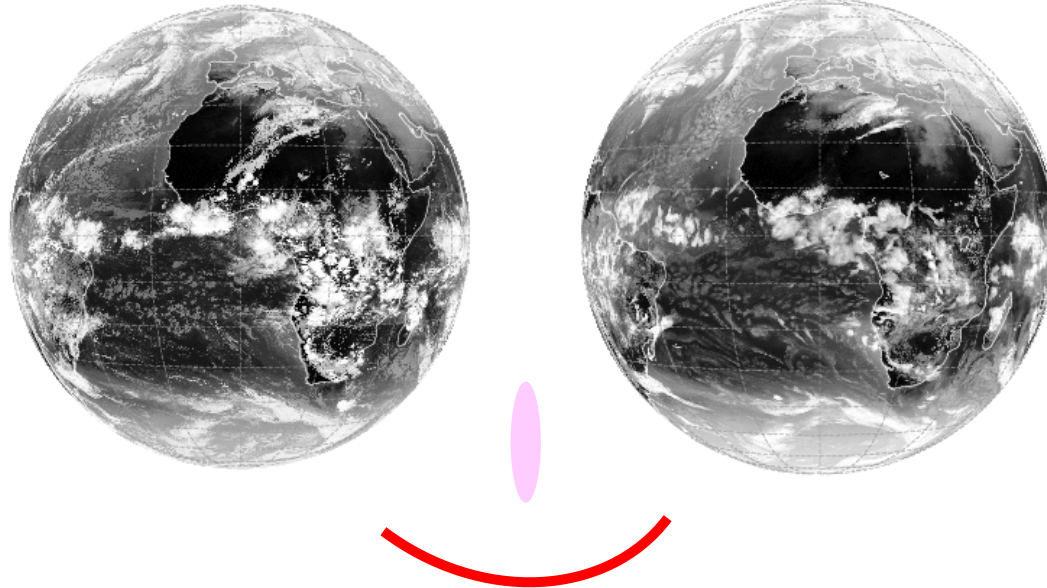


Convection and the Tropics

Meteosat 9 IR10.8 20120220 15 UTC

ECMWF oper Fc 20120220 00 UTC+15h:



Peter Bechtold

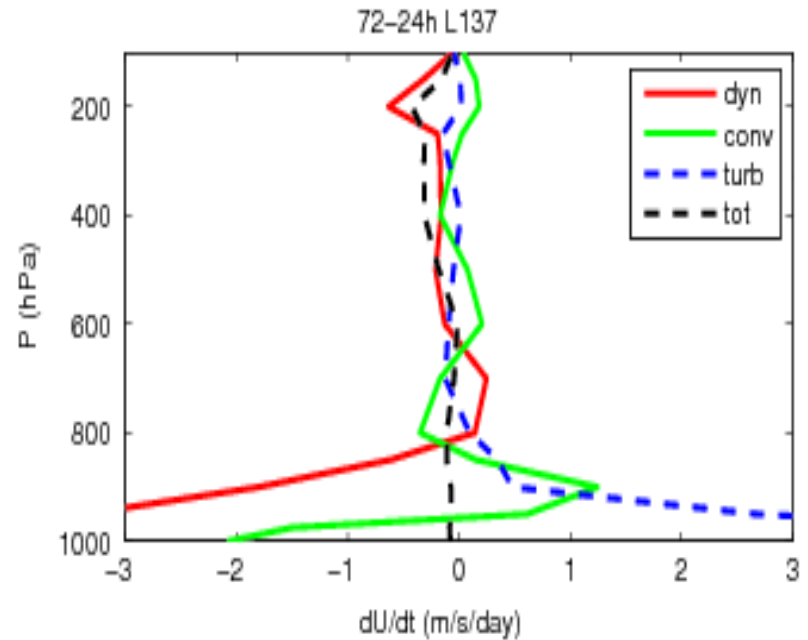
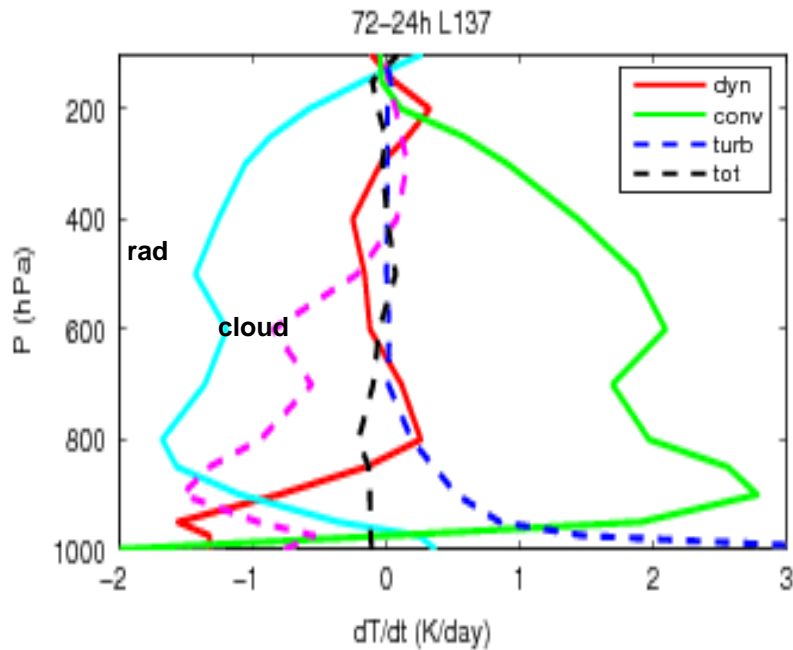
with help from:

P. Bauer, P. Berrisford, J. Bidlot, C. Cardinali., T. Haiden, L. Hirons, M. Janousek, D. Klocke, L. Magnusson, A. McNally, J-J. Morcrette, S. Malardel, F. Prates, M. Rodwell, I. Sandu., N. Semane, F. Vitart and discussions with many colleagues

