

# Diagnosis of Ensemble Forecasting Systems

M. Leutbecher



September 2009

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**mean**  $\longleftrightarrow$  **variance**

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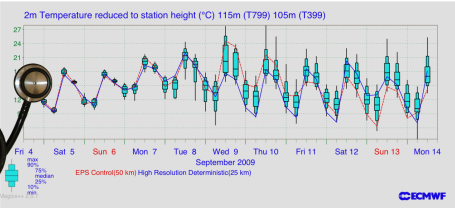
Acknowledgements: Renate Hagedorn, Hans Hersbach, Linus Magnusson, . . . ,  
Thorpex PDP IG5: O. Talagrand, F. Atger, . . .

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- 3 Assessing the spatio-temporal variation of the pdf-shape
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- 4 Diagnosis & Numerical Experimentation
- 5 Conclusions

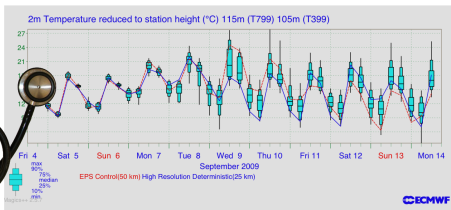
# Diagnosis of ensemble forecast systems



Why?



# Diagnosis of ensemble forecast systems

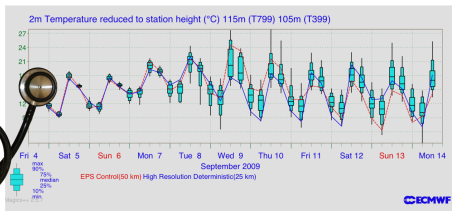


Why?

- Aid forecast system development:

- ▶ Quantify meteorologically relevant differences between different forecast systems.
- ▶ Identify deficiencies
- ▶ Provide guidance for refining the representation of initial uncertainty and model uncertainty

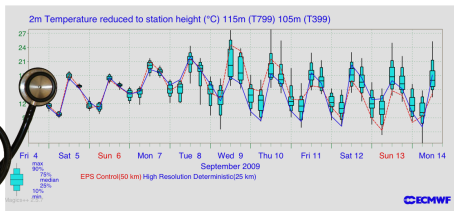
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  - ▶ Examine origin of large forecast errors

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Limitations: not exhaustive, not only new developments, some of the new things are work in progress



# Diagnosis of the numerical model

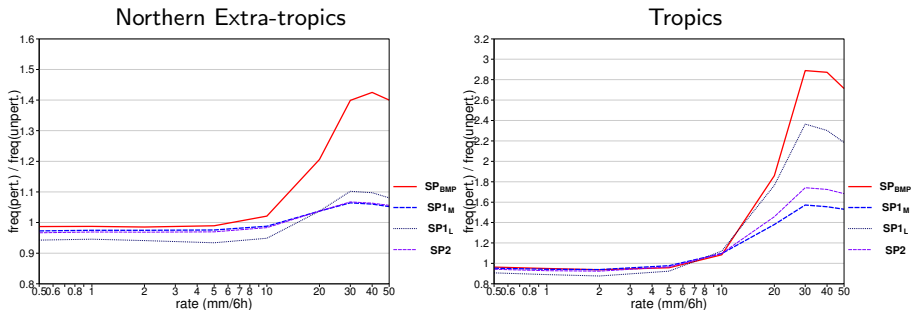
used in the ensemble forecast system

- ensemble forecast model  $\neq$  model used for “deterministic” forecast: resolution, timestep, ...
- look at performance of the control forecast (unperturbed member of ensemble)
- realism of model climate of perturbed forecast model (including impact of model perturbations)

Everything as would be done for the deterministic system (except for the model perturbations).

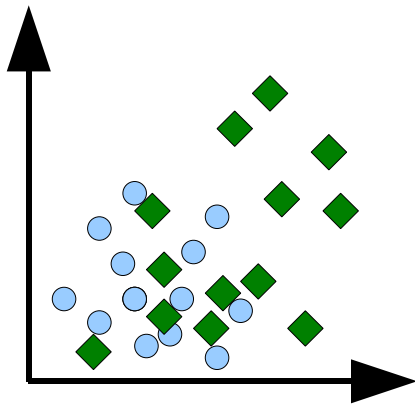
# Getting the climate right: An example

- ECMWF EPS uses Stochastically Perturbed Parametrization Tendencies (SPPT) (“stochastic physics”)
- **Operational SPPT** ( $\leq 35R2$ ) distorts the tail of the climatological distribution of precipitation.
- A recent **major revision of SPPT** has improved precipitation distribution

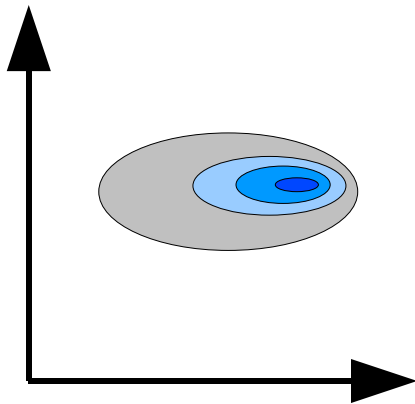


Precipitation frequency ratios between forecasts using tendency perturbations and forecasts without tendency perturbations. — **operational SPPT** - - **revised SPPT**

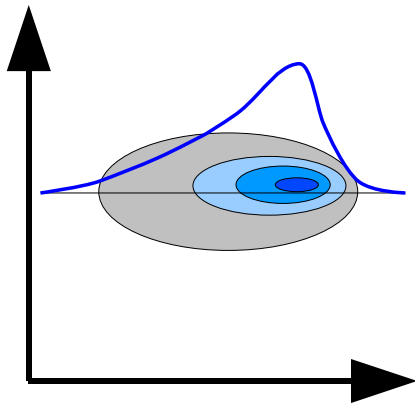
# Diagnosis of deterministic forecasts



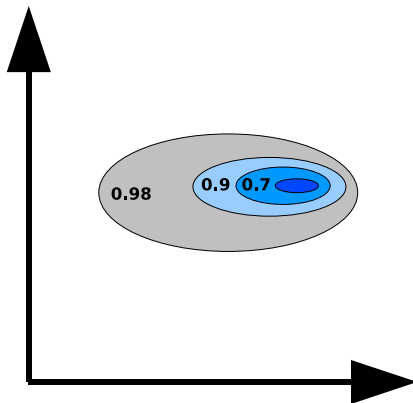
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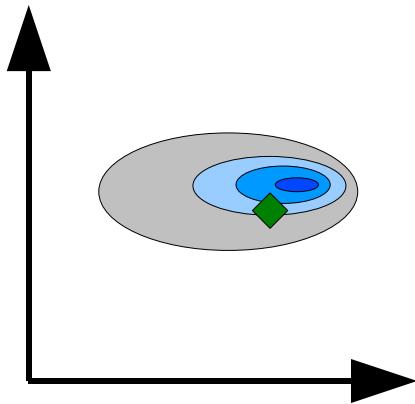
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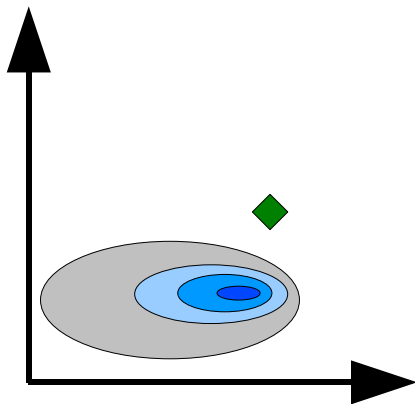
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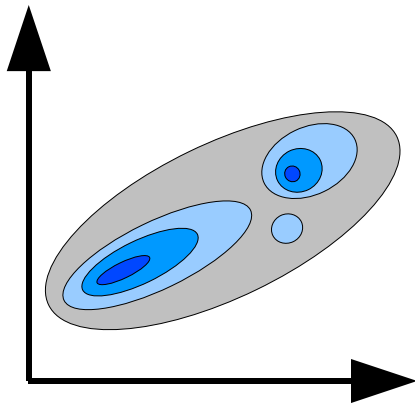
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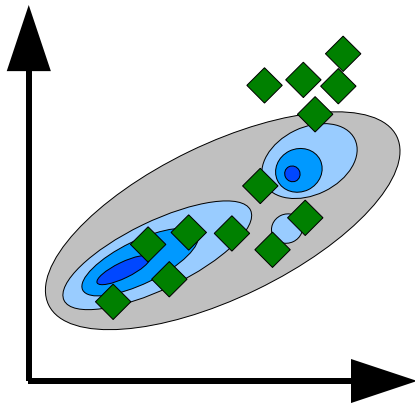
- conclusions for single cases only for exceptional failures
- forecast and verification are different objects



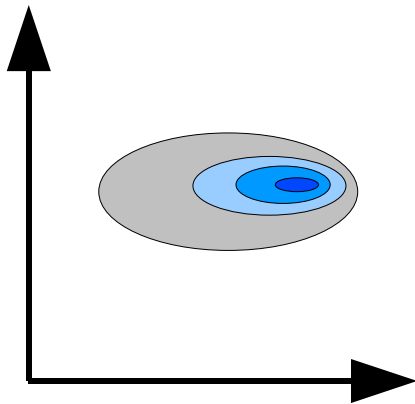
# Diagnosis of probabilistic forecasts: Reliability



