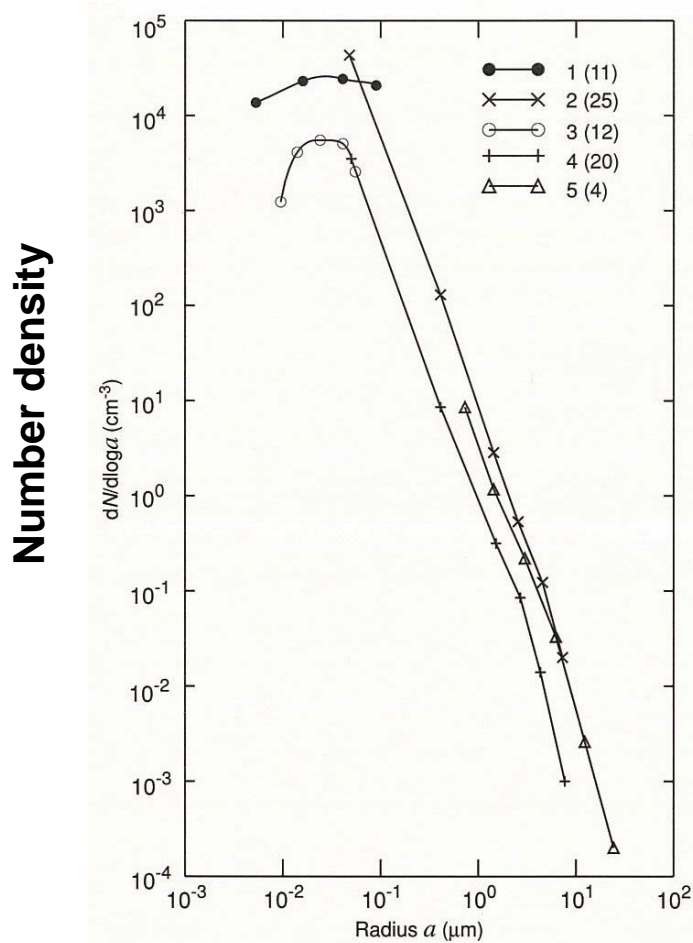
Met Office
Hadley Centre



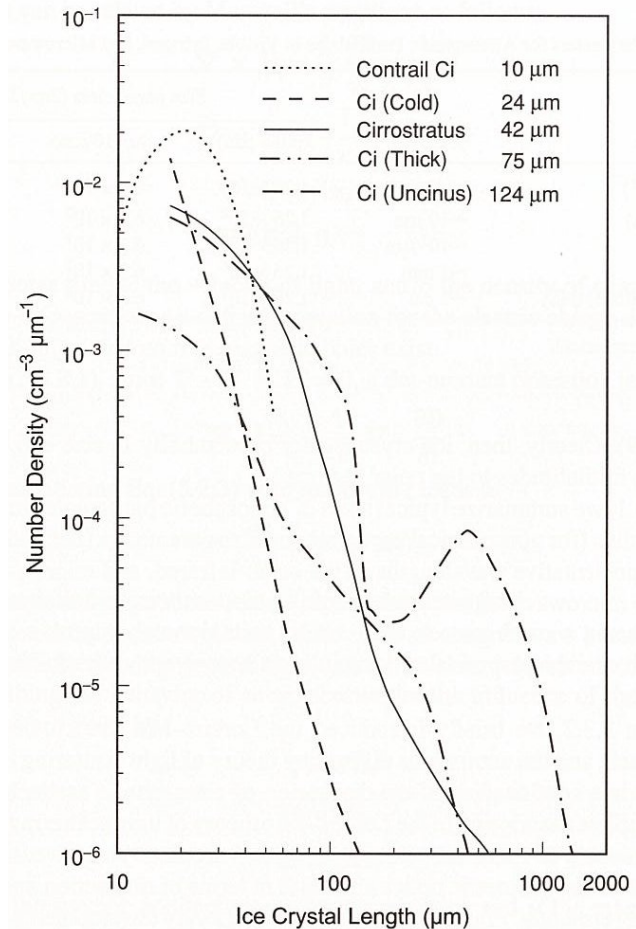
Thanks!