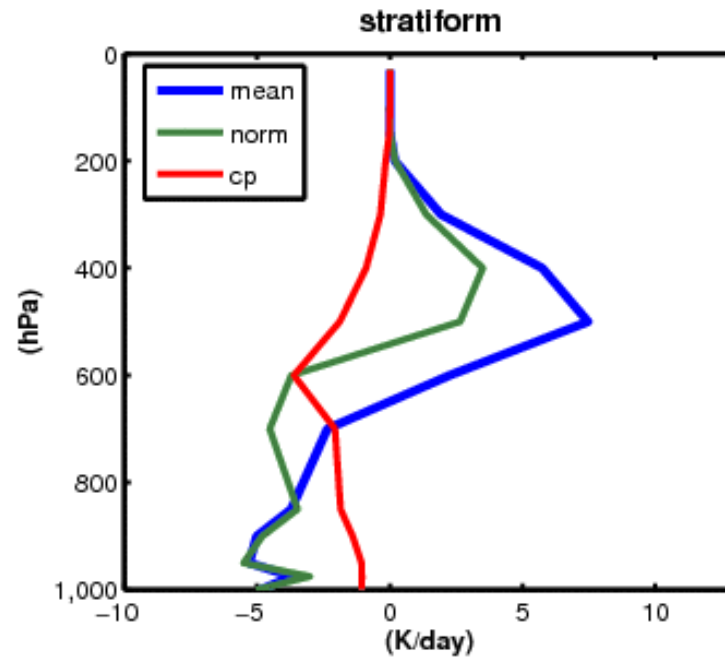
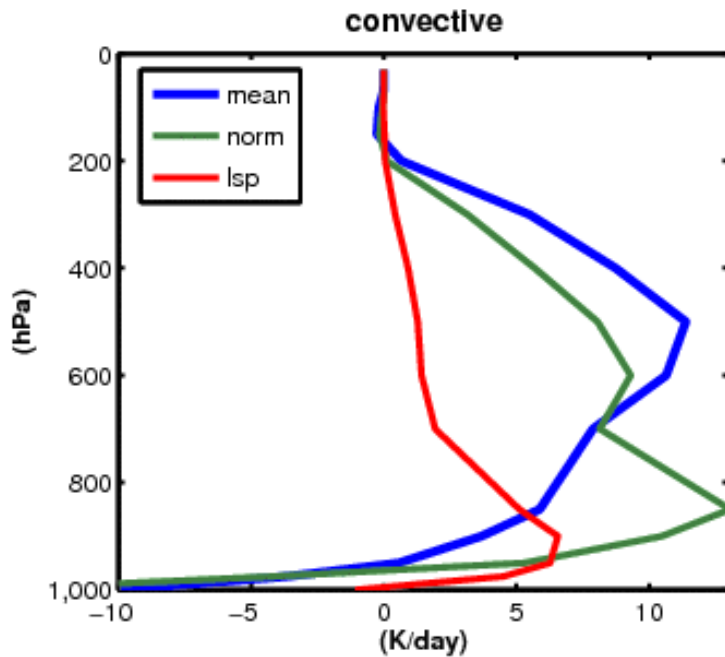
What to expect

- Convective heating – available energy - waves
- Regions with major forecast errors and uncertainty in the analysis
- Convection and seasonal forecasts: Mean state, diurnal cycle, tele-connections and the QBO
- The MJO, why better now
- The small planet, studying a future high-resolution system