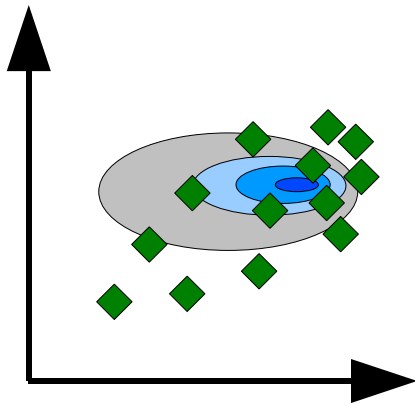
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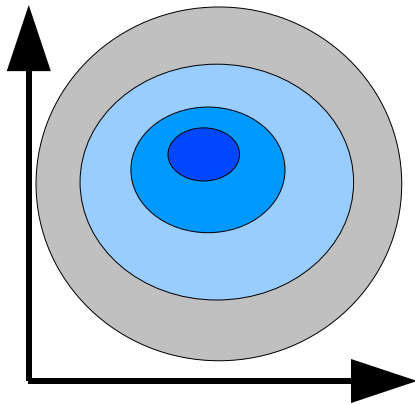
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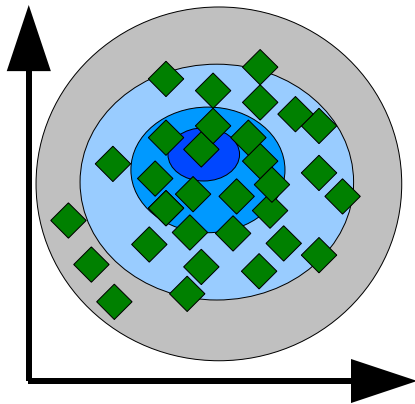
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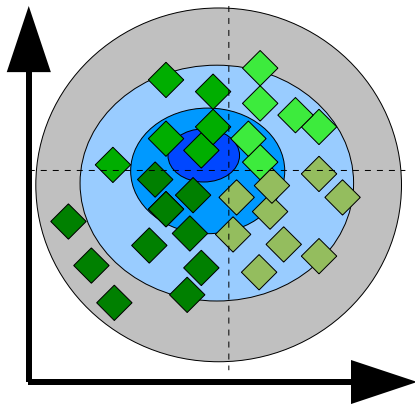
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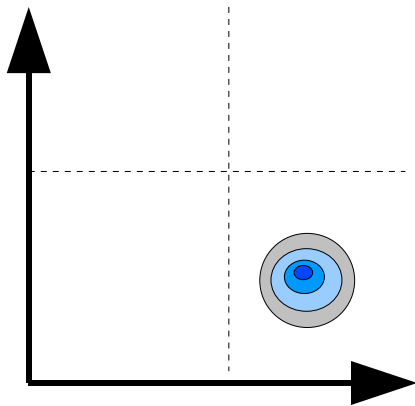
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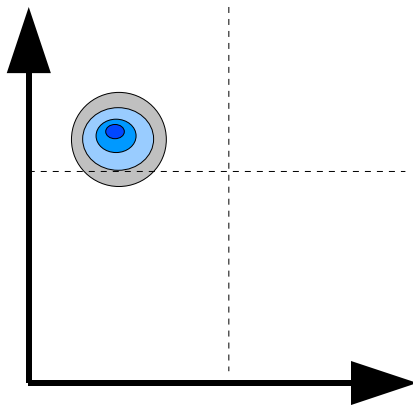


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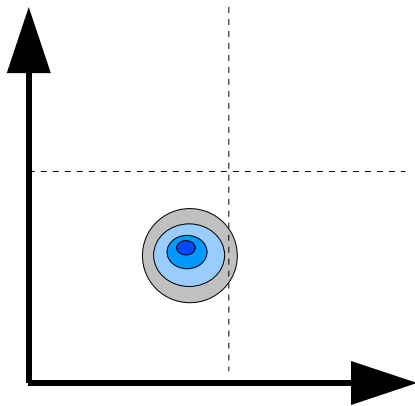




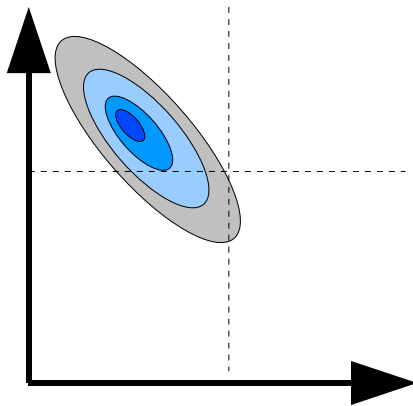
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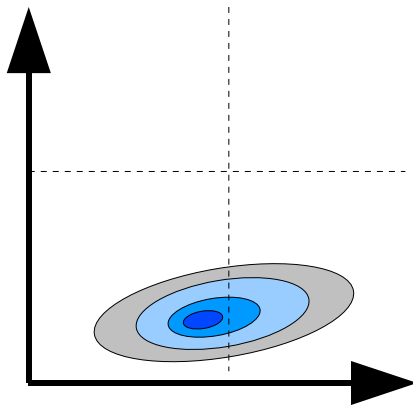
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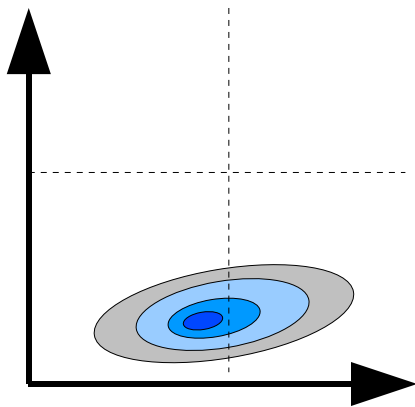
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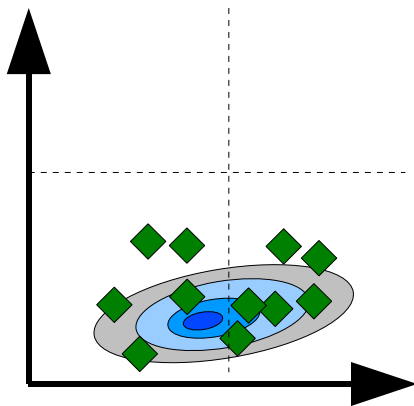
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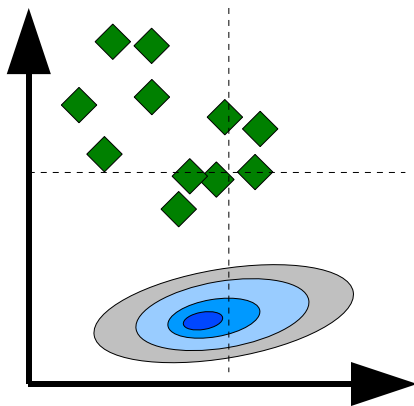
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# Standard methods to diagnose the pdf

Impractical to assess all aspects of a multivariate probabilistic prediction

How can the assessment be simplified?

- ① Limit assessment of probability distribution:
  - ▶ univariate prediction: e.g. geopotential at 500 hPa
  - ▶ binary events: does TC strike at  $x$ ; prediction of a cold anomaly



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  - ▶ Rank Histogram: reliability
  - ▶ Brier score, (Continuous) Ranked Probability Score: reliability and resolution

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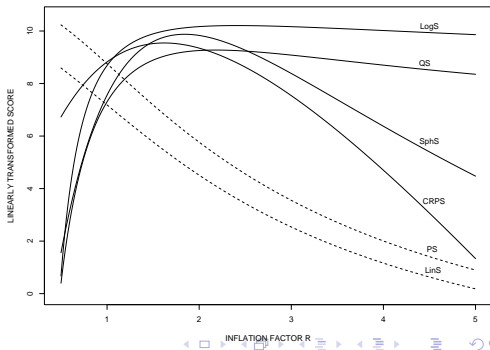
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  - ▶ Rank Histogram: reliability
  - ▶ Brier score, (Continuous) Ranked Probability Score: reliability and resolution (decomposition!)
  - ▶ Relative Operating Characteristic (ROC): (discrimination)
  - ▶ Logarithmic Score (Ignorance): reliability and resolution

# Proper scores

- **strictly proper** implies that optimizing the score leads to the correct probability distribution
- optimization of a score that is not proper is likely to lead to a wrong distribution
- concise mathematical definitions of *proper* and *strictly proper* are available (see Gneiting and Raftery, 2004)
- examples of proper scores: BS, RPS, CRPS, logarithmic score

Fig. 1,  
Gneiting and Raftery (2004)



## Spread-error relationship and ensemble size

Assume a perfectly reliable (statistically consistent)  $M$ -member ensemble: Ens. members  $x_j, j = 1, \dots, M$  and truth  $y$  are independent draws from a distribution with mean  $\mu$  and variance  $\sigma^2$ .

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Expected squared  
error of ensemble  
mean

$$\mathbb{E} \left( \underbrace{\frac{1}{M} \sum_{j=1}^M x_j}_{\mu + \text{sampling error}} - y \right)^2 = \left( 1 + \frac{1}{M} \right) \sigma^2$$

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Expected ensemble  
variance

$$\mathbb{E} \frac{1}{M} \sum_{j=1}^M \left( x_j - \underbrace{\frac{1}{M} \sum_{k=1}^M x_k}_{\neq \mu, \text{ contains } x_j} \right)^2 = \left( 1 - \frac{1}{M} \right) \sigma^2$$

# Spread-error relationship and ensemble size

Perfectly reliable M-member ensemble:

Ens. members  $x_j, j = 1, \dots, M$  and truth  $y$  are independent draws from a distribution with mean  $\mu$  and variance  $\sigma^2$ .

For large ensembles, e.g.  $M = 50$ ,

$$\overline{\text{ensemble variance}} = \overline{\text{squared ensemble mean error}}$$

in practice.



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For smaller ensemble size, e.g.  $M \leq 20$ ,

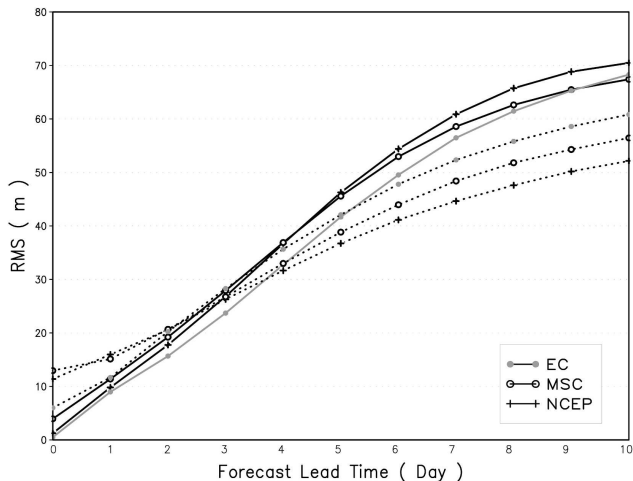
$$\left(1 - \frac{1}{M}\right)^{-1} \overline{\text{ens. variance}} = \left(1 + \frac{1}{M}\right)^{-1} \overline{\text{squared ens. mean error}}$$

# Spread versus error in 2002

**500 hPa  
geopotential height,  
N.-Hem.  
extra-tropics**

dashed: ens. stdev.  
solid: EM RMSE  
10 member ensembles

ECMWF( $T_{L255L40}$ ),  
MSC, NCEP

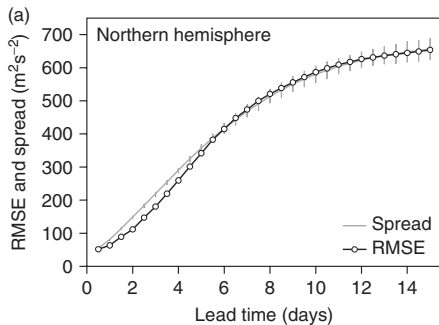


May–July 2002 *Fig. 5,*  
*Buizza et al. 2005*

FIG. 5. May–Jun–Jul 2002 average rms error of the ensemble mean (solid lines) and ensemble standard deviation (dotted lines) of the EC-EPS (gray lines with full circles), the MSC-EPS (black lines with open circles), and the NCEP-EPS (black lines with crosses). Values refer to the 500-hPa geopotential height over the Northern Hemisphere latitudinal band 20°–80°N.