Size of atmospheric particles

Aerosols



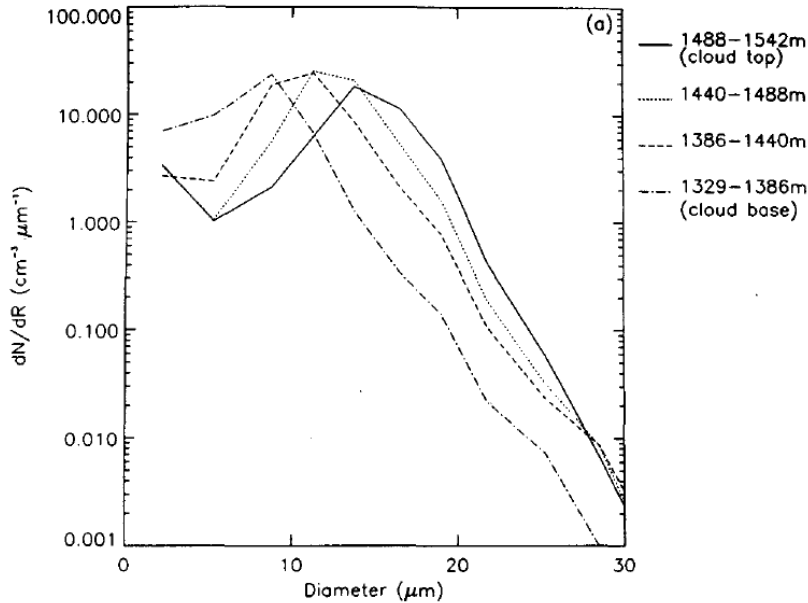
Ice crystals



Particle size

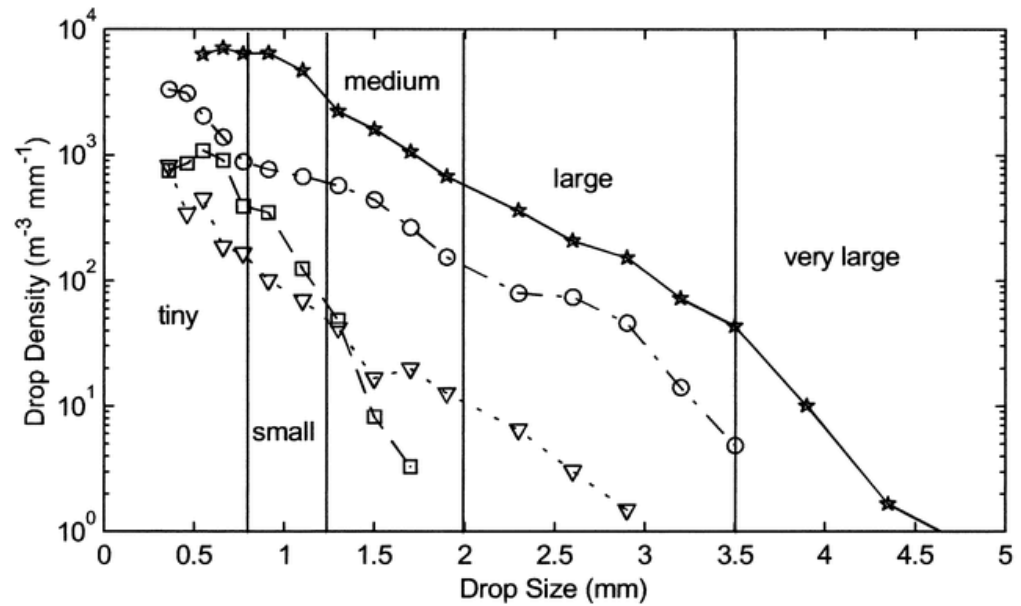
Size of atmospheric particles

Liquid cloud droplets



(Martin et al., *JAS*, 51, 1823-1842, 1994)

Rain drops

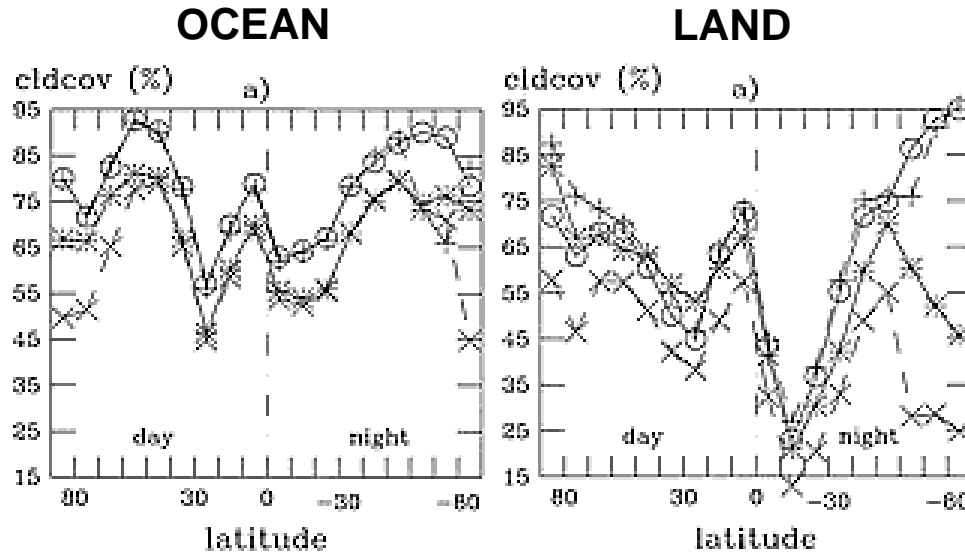


(Nystuen, *JAOT*, 18(10), 2001.
DOI: 10.1175/1520-0426(2001)018<1640:LTRFUA>2.0.CO;2)



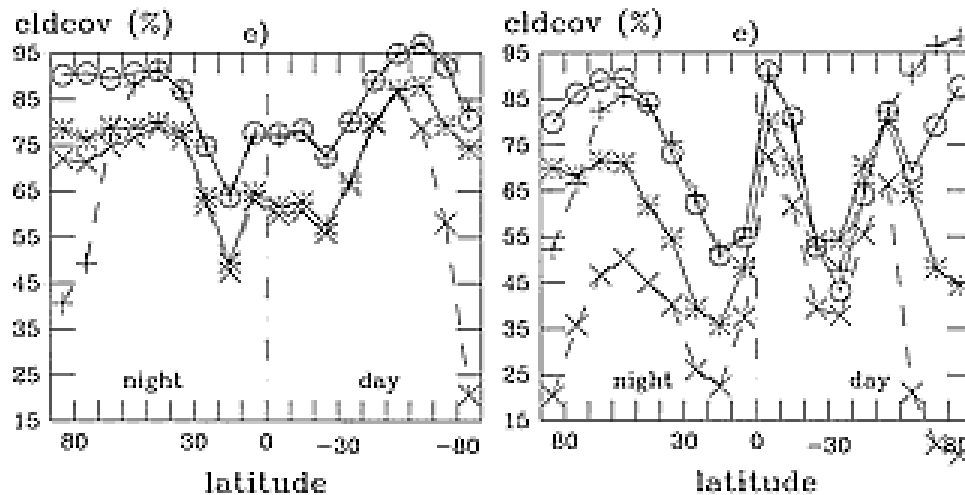
Met Office
Hadley Centre

And the obs?



○ 3Inew + 3Iold * DX × CX

Jul 1987 7:30am



Jan 1988 7:30am

T_B from TOVS

ECMWF

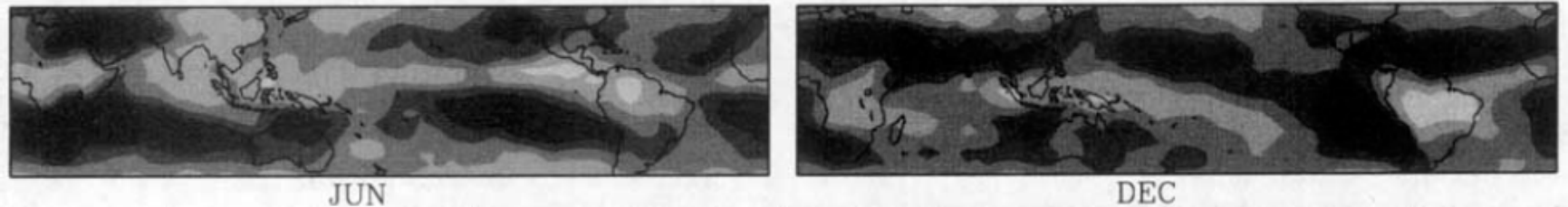


FIG. 3. Monthly averages of TOVS brightness temperature T^* computed from ECMWF analyses of tropospheric temperature and moisture in 1989. The same gray scale is used for comparison with the corresponding TOVS-observed radiances presented in Fig. 1, where the dry regions are more intense than in these simulated radiances.