Tropical T and U tendency budgets

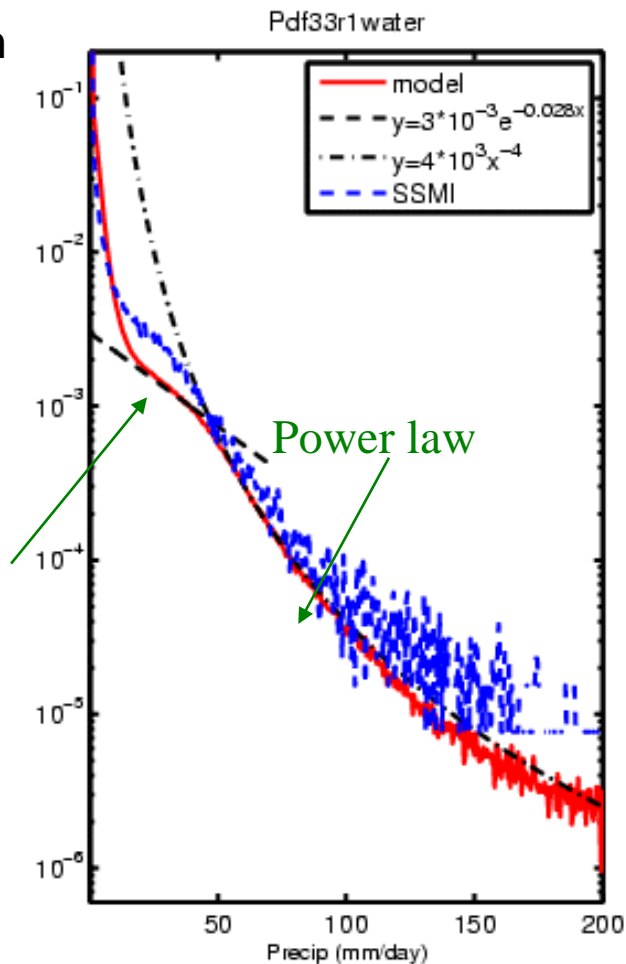


Normalized convective and stratiform heating profiles

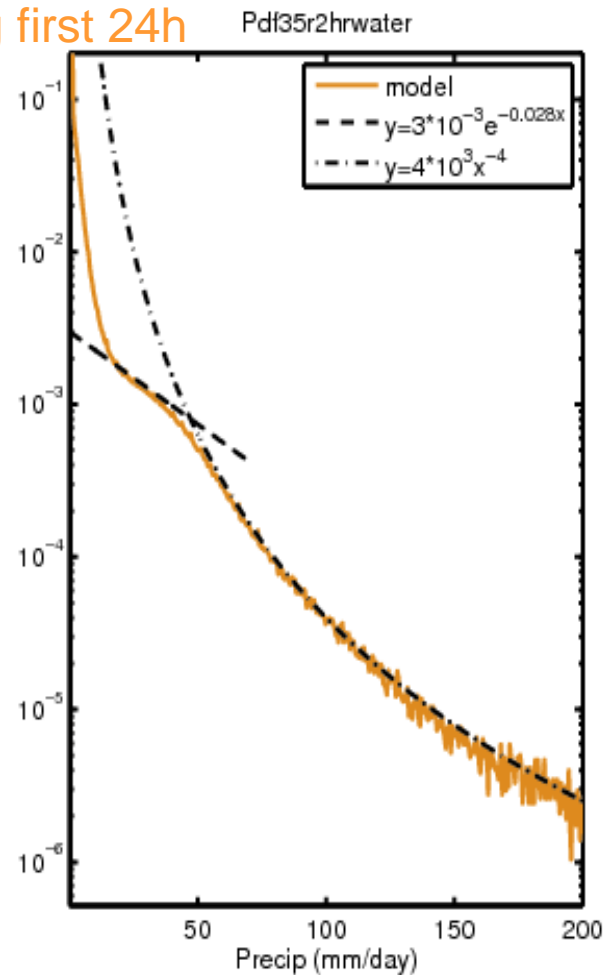


Some statistical properties of convection: Pdfs of instant. Rain rates

from T799 33r1
during first 24h

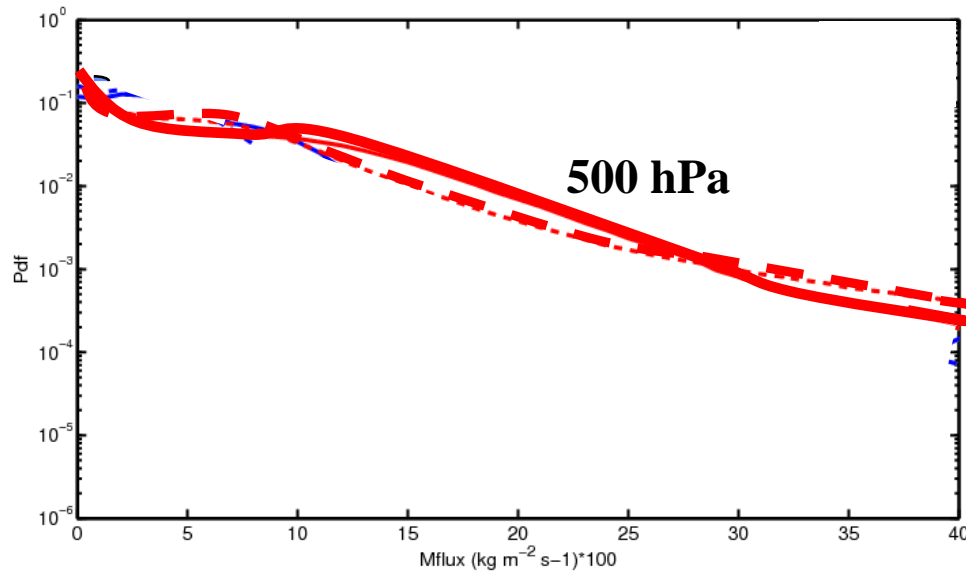


from T1279 35r2
during first 24h



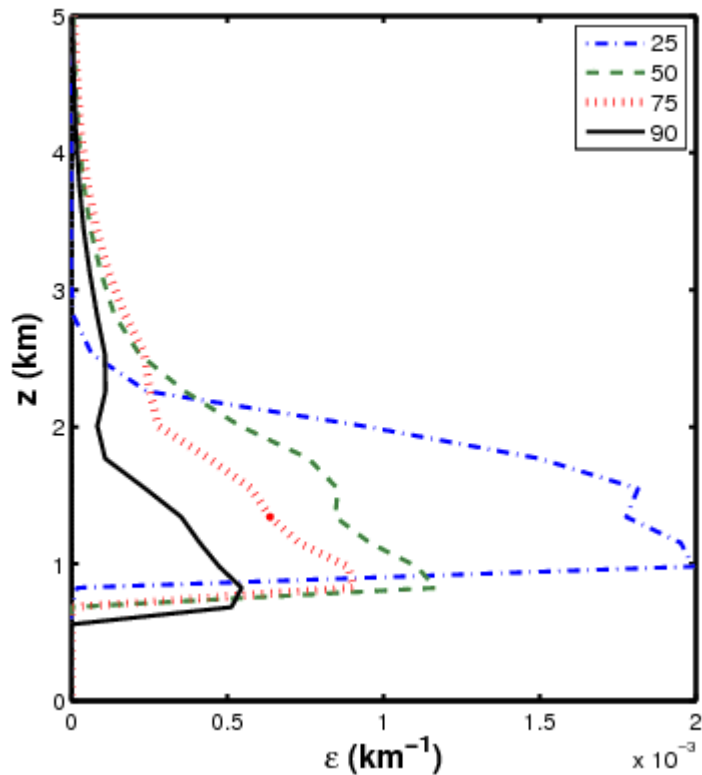
SSMI is from 1D-Var

Some statistical properties of convection: Pdfs of mass fluxes T159 & T1279

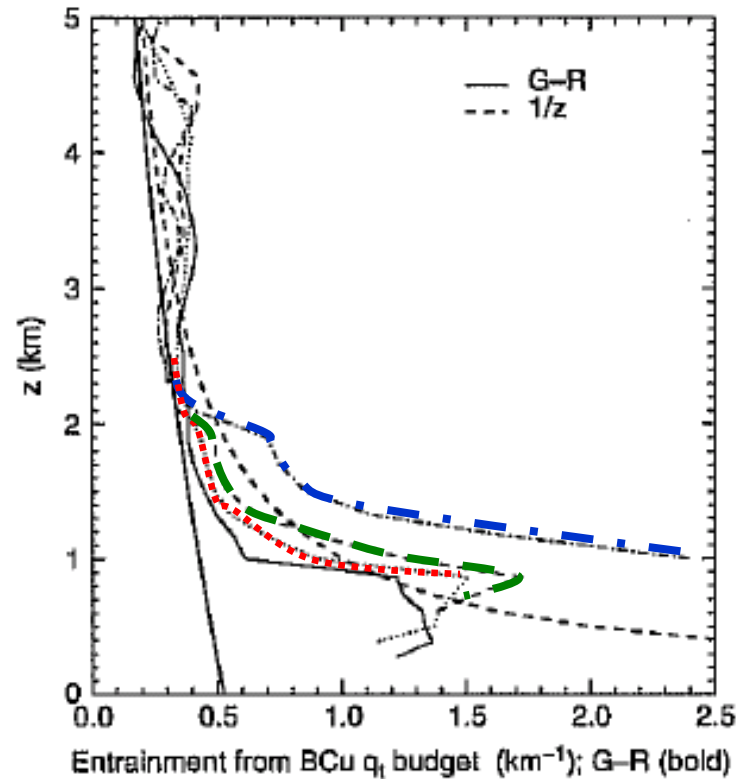


From theory expect exponential Pdf of mass flux, Cohen and Craig (2006)

Entrainment rates versus CRM (UKMO)



See also deRooy (QJ 2012)