# Spread versus error in 2007

500 hPa geopotential, N.-Hem. extra-tropics  
ECMWF, cycle 32r2



cycle 32r3

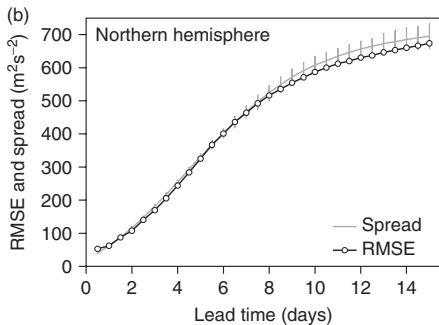


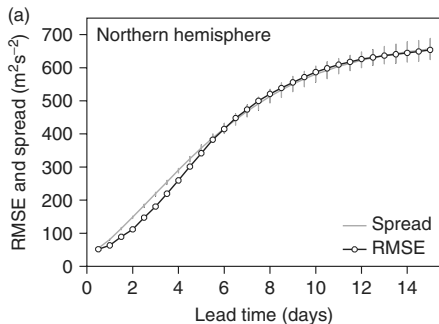
Fig. 12, Bechtold et al. 2008

50 member,  $T_L399L62$

69 cases, June-Sept. 2007

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500 hPa geopotential, N.-Hem. extra-tropics  
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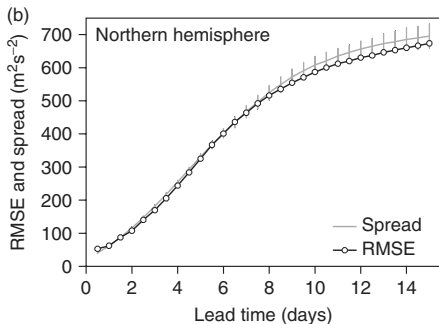


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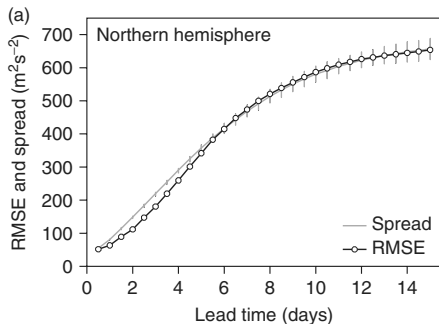
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*Improved match due to revised model physics together with a 30% reduction of the initial perturbation amplitude*

# Spread versus error in 2007

## 500 hPa geopotential, N.-Hem. extra-tropics ECMWF, cycle 32r2



## cycle 32r3

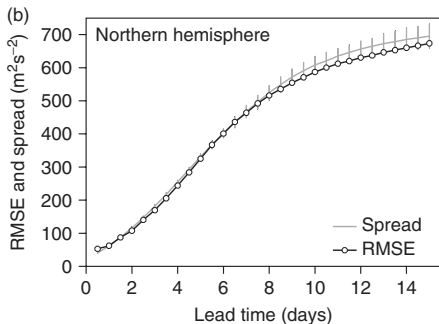


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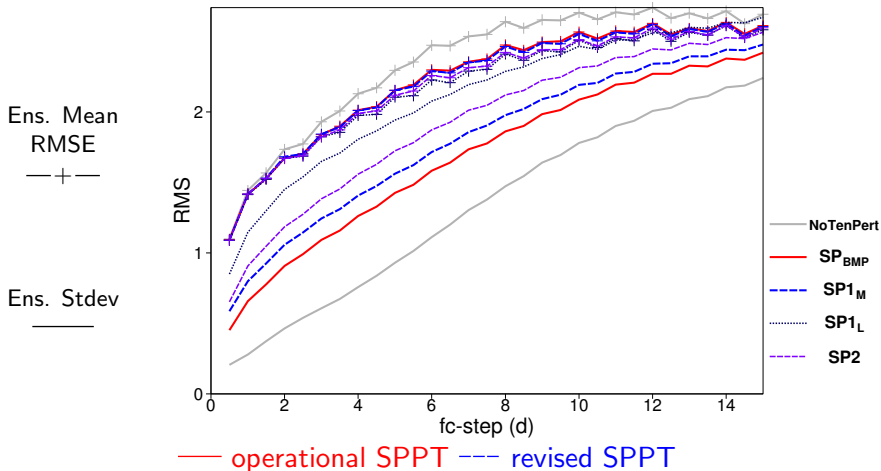
*Improved match due to revised model physics together with a 30% reduction of the initial perturbation amplitude* **Time to retire?**

## Explore more directions in phase space

- Agreement for other variables (e.g. T850)?
- Other regions (e.g. tropics)?
- Detailed geographical distribution of spread?
- Different spatial scales?

# Example: v850 tropics (20°S–20°N)

SPPT revision

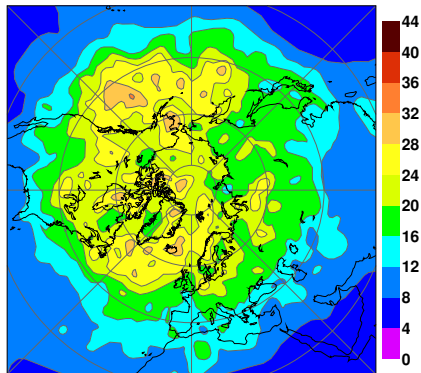


40 cases Nov/Dec 2007 + Jul/Aug 2008,  $T_L399(255$  from D+10)

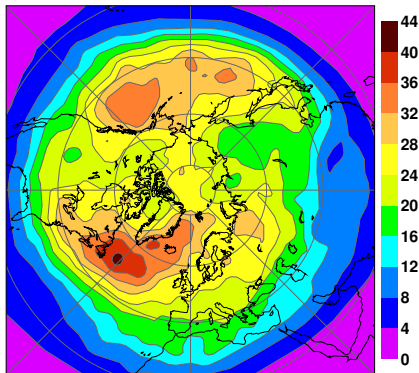
# Scale-dependent spread-error diagnostic

spread map D+2: unfiltered fields

D+2 RMSE Z500 UF DJF 2006



D+2 Spread Z500 UF DJF 2006



region	spread : RMSE
20°-90°N	1.16
35°-65°N	1.22

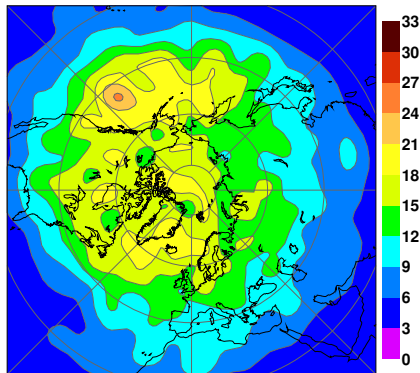
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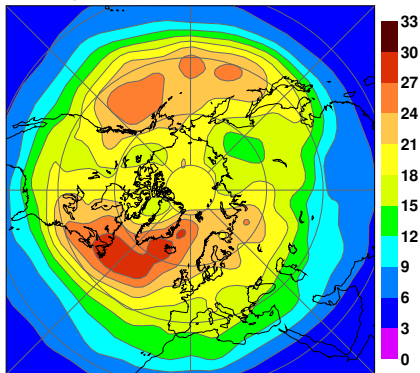
# Scale-dependent spread-error diagnostic

spread map D+2: total wavenumber 8–21

D+2 RMSE Z500 T8-21 DJF 2006



D+2 Spread Z500 T8-21 DJF 2006



region	spread : RMSE
20°–90°N	1.32
35°–65°N	1.37

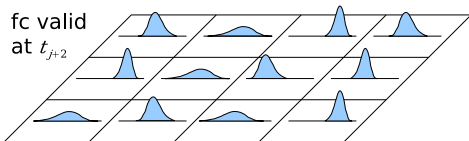
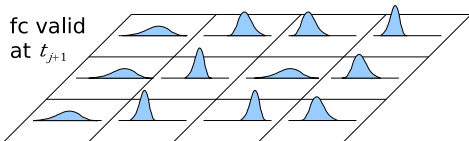
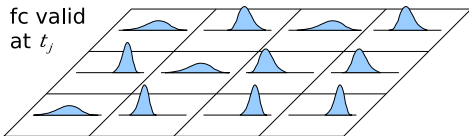
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# Assessing flow-dependent and data-dependent variations in pdf-shape

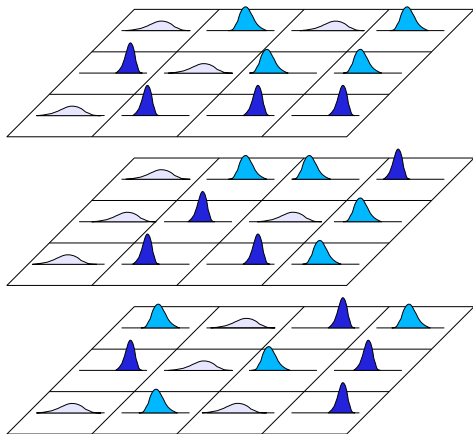
sample size limited: therefore initially focus on 2nd moment of pdf, i.e. variance

- 1 Stratification by ensemble standard deviation: spread-reliability
- 2 Modified event definition: EM error  $> \theta$ , where threshold  $\theta$  depends on a “climatological” stdev of the EM error
- 3 Gaussian centred on CF or EM as reference. Stdev of Gaussian can vary geographically and seasonally.