TOVS

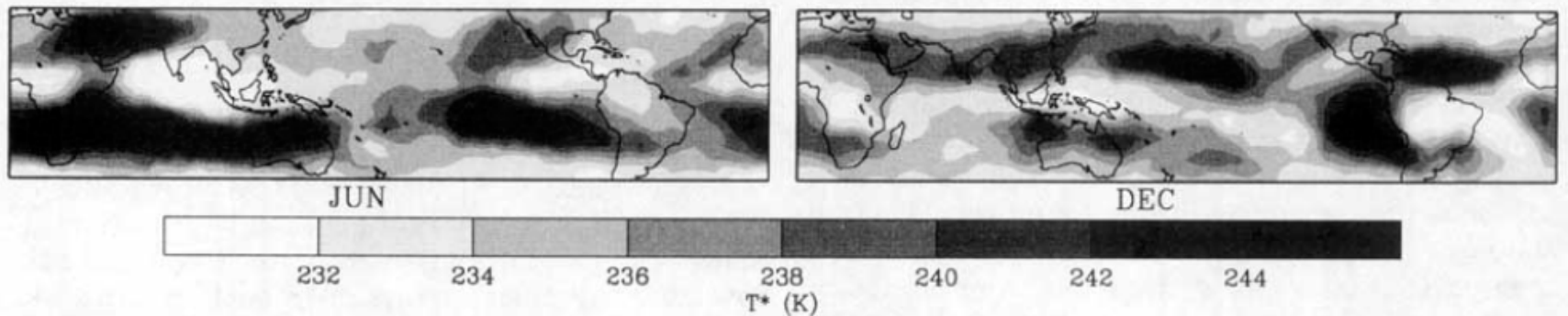
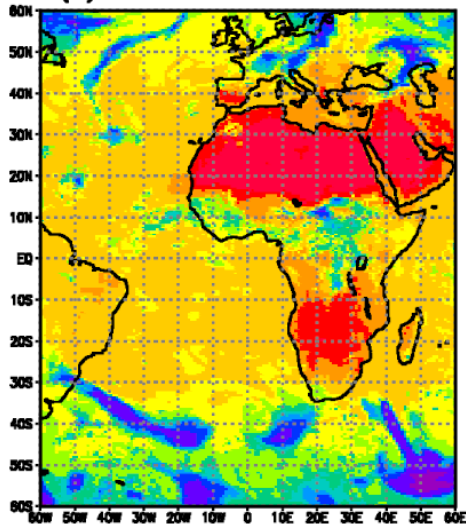


FIG. 2. Monthly averaged fields of TOVS brightness temperature T^* observed by the NOAA weather satellites during 1989. Darker shades indicate the higher brightness temperatures observed where persistent dry regions allow heat to escape from deeper in the troposphere, principally in the winter subtropical regions.