Derbyshire et al. (QJ 2011)

The global Lorenz Energy cycle

Generation Conversion

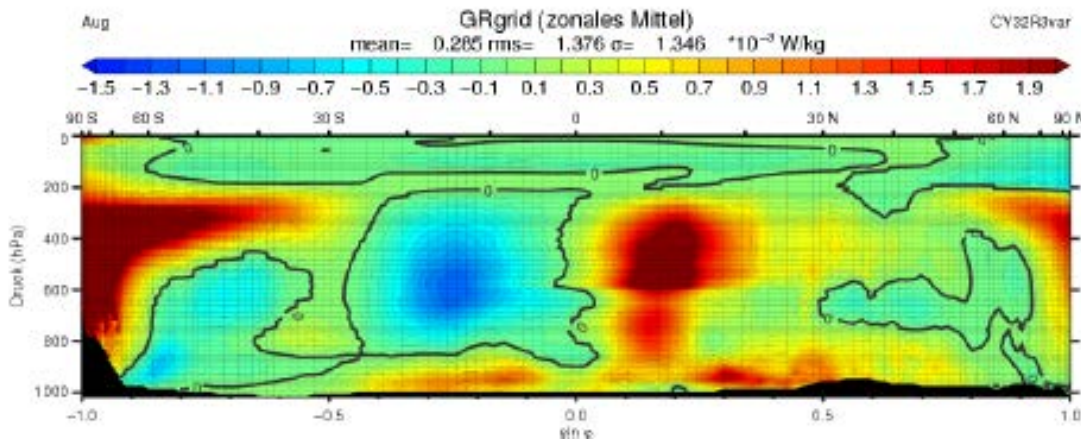
$$\frac{da}{dt} = \boxed{NQ} + \boxed{\alpha\omega} = N\bar{Q} + \bar{\alpha}\bar{\omega} + \overline{\alpha'\omega'}$$

Lorenz efficiency factor Net heating

$$\overline{\alpha'\omega'} = \frac{R}{P} [1 + (\varepsilon^{-1} - 1)] \overline{T'\omega'} + (\varepsilon^{-1} - 1) \bar{\alpha} \overline{q'\omega'}$$

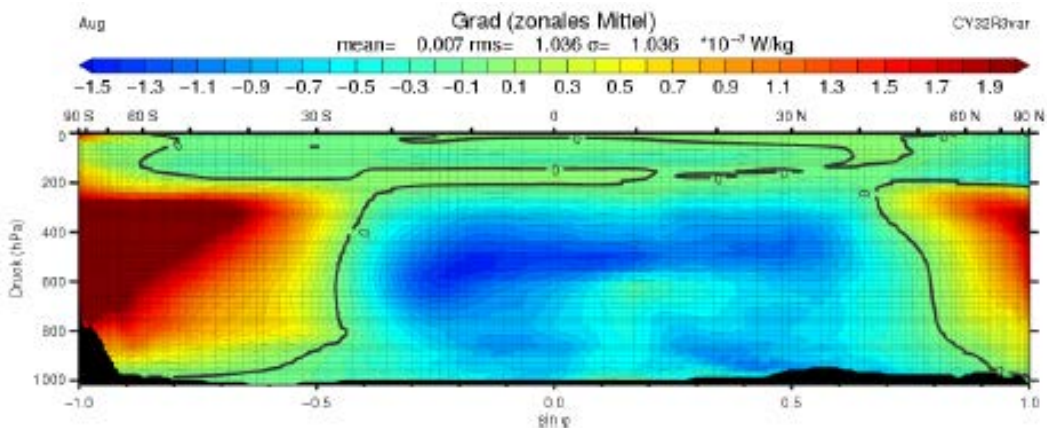
Generation rates

Total Generation rate (W/kg)



- Generation rates maximum in upper tropical troposphere

Generation rate - radiation

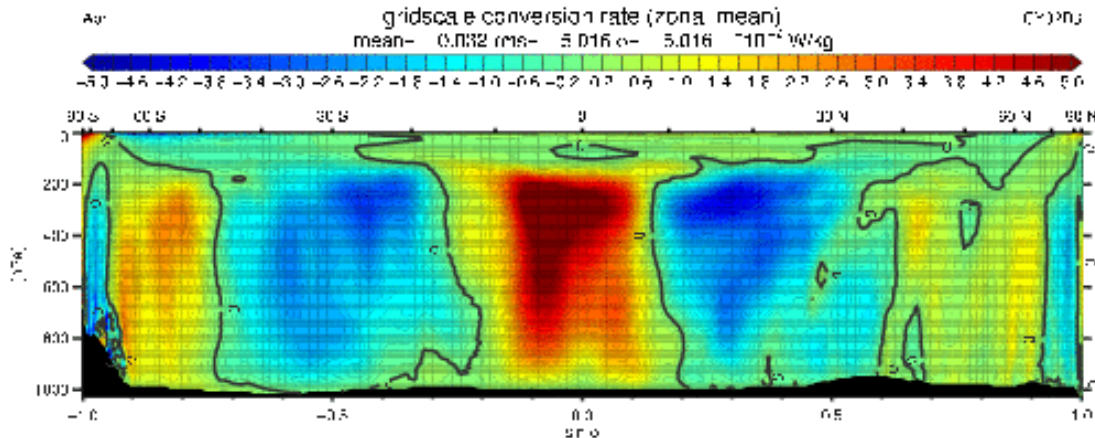


- Radiation does not contribute to the conversion rates but to the generation rate, but even there has only at poles a positive contribution (cooling at cold places) but globally a negative contribution (as in Tropics it is cooling where it is warm)

Steinheimer et al. 2008, Tellus

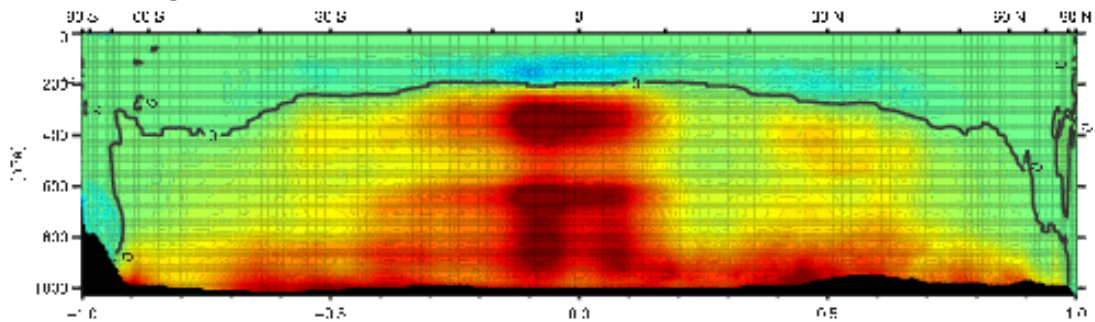
Conversion rates and convection

Grid-scale conversion rate (W/kg)