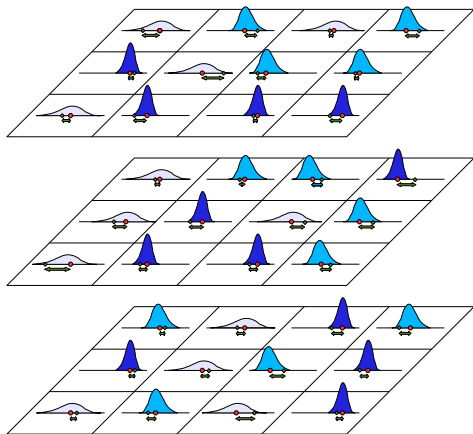
# spread-reliability: methodology



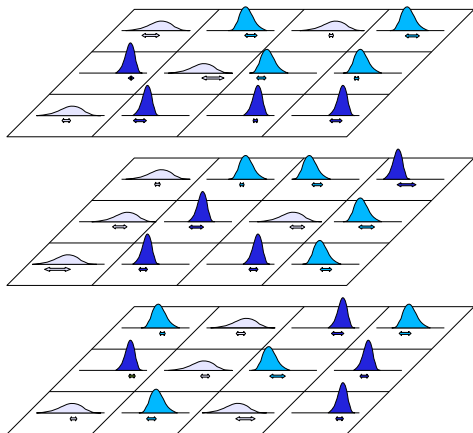
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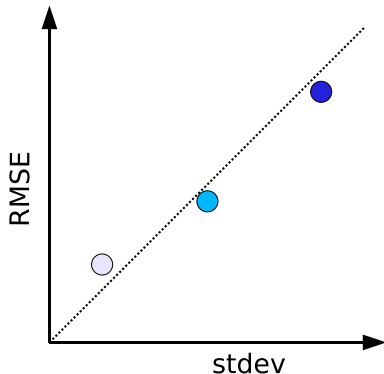
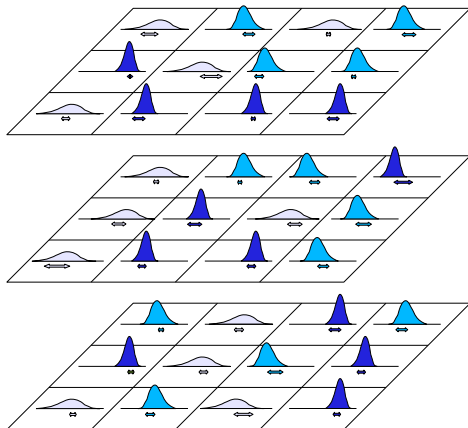
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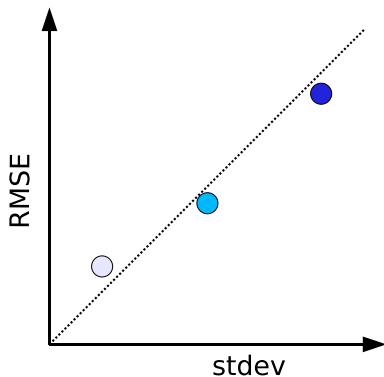
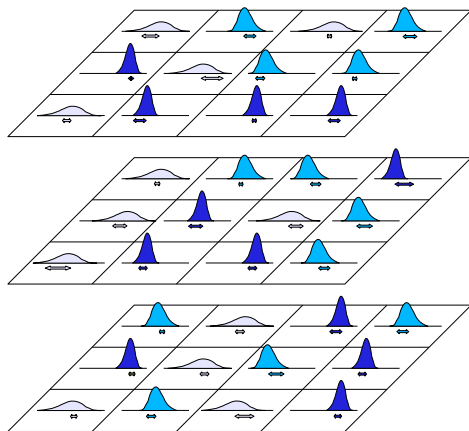
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see also Leutbecher and Palmer (2008); Leutbecher et al. (2007)



# spread-reliability: Z500 DJF06/07

Z500 Stdev and ens. mean RMSE, 35N–65N, DJF06/07

$t = 24$  h

$t = 48$  h

$t = 120$  h

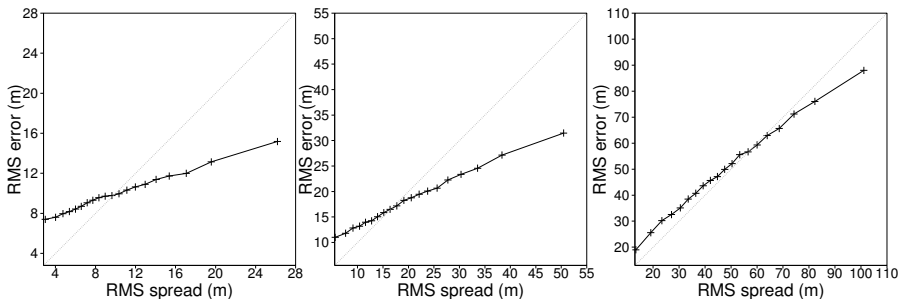


Fig. 7 from Leutbecher, Buizza & Isaksen (2007)

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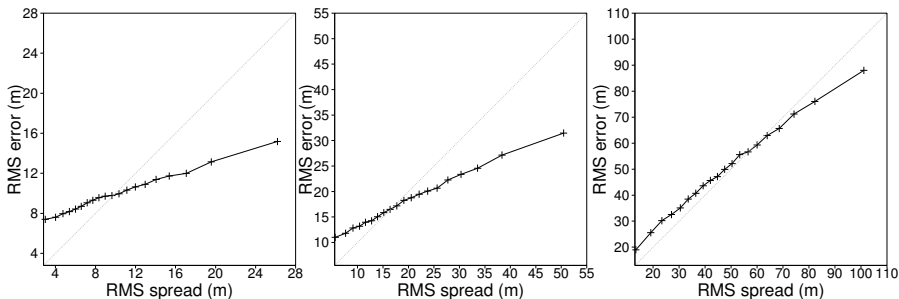


Fig. 7 from Leutbecher, Buizza & Isaksen (2007)

deficiency at the early forecast ranges; reliability improves with lead time

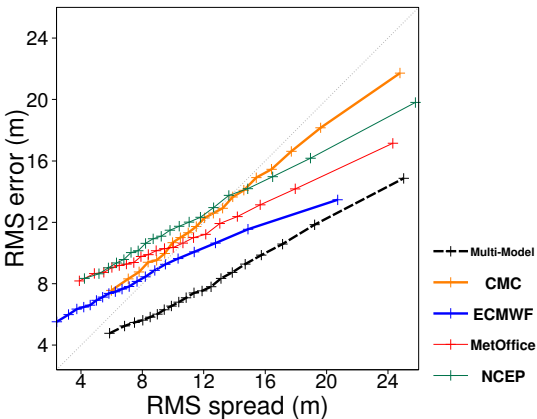
# Comparison with other ensembles in TIGGE

- data provided by Renate Hagedorn
- direct model output
- verified with quasi-independent analysis: ERA-Interim
- period: DJF2008/2009 (0 UTC, 90 start dates)
- region: N.-Hem midlatitudes ( $35^{\circ}$ – $65^{\circ}$ N)

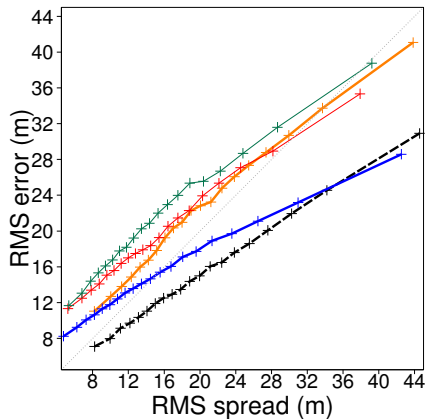
# Spread-reliability: GH 500 hPa

TIGGE comparison

24 h

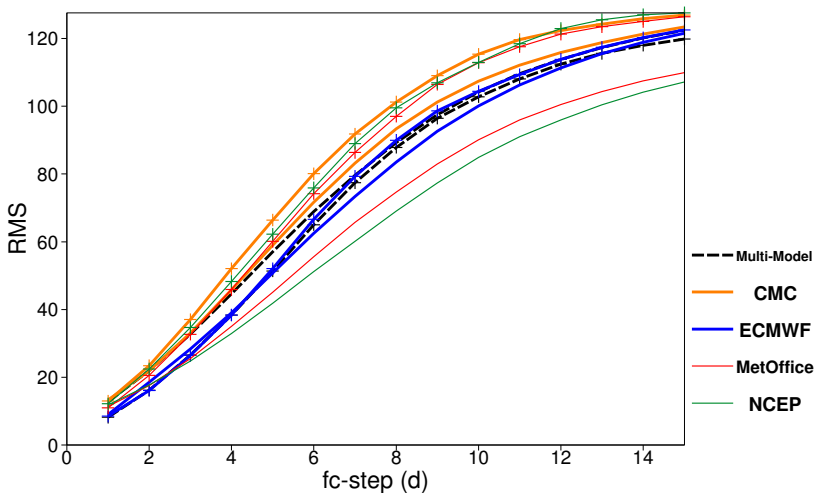


48 h



# Ens. stdev and EM RMS error: 500 hPa geopotential

TIGGE comparison

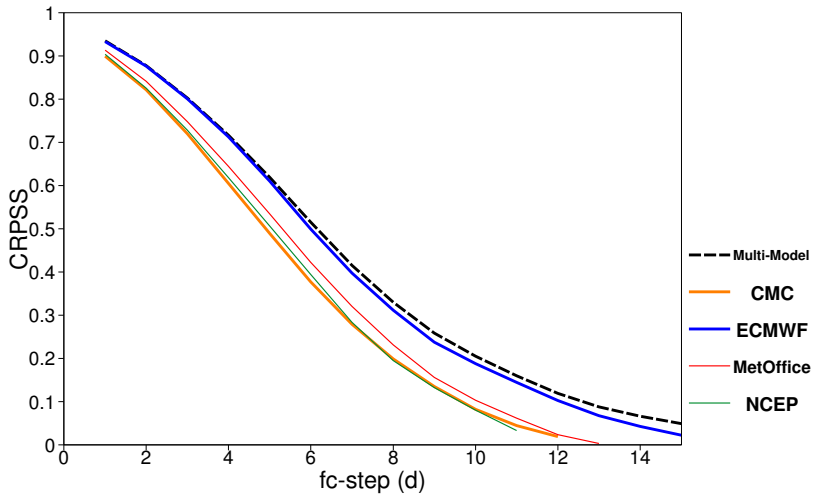


—+— Ens. Mean RMSE

— Ens. stdev

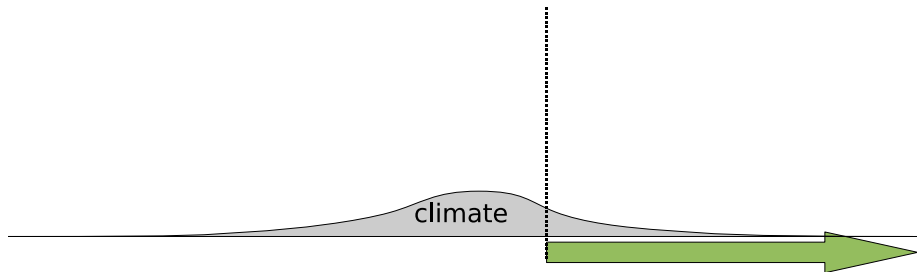
# CRPS: 500 hPa geopotential

TIGGE comparison



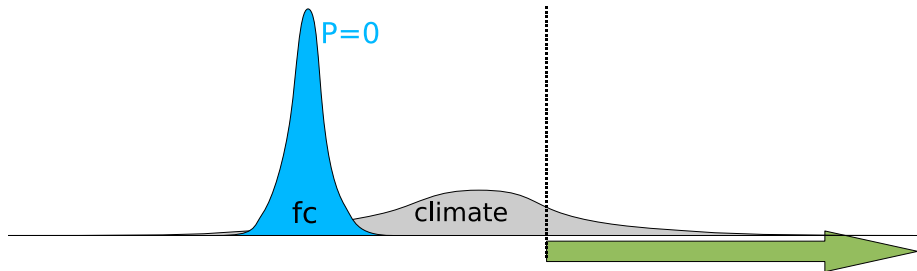
# Binary events based on the climate

consider a short lead time ...