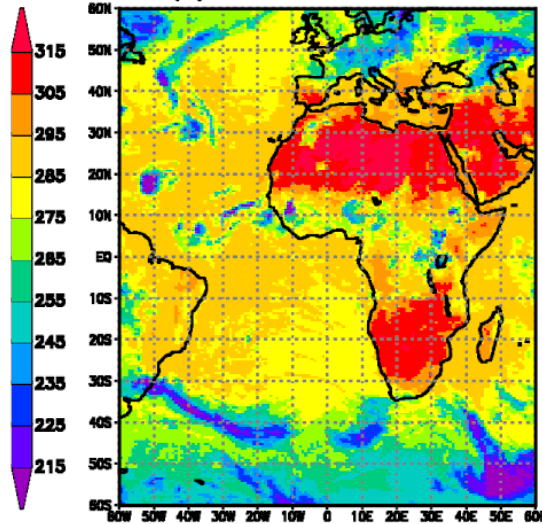


IRW&WV from METEOSAT

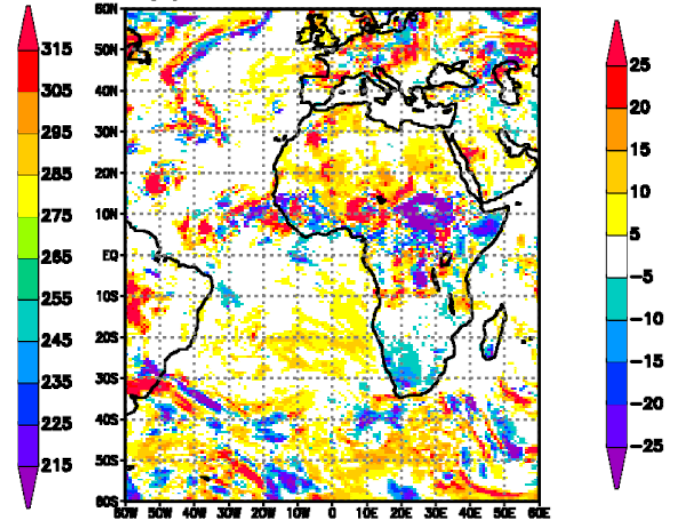
(a) IR Window:Unified Model



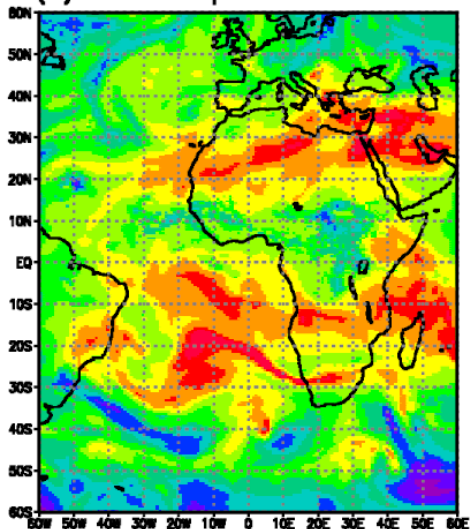
(b) IR Window:Meteosat



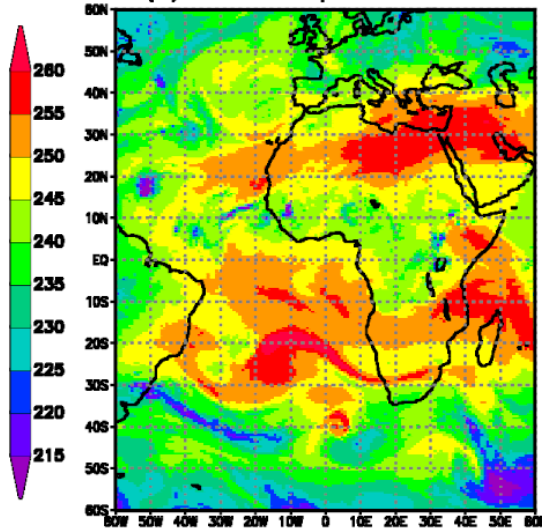
(c) IR Window:Model - Obs



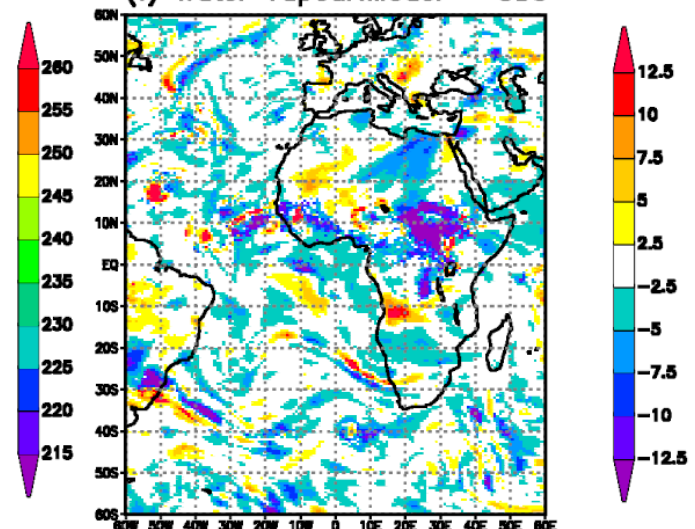
(d) Water Vapour:Unified Model



(e) Water Vapour:Meteosat

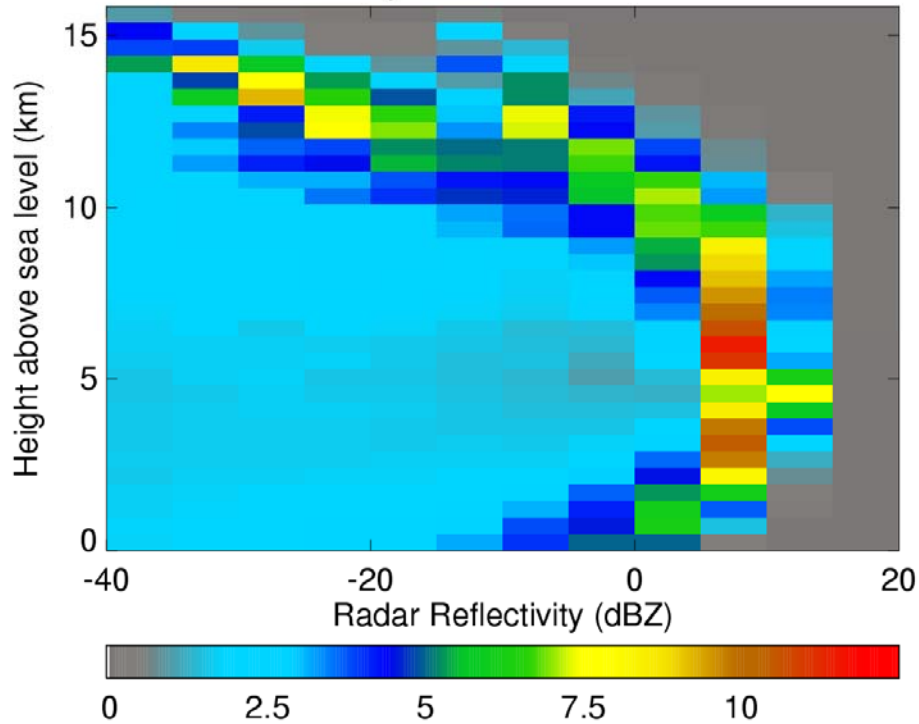


(f) Water Vapour:Model - Obs

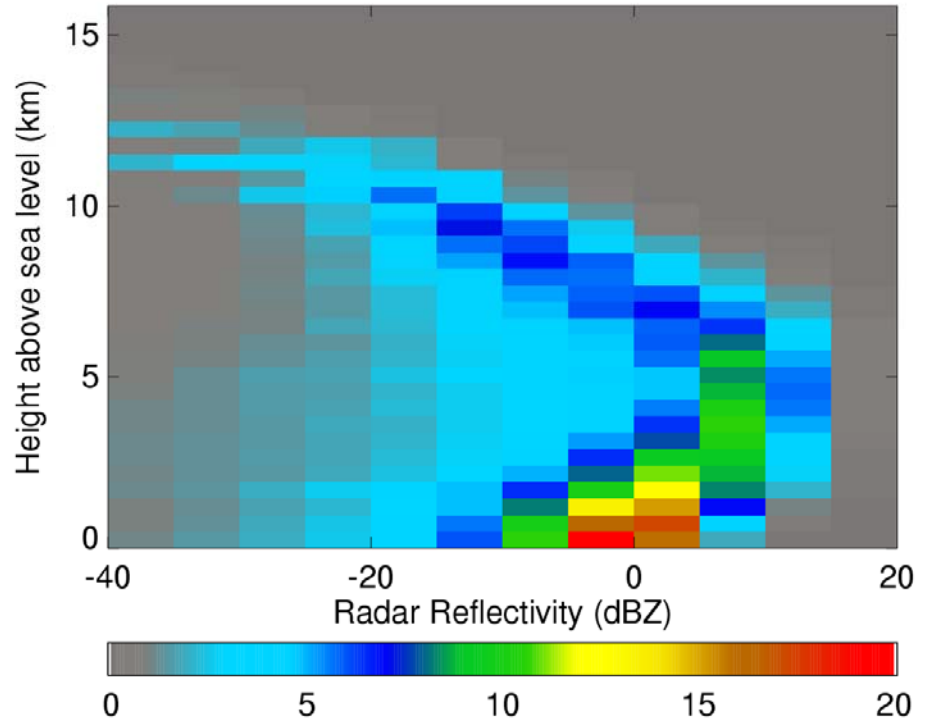