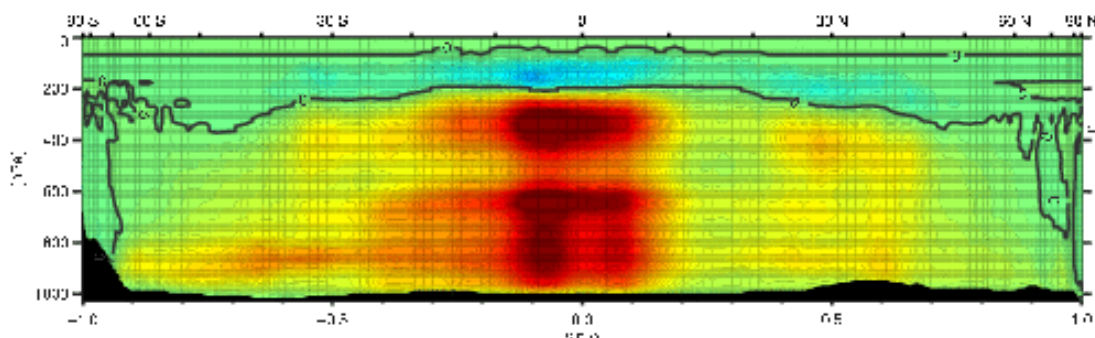


- Grid-scale has positive and negative contributions to kinetic energy conversion rate, maximum in upper-tropical troposphere

Subgrid conversion rate



Subgrid conversion rate - convection



- Convection so important because contribution always positive !

Shallow water system and linear waves

$$V = 0 \Rightarrow U = U_0 e^{-y^2/2} e^{ik(x-ct)} \quad G(z, m)$$

Kelvin wave, geostrophic

$$c = \frac{\omega}{k} = \sqrt{gh}$$

$$V \neq 0 \Rightarrow V(y) = \begin{bmatrix} 1 \\ 2y \\ 4y^2 - 1 \\ \vdots \\ H_n(y) \end{bmatrix} e^{-y^2/2} \quad G(z, m)$$

General, Hermite Polynomials
Modes alternate asymm./symmetric

$$\frac{\omega^2}{c^2} - \frac{k}{\omega} - k^2 = \frac{(2n+1)}{c}, \quad c = \sqrt{gh}; \quad n = 0, 1, 2, \dots$$

Dispersion relation

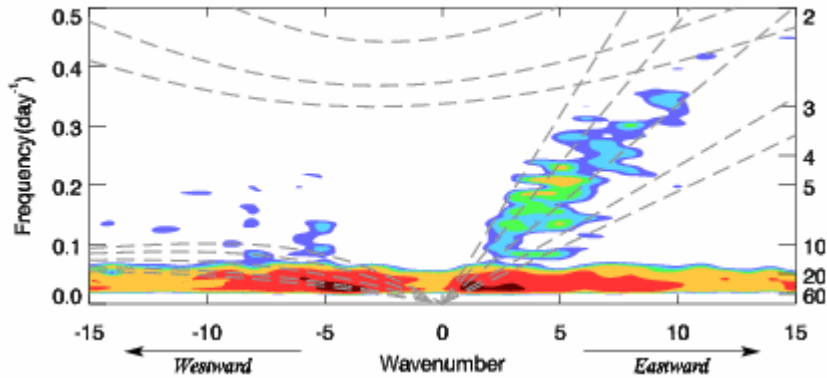
$$G(z, m) = e^{-z/(2H_s)} \operatorname{Re}(e^{-imz})$$

see T. Matsuno. Quasi-geostrophic motions in the equatorial area. Journal of the Meteorological Society of Japan, 44:25-42, 1966.

Wave number Frequency Spectra OLR

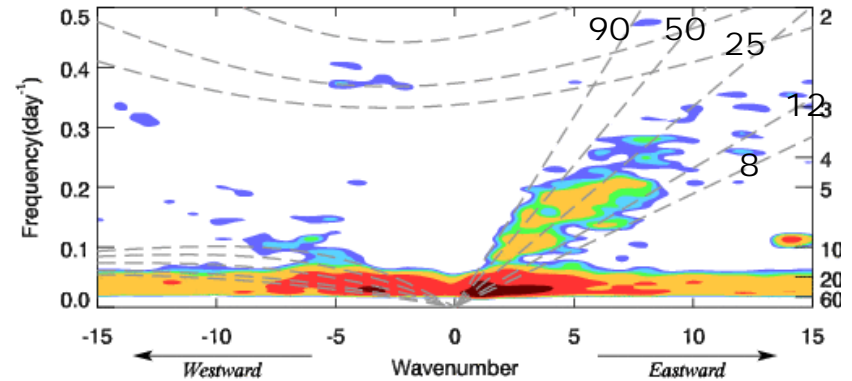
Cy38r1 (2012)