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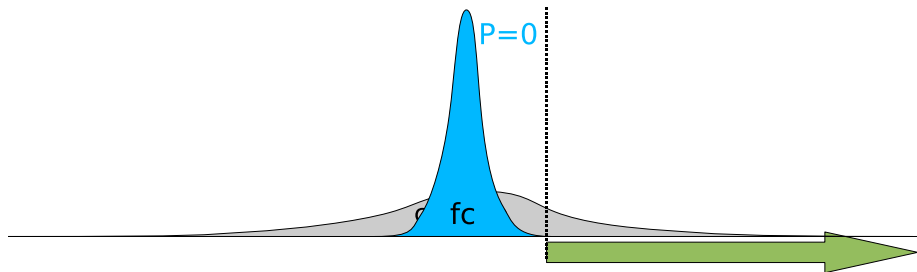
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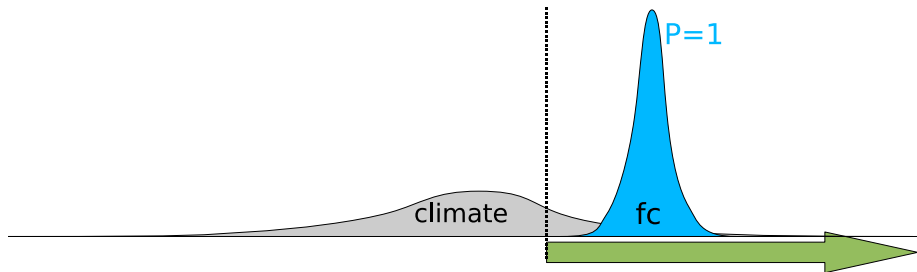
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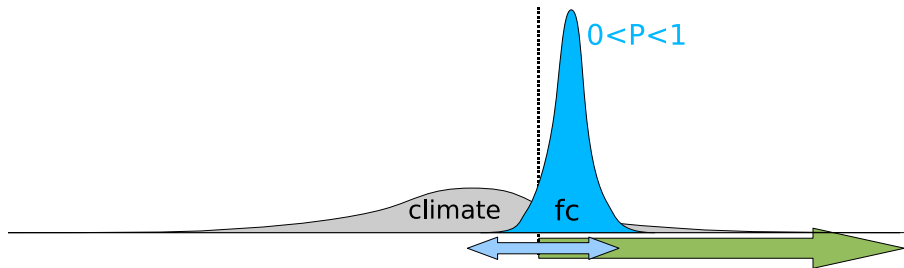
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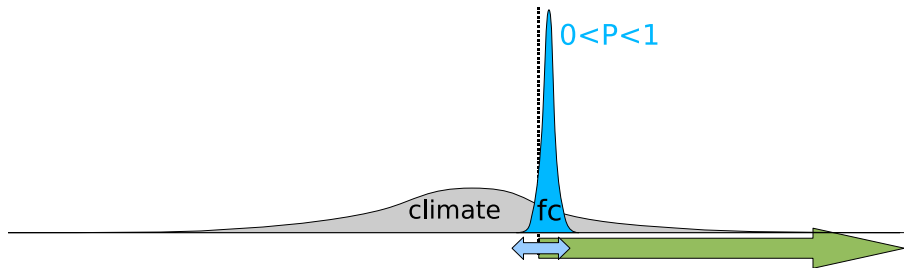
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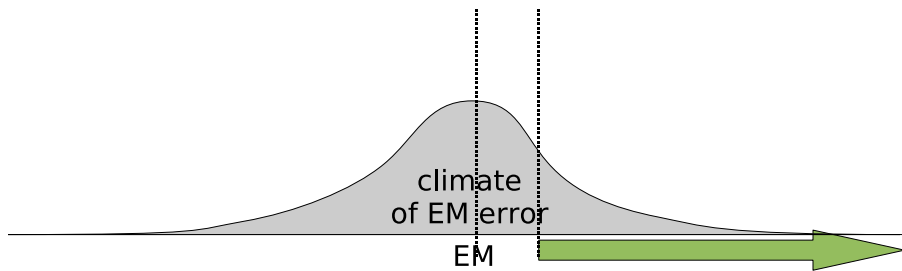


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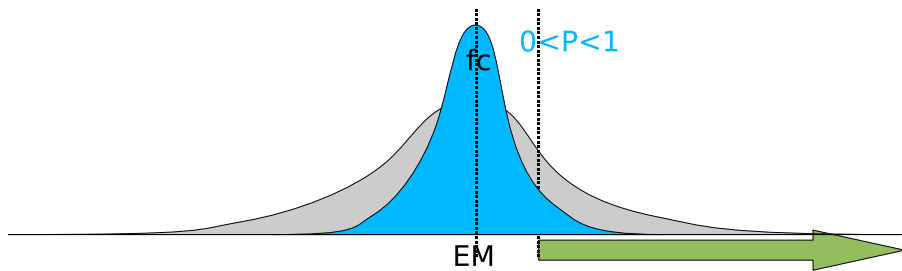
consider an *even shorter* lead time ...



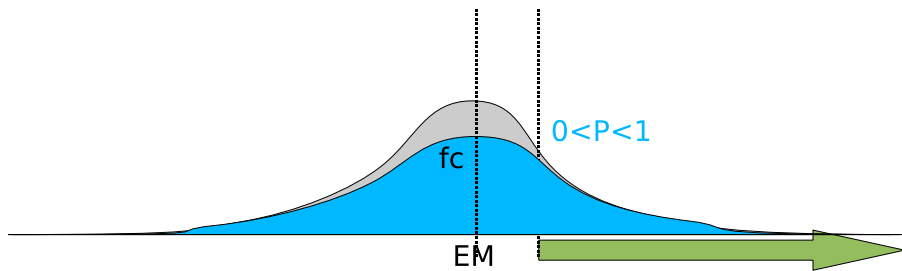
# Binary events based on Ens. mean and its error climate



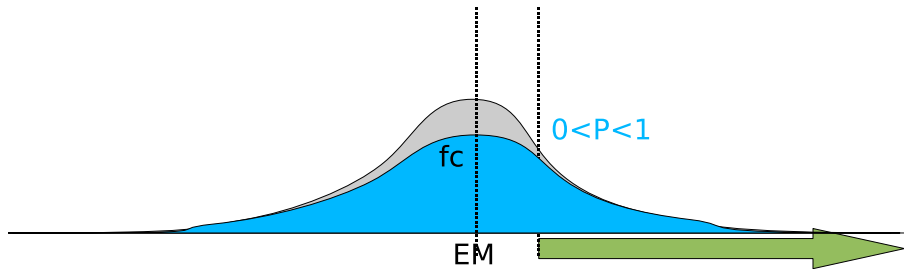
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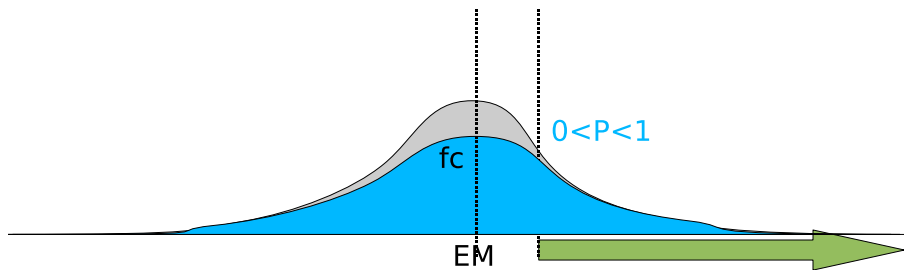
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scales naturally with lead time, expect to be better suited to diagnose skill of variations in pdf-shape



## Binary events based on Ens. mean and its error climate



scales naturally with lead time, expect to be better suited to diagnose skill of variations in pdf-shape  
can also use CF and climate of CF errors ...

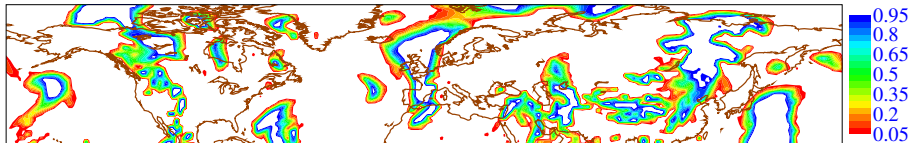
# An error climatology based on reanalyses and reforecasts

- reforecast started from ERA-40 and operational analyses (reforecasts from ERA-Interim operational since March 2009)
- 5 members (CF + 4 PF)  $\Rightarrow$  Ens. mean slightly less accurate
- 9 weeks centred on week of interest
- 18 years, once weekly  $\Rightarrow 18 \times 9 = 162$  errors
- errors for climatology computed with ERA-Interim analyses
- verification for DJF09 with operational analyses

# Probability of different kinds of events

48-hour fc of 850 hPa meridional velocity  
valid at 0 UTC on 31 January 2009

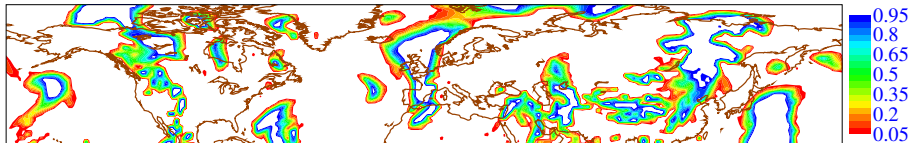
$$P(x > \mu_{\text{clim}} + \sigma_{\text{clim}})$$



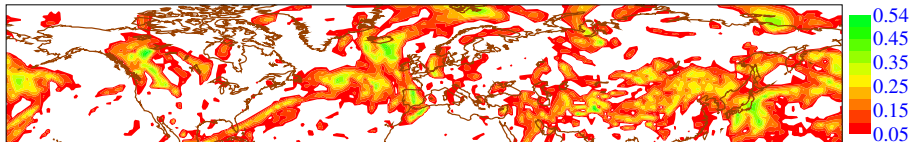
# Probability of different kinds of events

48-hour fc of 850 hPa meridional velocity  
valid at 0 UTC on 31 January 2009

$$P(x > \mu_{\text{clim}} + \sigma_{\text{clim}})$$



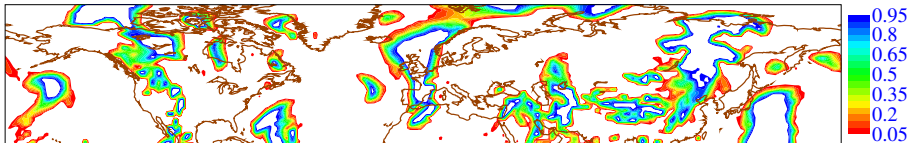
$P(x > EM + \sigma_{\text{err}})$  expect on average 0.16