Application to HadGAM1

Tropical Warm Pool

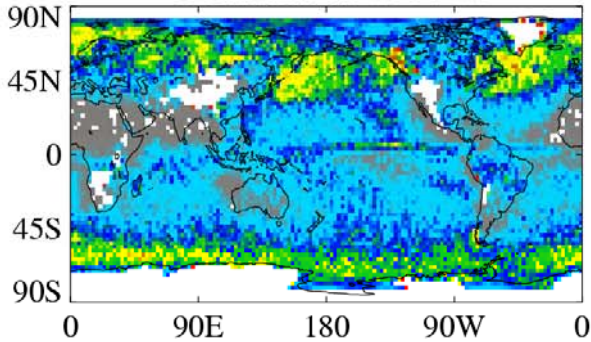


North Atlantic

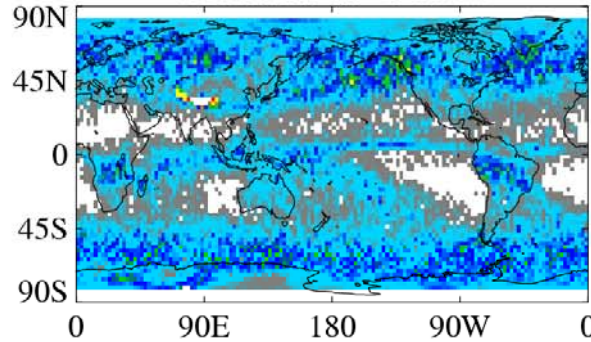


Hydrometeor frequency of occurrence

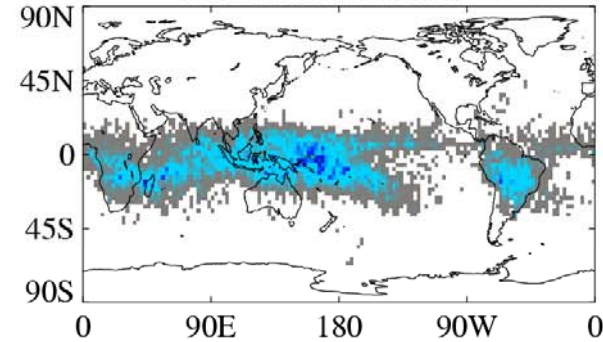
CloudSat at 1500 m



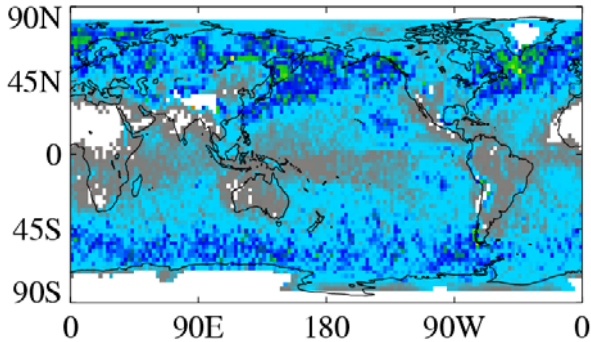
CloudSat at 4500 m



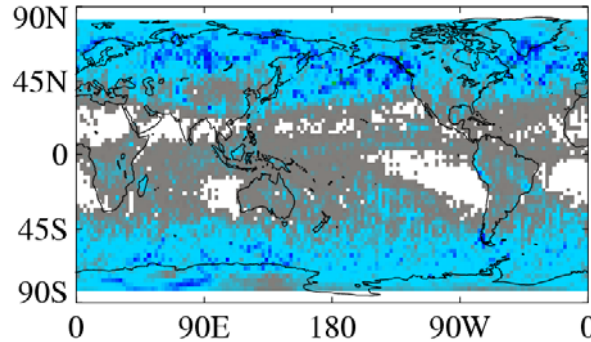
CloudSat at 13500 m



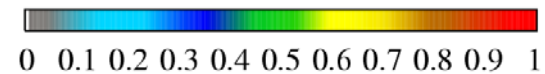
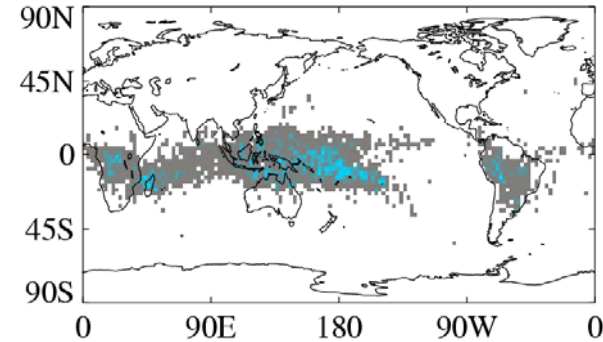
MetUM N320L50 at 1500 m