S-B



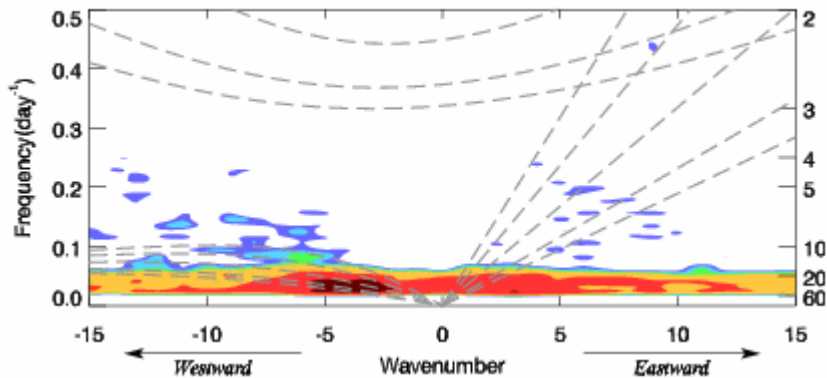
NOAA

S-B

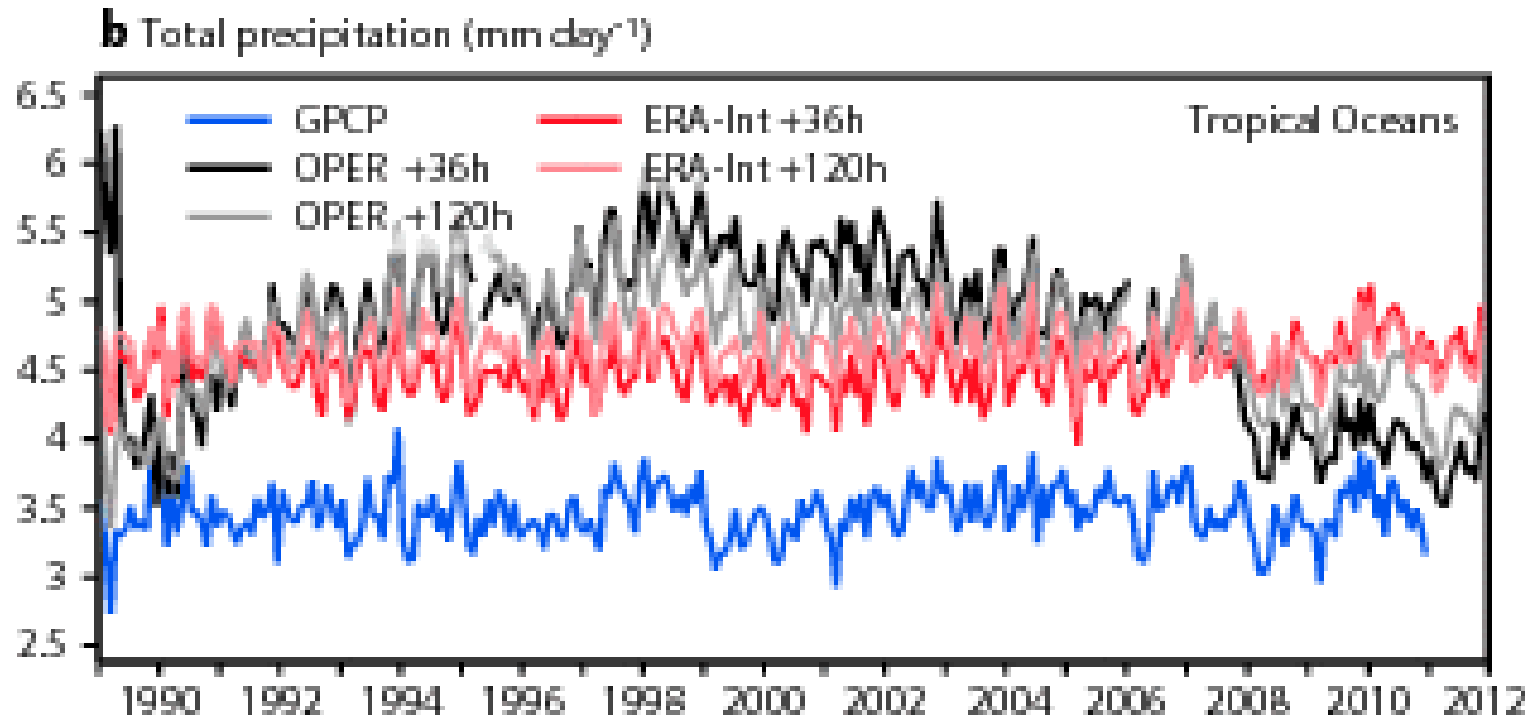


Cy31r1 (ERA-Interim)