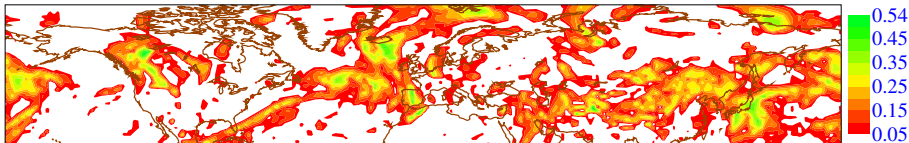
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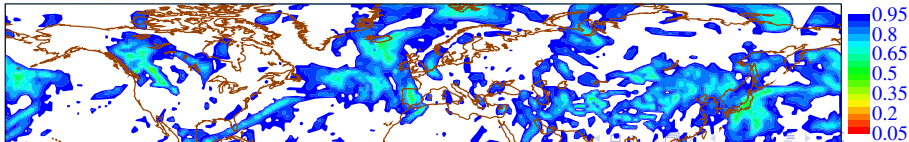
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$P(x > EM + \sigma_{\text{err}})$  expect on average 0.16



$P(x > EM - \sigma_{\text{err}})$  expect on average 0.84





# Probabilistic scores for new types of events

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$$p^{(W)}(n) = \frac{n + 2/3}{M + 4/3} \in \left[ \frac{2}{3M + 4}, \frac{3M + 2}{3M + 4} \right]$$

with  $n$  being the number of members predicting the event and  $M$  being the ensemble size. The  $p^{(W)}(n)$  are known as Tukey plotting position; cf. also Cromwell's rule and Wilks (2006).



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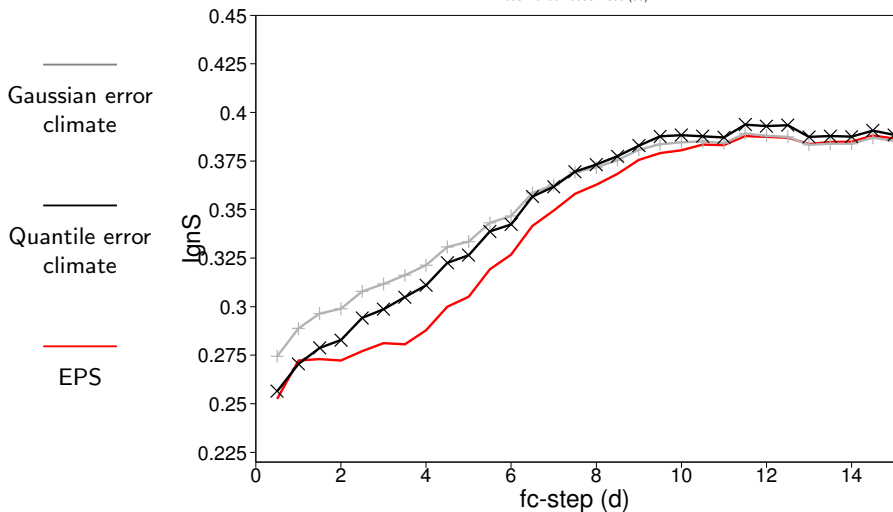
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- **ROC-area:**  $\int_0^1 H dF \in [0.5, 1]$ , where  $H$  and  $F$  denote Hit Rate and False Alarm Rate, respectively.

# Logarithmic Score for $x > EM + \sigma_{err}$

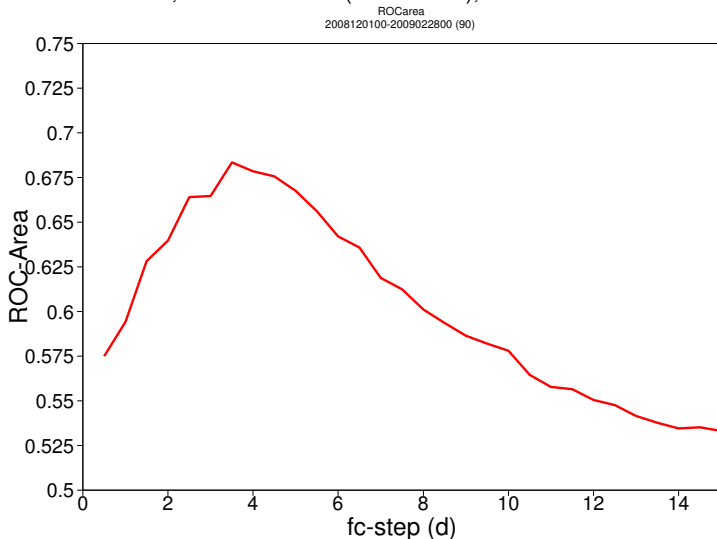
v850hPa, anom.>1 stdev (em-err-clim), Northern Mid-latitudes

IgnoranceScore, IgnoranceScoreGaussianClimate, IgnoranceScoreQuantileClimate  
2008120100-2009022800 (90)



# Area under the ROC for $x > \text{EM} + \sigma_{\text{err}}$

v850hPa, anom.>1 stdev (em-err-clim), Northern Mid-latitudes



## Intermediate Summary: Events relative to EM/CF

- Overall, scores (BS, IgnS, ROC-area) indicate that EPS has more skill in predicting variations in pdf shape than climatological error pdf
  - ▶ However, additional EPS skill tends to be relatively small initially.
  - ▶ It increases to max typically around  $t \approx 4 \pm 2$  d.
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- Initial skill increase consistent with fact that spread-error reliability improves with lead time
- Work in progress . . .
  - ▶ What should be expected from a good EPS system?
  - ▶ Can we get additional insight by using this technique to compare different ensemble configurations?
  - ▶ What can we learn from this for ensemble calibration techniques?

# Evaluation of the pdf $p(x)$ of a continuous variable

- Two proper scores
  - ▶ Continuous Ranked Probability Score (CRPS)
  - ▶ Continuous Ignorance Score (ClgnS)
- Two reference forecasts are considered:
  - ▶  $N(\text{CF}, \sigma_{\text{err}}^2(\text{CF}))$ :  
 $\Delta$ score between EPS and  $N(\text{CF}, \sigma_{\text{err}}^2)$  evaluates **all moments** of pdf
  - ▶  $N(\text{EM}, \sigma_{\text{err}}^2(\text{EM}))$ :  
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- What difference should be expected?
  - ▶ define two kinds of “perfect probabilistic forecast”
  - ▶ an analytical example
- Results for the operational ECMWF EPS



# The Continuous Ranked Probability Score

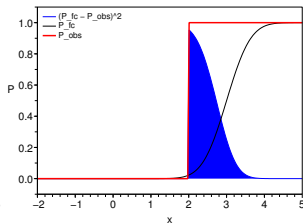
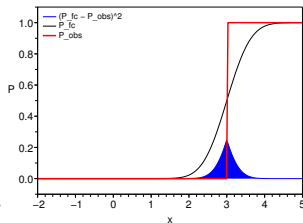
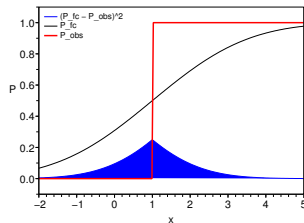
CRPS = Mean squared error of the cumulative distribution  $P_{fc}$

$$\text{cdf of truth} \quad P_y(x) = P(y \leq x) = H(x - y) \quad (1)$$

$$\text{cdf of forecast} \quad P_{fc}(x) = P(x_{fc} \leq x) \quad (2)$$

$$\text{CRPS} = \int (P_{fc}(x) - P_y(x))^2 dx \quad (3)$$

$$= \int \text{BS}_x dx \quad (4)$$



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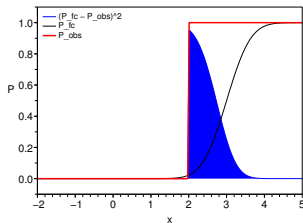
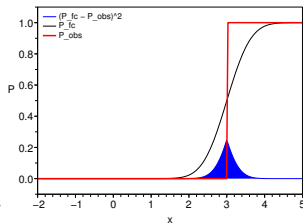
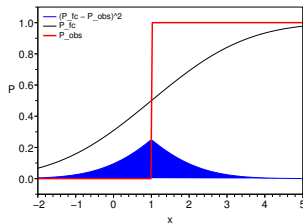
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equal to Mean Absolute Error for a deterministic forecast

# Continuous Ignorance Score

or Continuous Logarithmic Score

Let  $y$  denote truth and  $p$  the forecasted probability density

$$\text{CIgnS} = -\log p(y)$$

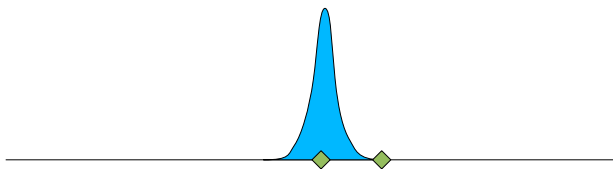


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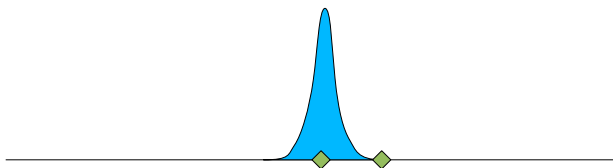


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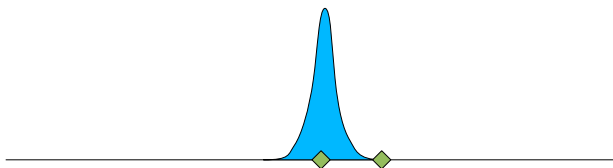
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Mean squared error of reduced centred variable plus  
logarithmic penalty term for the spread ( $\sigma$ ).

# Perfect probabilistic forecasts

- Usually: skill score = 0  $\Rightarrow$  as good as climatology  
skill score = 1  $\Rightarrow$  perfect deterministic forecast
- We may still get closer to 1 but will *never* reach it!

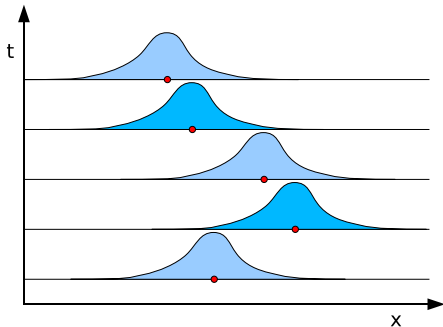
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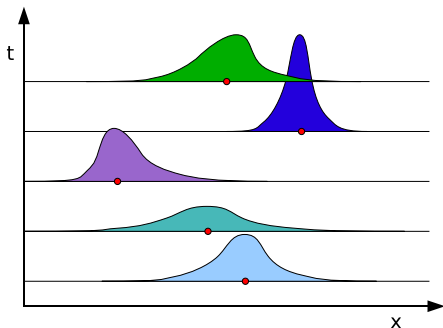
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# Levels of perfection

use label  $t$  to refer to different valid times of the forecast (lead time fixed)

**Perfect dynamic forecast:** Perfect flow- and data-dependent variations in pdf-shape

$$p_t(x) = p_d(x - \mu_t, t)$$

with  $p_d$  statistically consistent with error of the mean  $\mu_t$  for each  $t$ ,  
given average variance  $\mathbb{E}_t \int x^2 p_d(x, t) = \bar{v}$  and mean zero for each  $t$ .