MetUM N320L50 at 4500 m



MetUM N320L50 at 13500 m

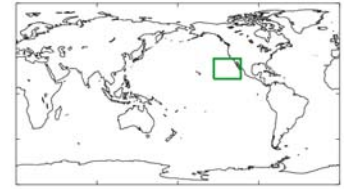




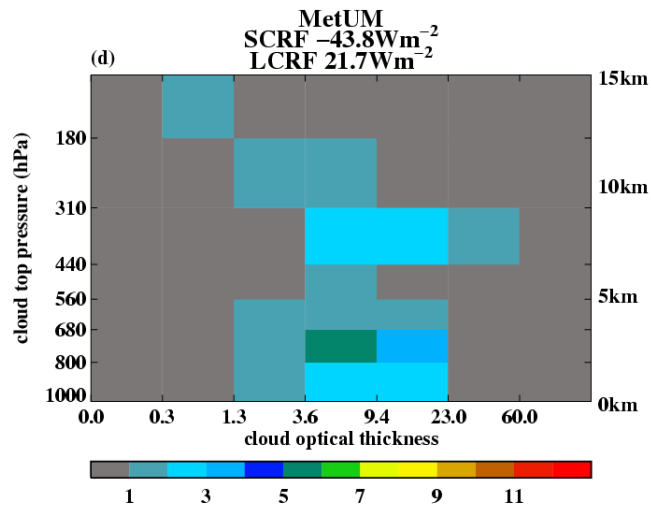
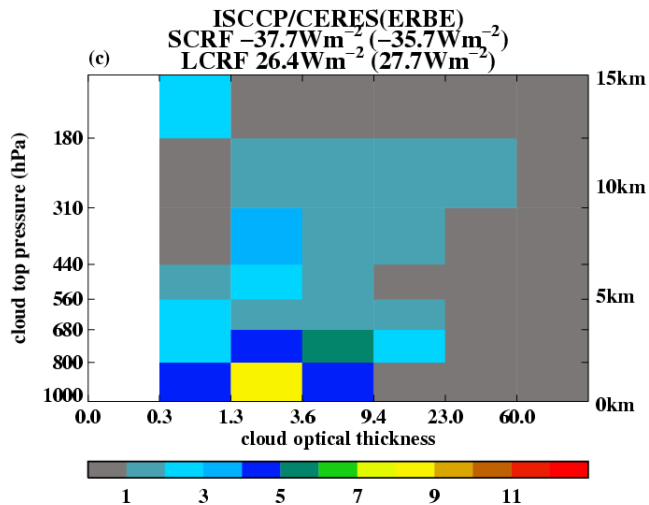
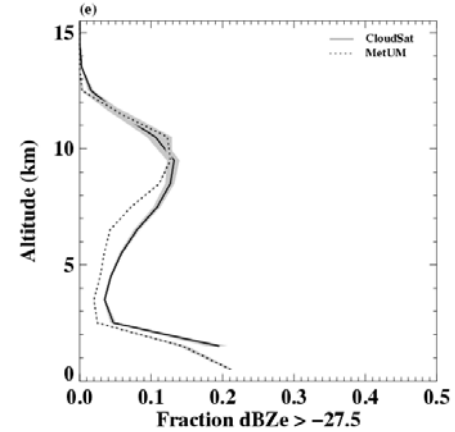
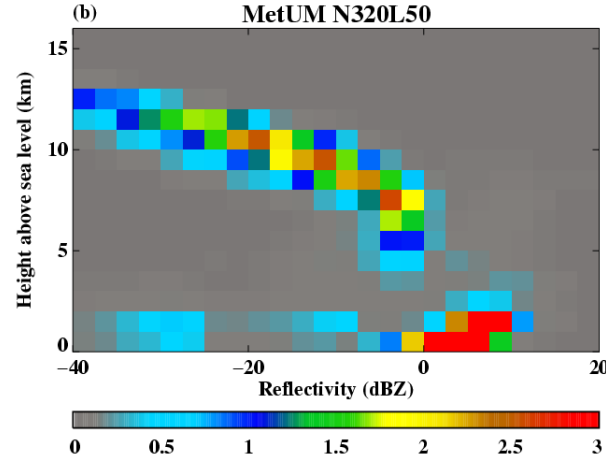
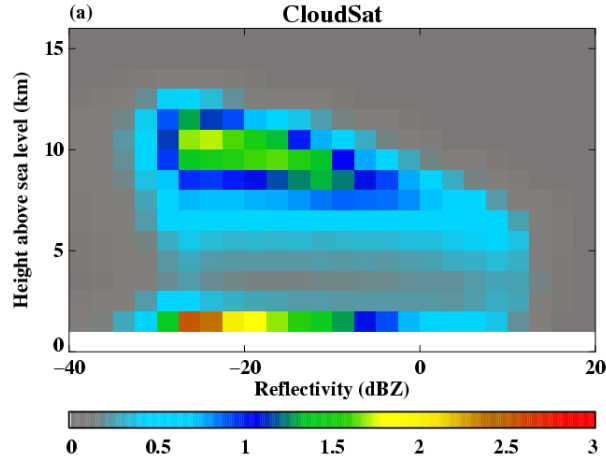
Californian Sc

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Hadley Centre

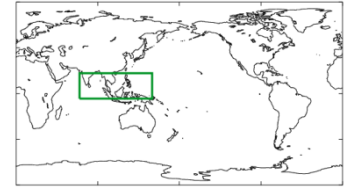
Observations



MetUM



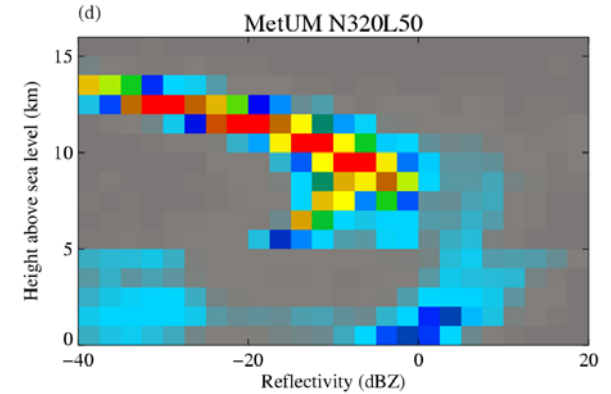
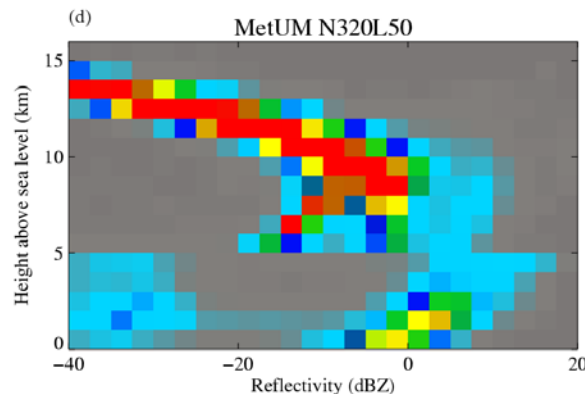
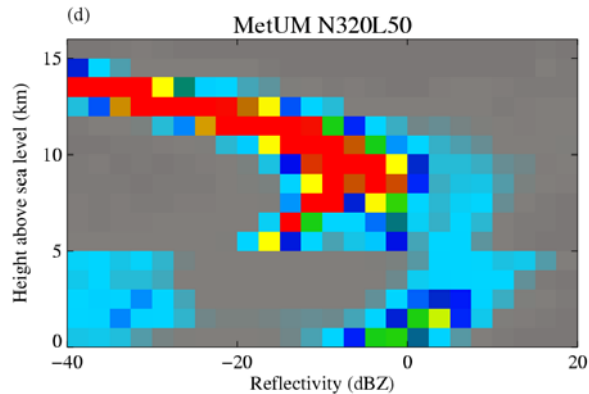
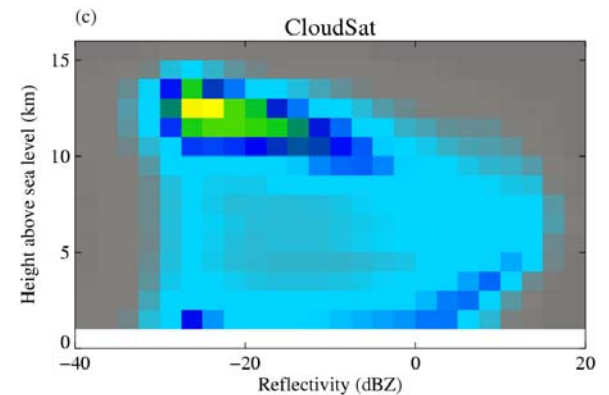
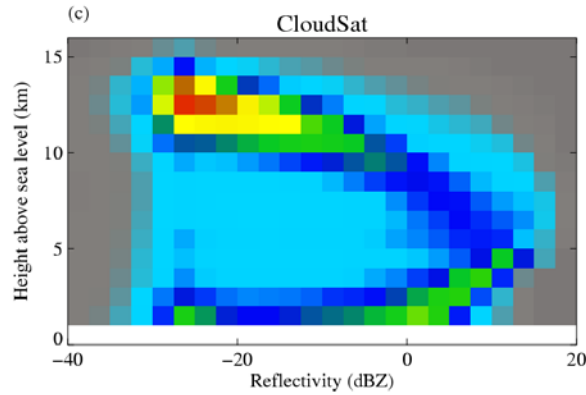
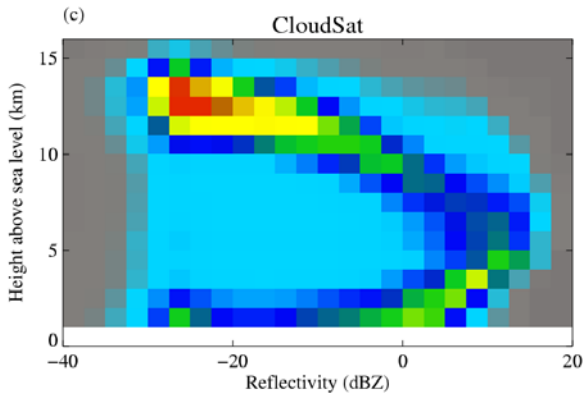
Impact of sampling