S-B

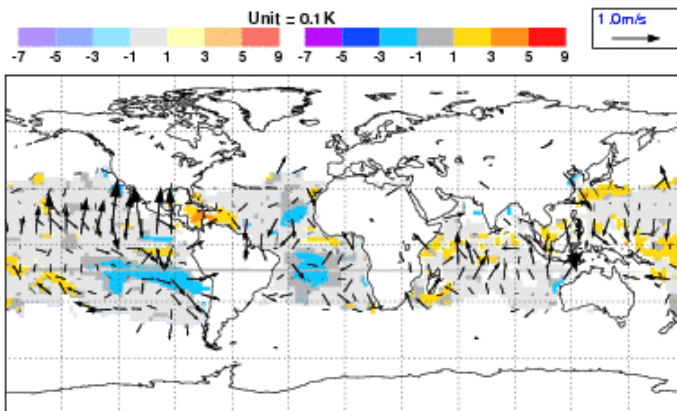


Time-series: Precipitation

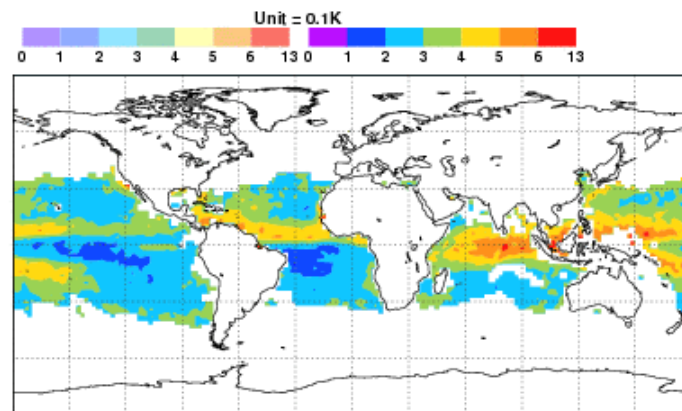


Analysis increments SON 2011

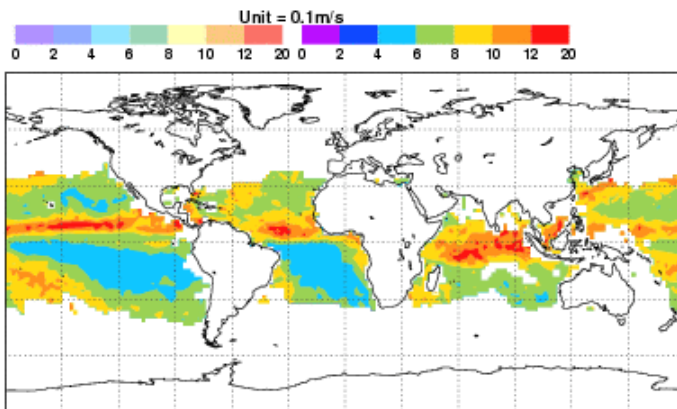
(a) 1000 hPa Mean T,v



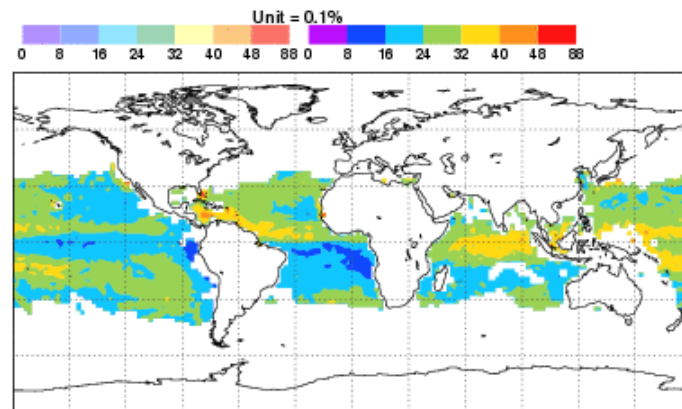
(b) 1000 hPa StDev T



(c) 1000 hPa StDev Wind.Speed

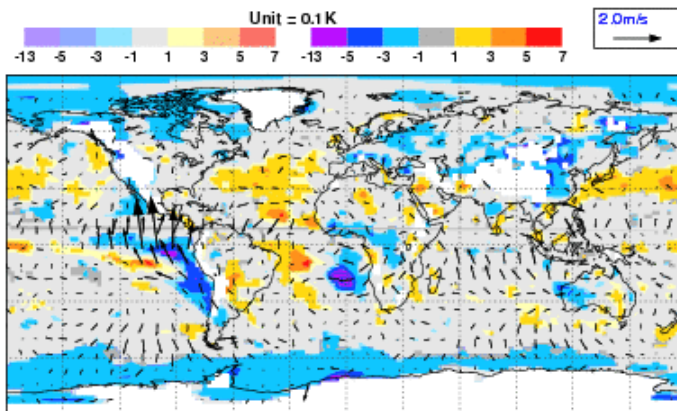


(d) 1000 hPa RMS Rel.Hum.