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**Perfect static forecast:** Constant (or seasonally varying) flow- and data-independent pdf-shape which is perfect:

$$p_t(x) = p_s(x - \mu_t)$$

with  $p_s$  statistically consistent with the error of the mean  $\mu_t$  in the time-average sense, and  $\int x p_s dx = 0$ , and  $\int x^2 p_s dx = \bar{v}$ .

# An idealized example with Gaussian distributions

Now, focus on variance prediction.

Let Ens. Mean error be a random variable distributed according to

$$p^*(x, t) = \frac{1}{\sigma(t)\sqrt{2\pi}} \exp\left(-\frac{x^2}{2\sigma^2(t)}\right)$$

- perfect dynamic forecast: issue  $p_d = N(\mu_t, \sigma^2(t))$
- perfect static forecast: issue  $p_s = N(\mu_t, \overline{\sigma^2})$  with  $\overline{\sigma^2} = \mathbb{E}_t \sigma^2(t)$

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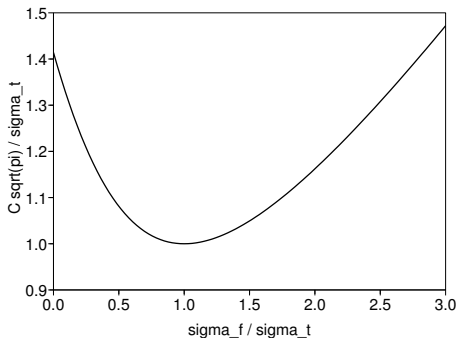
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- perfect static forecast: issue  $p_s = N(\mu_t, \overline{\sigma^2})$  with  $\overline{\sigma^2} = \mathbb{E}_t \sigma^2(t)$

What is the difference in probabilistic scores (CRPS, CIgnS) between the perfect dynamic forecast and the perfect static forecast?

## Expected value of CRPS

Let  $y$  denote the true value of the EM error. We see that the expected value of the CRPS is

$$\mathbb{E}_y \text{CRPS}(N(0, \sigma_f^2), y) = \frac{\sigma_t}{\sqrt{\pi}} \left[ -\frac{\sigma_f}{\sigma_t} + \sqrt{2 + 2\sigma_f^2/\sigma_t^2} \right]$$

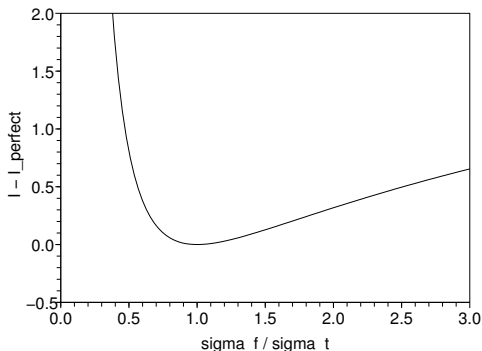


The CRPS has a minimum value at  $\sigma_f = \sigma_t$ . This is not surprising as the CRPS is a proper score.

## Expected value of ClgnS

Again let  $y$  denote the true value of the EM error. The expected value of the Continuous Ignorance Score is

$$\mathbb{E}_y \text{ClgnS}(N(0, \sigma_f^2), y) = \frac{1}{2} [\ln(2\pi\sigma_f^2) + (\sigma_t/\sigma_f)^2]$$



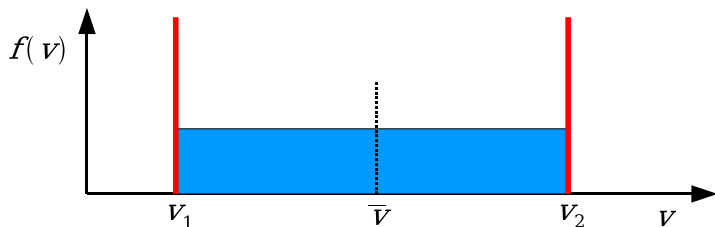
The minimum is again at  $\sigma_f = \sigma_t$ ; ClgnS is proper!



## Two particular distributions of variance

Let  $v = \sigma^2$  denote the variance

- continuous uniform distribution:  $v \sim U(v_1, v_2)$
- discrete uniform distribution:  $v \sim \frac{1}{2}\delta(v - v_1) + \frac{1}{2}\delta(v - v_2)$

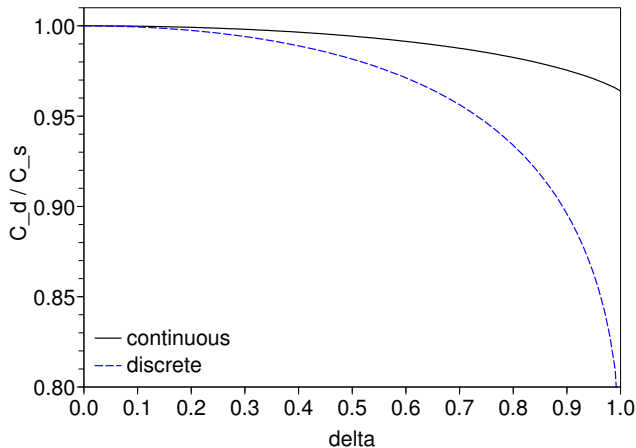


Introduce dimensionless parameter

$$\delta = \frac{v_2 - v_1}{2\bar{v}} \in [0, 1]$$

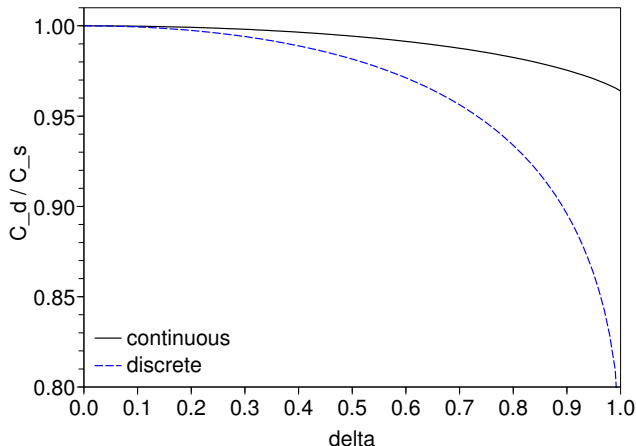
# Expected CRPS for uniform variance distributions

CRPS ratio: dynamic forecast / static forecast



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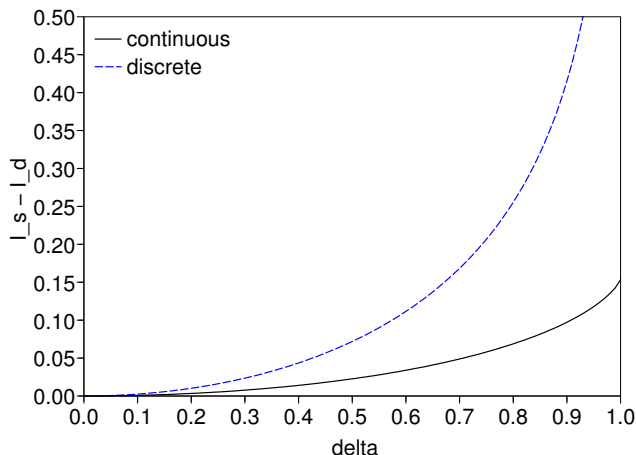


for 48-hour error doubling:

3% (20%) reduction in CRPS  $\implies$  2-hour (13-hour) gain in lead time

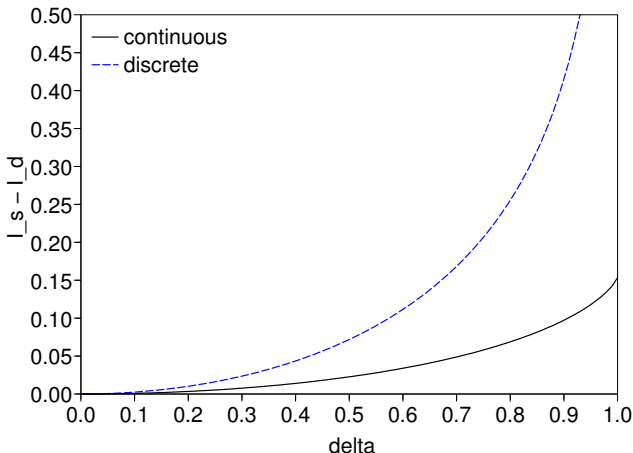
# Expected ClgnS for uniform variance distributions

$\text{ClgnS}(\text{ static forecast } ) - \text{ClgnS}(\text{ dynamic forecast } )$



# Expected ClgnS for uniform variance distributions

$\text{ClgnS}(\text{ static forecast } ) - \text{ClgnS}(\text{ dynamic forecast } )$



for 48-hour error doubling:

reduction of ClgnS by 0.15 (0.7)  $\implies$  10-hour (48-hour) gain in lead time

# Dressed ens. mean forecast: v 850 hPa, 35°–65°N, DJF09

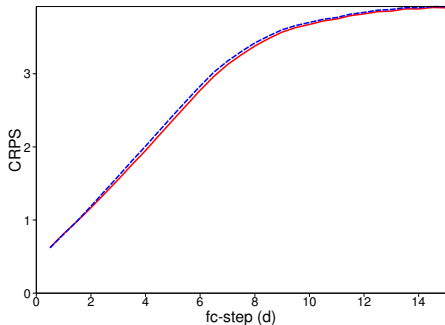
— EPS

- - -  $N(\text{EM}, \sigma_{\text{err}}^2(\text{EM}))$

raw prob. for CRPS; Gaussian for ClgnS

$\sigma_{\text{err}}$  estimated from reforecasts

## CRPS



# Dressed ens. mean forecast: v 850 hPa, 35°–65°N, DJF09

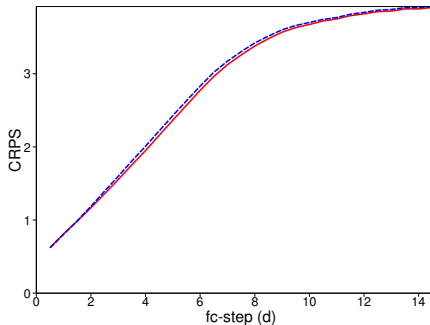
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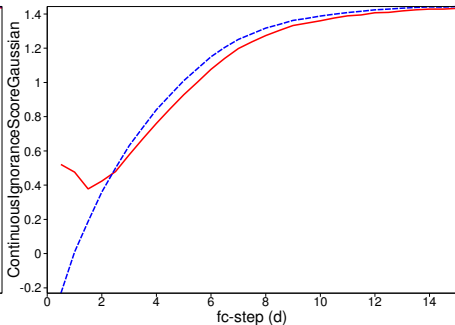
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## CRPS