December

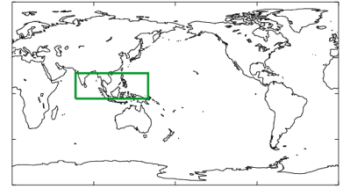
January

February

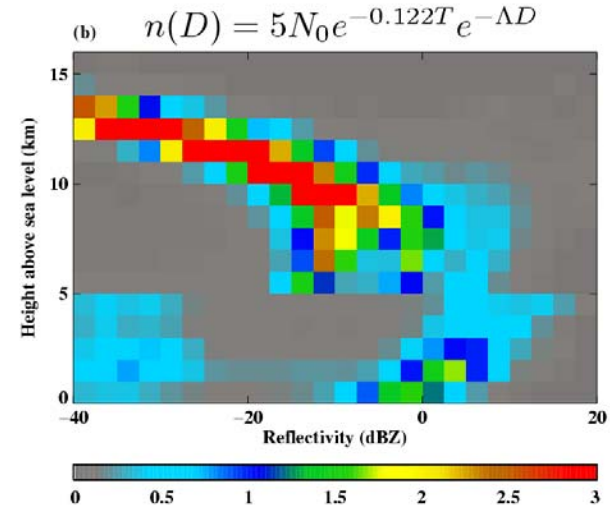
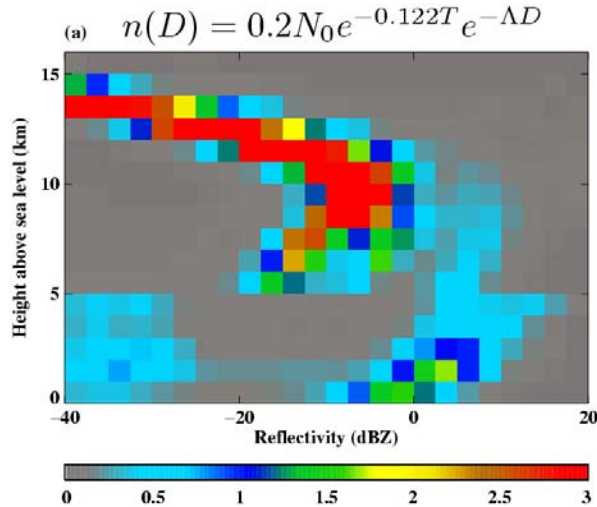


Sensitivity to PSD

$$n(D) = N_0 e^{-C_0 T} e^{-\Lambda D}$$

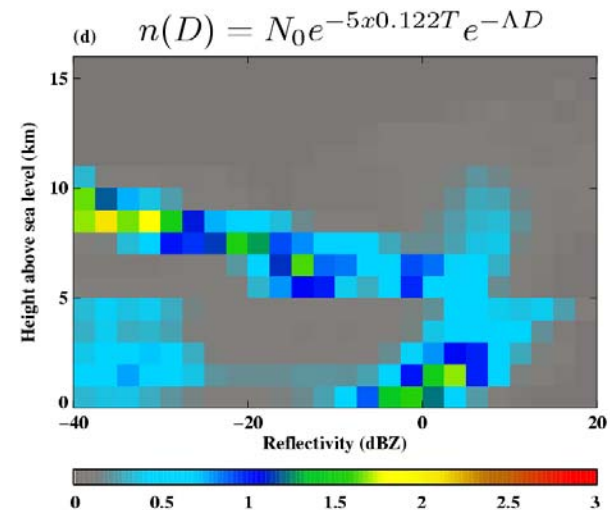
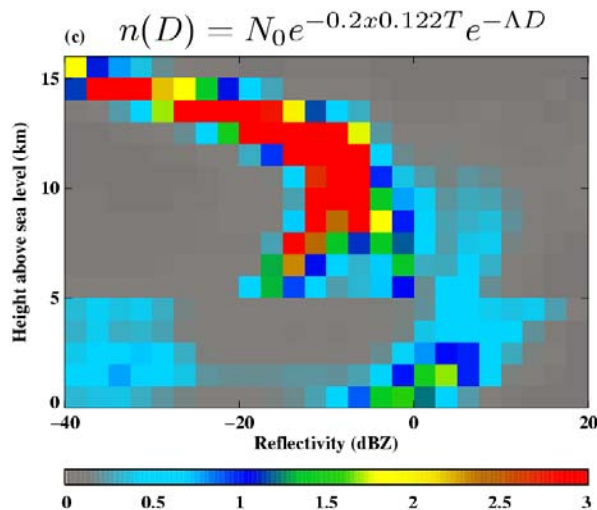