## ClgnS



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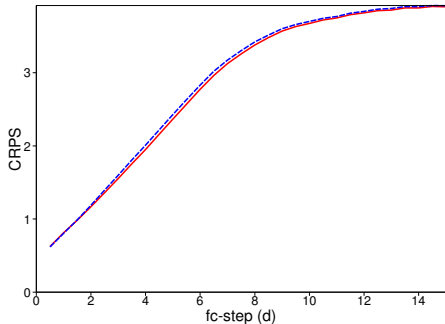
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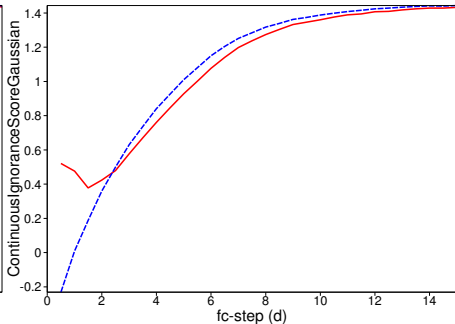
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$\sigma_{\text{err}}$  estimated from reforecasts

CRPS



ClgnS



- T850, Z500, qualitatively similar but ...
- Deficiencies in the short-range can be addressed via calibration (to a certain extent).



# Dressed control forecast: v 850 hPa, 35°–65°N, DJF09

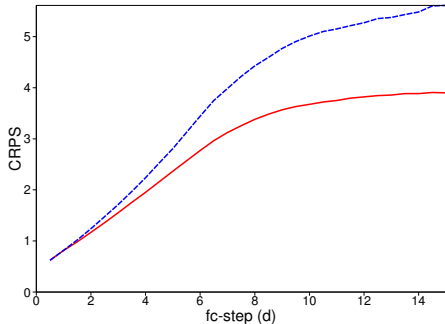
— EPS

- - -  $N(\text{CF}, \sigma_{\text{err}}^2(\text{CF}))$

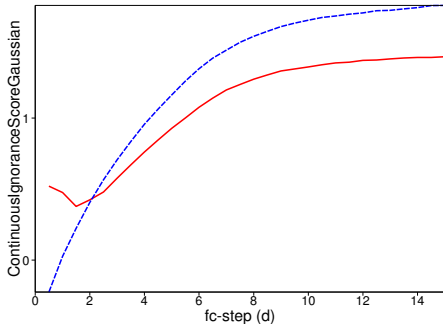
raw prob. for CRPS; Gaussian for ClgnS

$\sigma_{\text{err}}$  estimated from reforecasts

CRPS



ClgnS



# Diagnosis & Numerical Experimentation

## Initial uncertainty

- Deeper understanding from applying diagnostic techniques to clean numerical experimentation designed to answer specific questions

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- In the early ranges, say up to day 2, EM dressed with a climatological error distribution as good as or better than EPS.
- If CF/EM + past errors provide skilful probabilistic forecasts, then one may ask whether past errors might be a successful EPS perturbation strategy

# Flow-independent perturbations

- Mureau, Molteni & Palmer (1993)
  - ▶ initial perturbations based on 6-hour forecast errors from past 30 days & Gram-Schmidt-orthonormalisation
  - ▶ assimilation OI
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- Here: use random sample from past 24-hour forecast errors as initial perturbations (advantage: characteristics of short-range fc errors are closer to those of analysis errors than scaled differences of full fields)

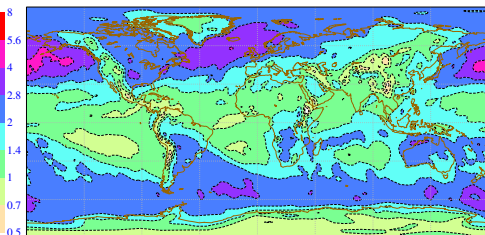
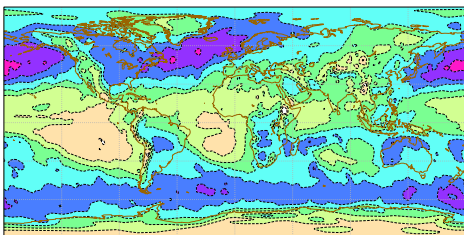


# Time mean spread vs. RMSE of Ens. mean

Meridional wind component ( $\text{m s}^{-1}$ ) at 850 hPa,  $t=48 \text{ h}$

singular vector init. perts.

24-hour fc. error init. perts.

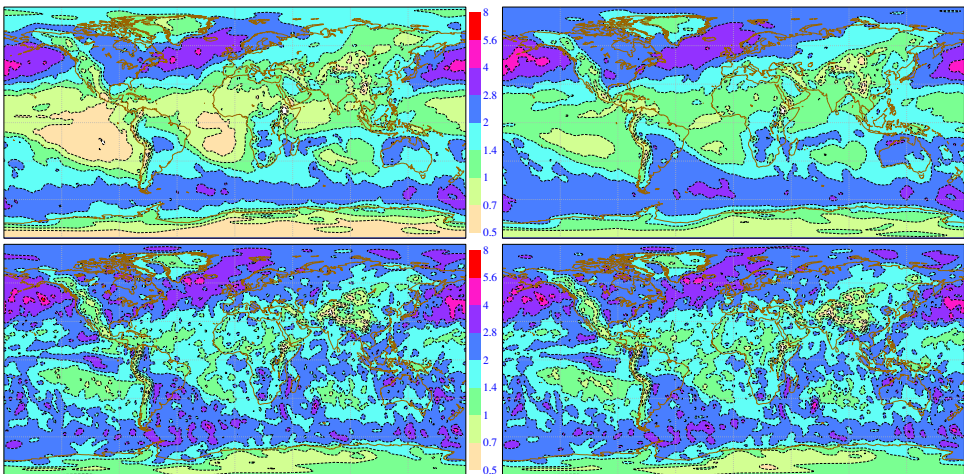


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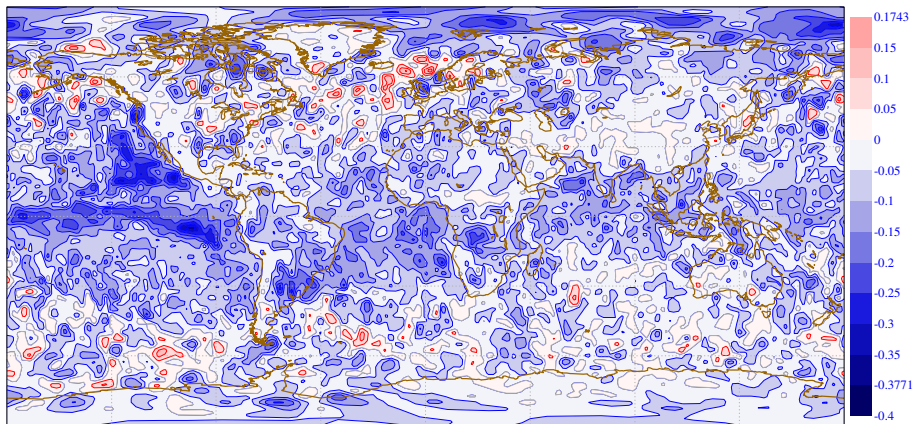
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**top:** ens. stdev.; **bottom:** ens. mean RMS error; 50 cases: 23 Nov '07–29 Feb '08  
 $T_{L255}$ , 32r3, unscaled 24-hour FCEs

# CRPS difference: FCE – SV

Meridional wind component ( $\text{m s}^{-1}$ ) at 850 hPa,  $t=48$  h

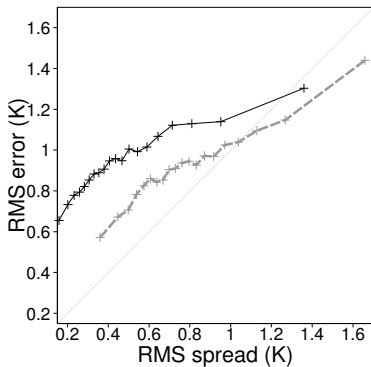


- CRPS (Continuous Ranked Probability Score  $\equiv$  mean squared error of the cumulative distribution)
- Blue means EPS based on short-range forecast errors is more skilful.
- 50 cases: 23 Nov 2007 – 29 Feb 2008

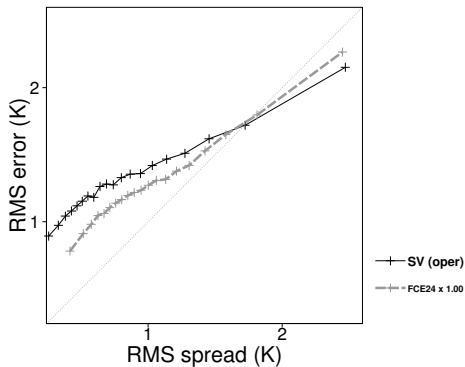
# Spread-reliability

850 hPa temperature, 35°–65°N

t850hPa, t=+24h, N.hem.mid  
N50/2007112300TO2008022900



t850hPa, t=+48h, N.hem.mid  
N50/2007112300TO2008022900



# Conclusions

- Comparison of Spread and EM-error continues to be an essential tool
  - ▶ Have not achieved a well tuned system for all variables and regions.
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  - ▶ Continuous Ignorance Score appears better suited than CRPS to evaluate flow- and data-dependent variations in spread.
- Initial perturbations based on past short-range forecast errors
  - ▶ → challenging benchmark for flow-dependent initial perturbations (e.g. singular vectors).
  - ▶ Rare locally large perturbations that are inconsistent with the flow may be an obstacle for operational implementation of the benchmark system.
  - ▶ Further diagnostic work in this area is expected to help the development of ensemble prediction system with improved flow-dependent variations of the pdf (in particular in the earlier forecast ranges).

# Issues to verify “initial uncertainty”

- truth not available
- short-range fc error correlated with analysis error → require obs (or perhaps an independent analysis)
- obs uncertainty/analysis uncertainty needs to be accounted for if ensemble spread is smaller than or of similar magnitude as the obs/an uncertainty.



# Multivariate verification

- Why? Ensemble of assimilations should provide a varying background error covariance to the assimilation system. One should attempt to verify the flow-dependent covariances.
- There are user applications that will be dependent on the joint pdf of several variables (e.g. distributed in space or several variables). Implications for suitable ensemble size can be different from univariate case.