$N_0/5$



$5 \times N_0$

$C_0/5$



$5 \times C_0$

Lidar backscatter

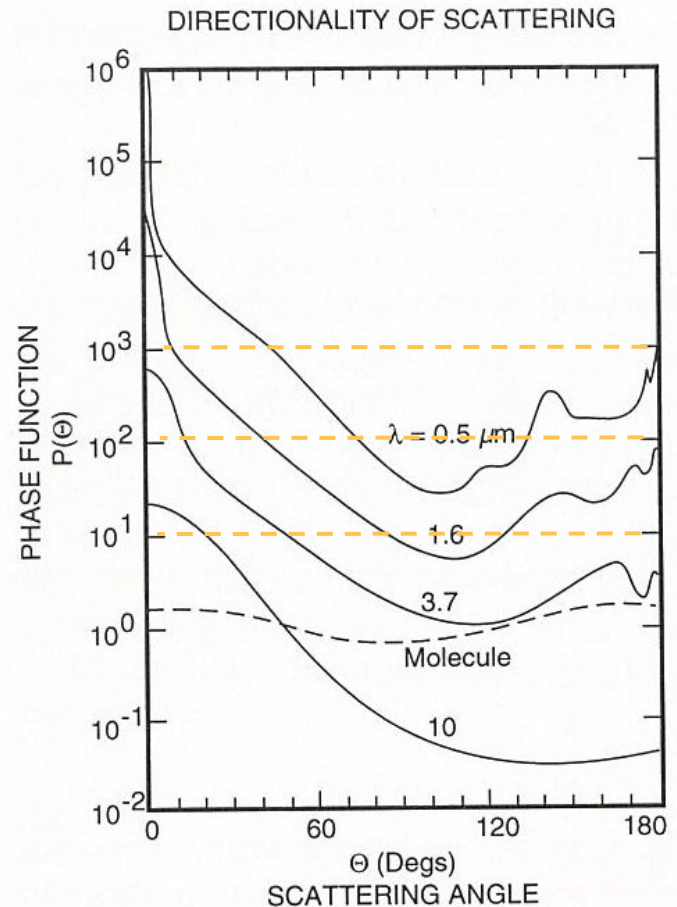
$$ATB(z) = \beta_s(z) e^{-2\alpha} \int_0^z \sigma_s(z) dz$$

$$\beta_s = \sigma_s \frac{P(\pi)}{4\pi}$$

$$\int_{\Omega} \frac{P(\Theta)}{4\pi} d\Omega = 1$$

$$\sigma_s(z) = \int \pi r^2 Q(r) n(r; z) dr$$

(Chepfer et al., *GRL*, 35, 2008. DOI:10.1029/2008GL034207)



(Salby, *Fundamentals of Atmos. Physics*, 1996. Fig 9.27)



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Radar reflectivity

$$\eta = \int \beta(D)n(D)dD$$

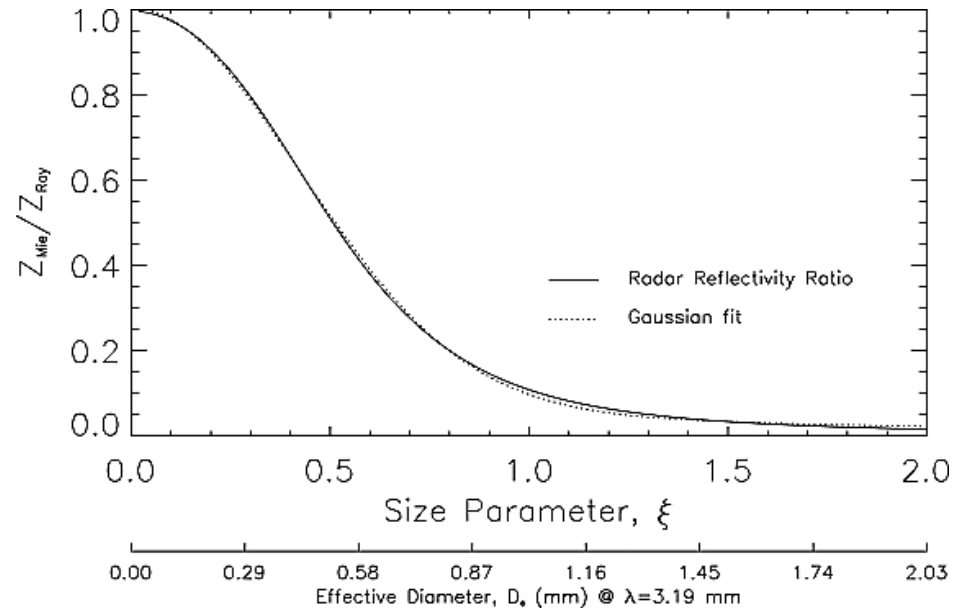
$$Z_e = \frac{|K(\nu, T)|^2}{|K_w|^2} Z f_{Mie}(D_e) e^{-2 \int_0^z \sigma_e(z) dz}$$

$$\beta(D) = \frac{\pi^5}{\lambda^4} |K|^2 D^6$$

$$\eta = \frac{\pi^5}{\lambda^4} |K|^2 \int D^6 n(D) dD$$

$$Z = \int D^6 n(D) dD$$

$$Z_e = \frac{\lambda^4}{\pi^5 |K_w|^2} \eta$$



(Benedetti et al., *JGR*, 108, 2003. DOI:10.1029/2002JD002693)



COSP v2.0

- **New instruments**
 - **RTTOV. Fast RT code for IR and PMW (Optional)**
 - **MODIS. Daytime only**
 - **CF, CTP, size, and OT**
 - **All, liquid and ice**
 - **Joint CTP/OT histogram**
 - **Matching Level3 product to be produced by NASA**
 - **TRMM PR sf-conv precip**
 - **Others...**
- **Compatibility with CMOR 2**
- **New diagnostics**
- **Expected to be released by Jan/Feb 2009**