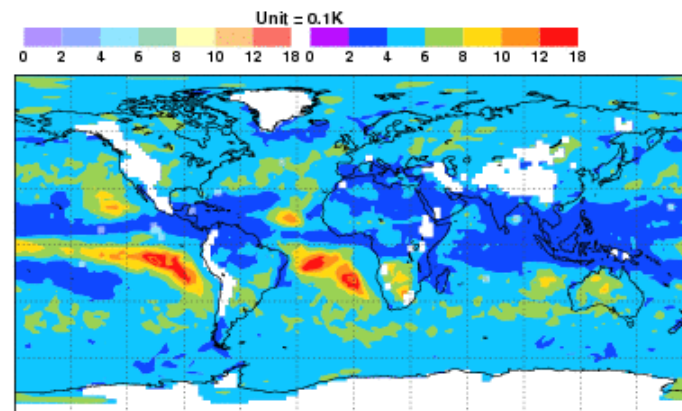


Analysis increments SON 2011

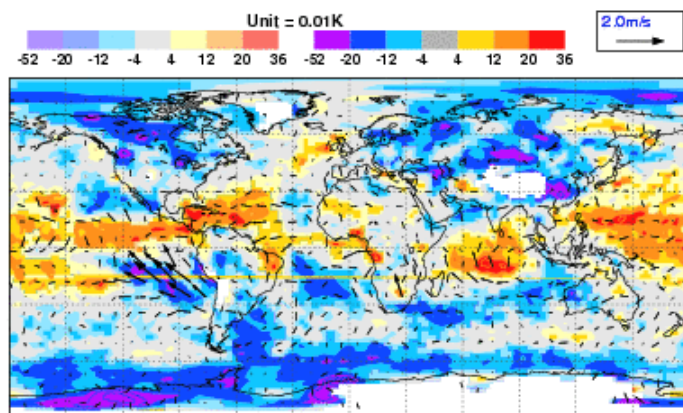
(e) 850 hPa Mean T,v



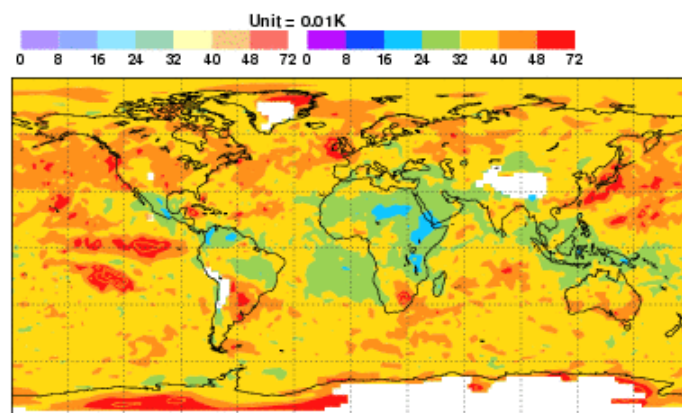
(f) 850 hPa StDev T



(a) 700 hpa Mean T,v

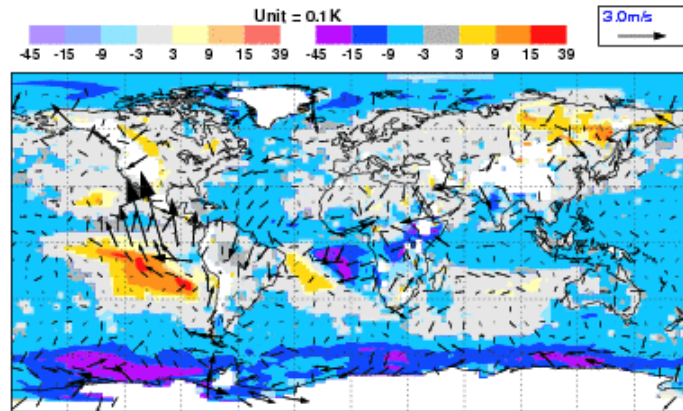


(b) 700 hPa StDev T

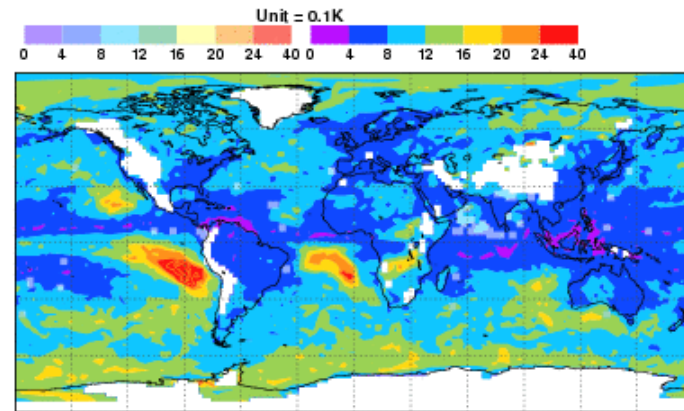


Analysis difference ECMWF-UKMO for OND 2011

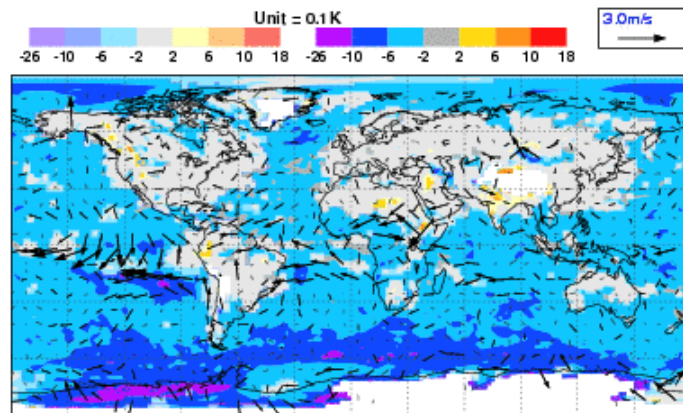
(a) 850 hPa Mean T,v



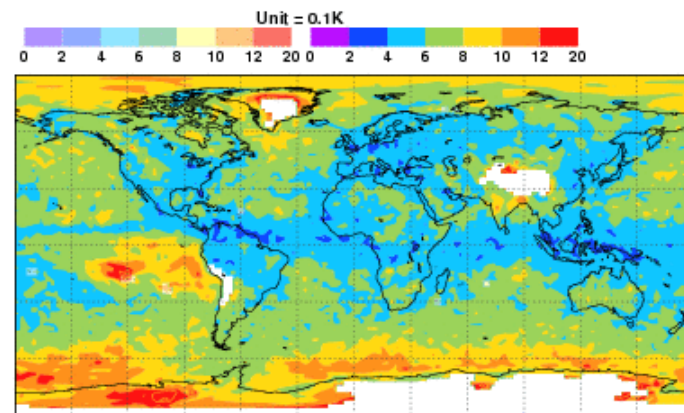
(b) 850 hPa StDev T



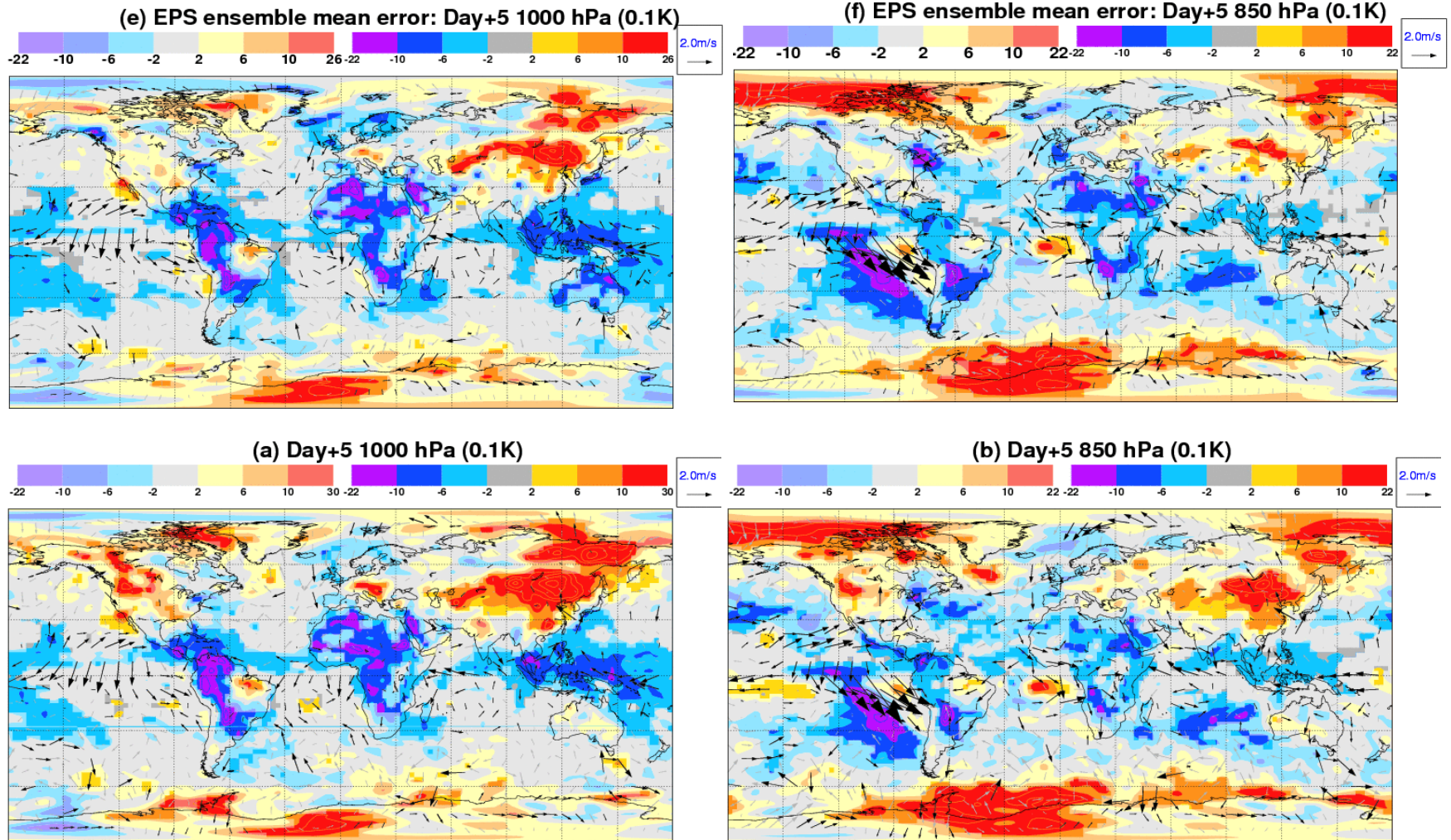
(a) 700 hPa Mean T,v



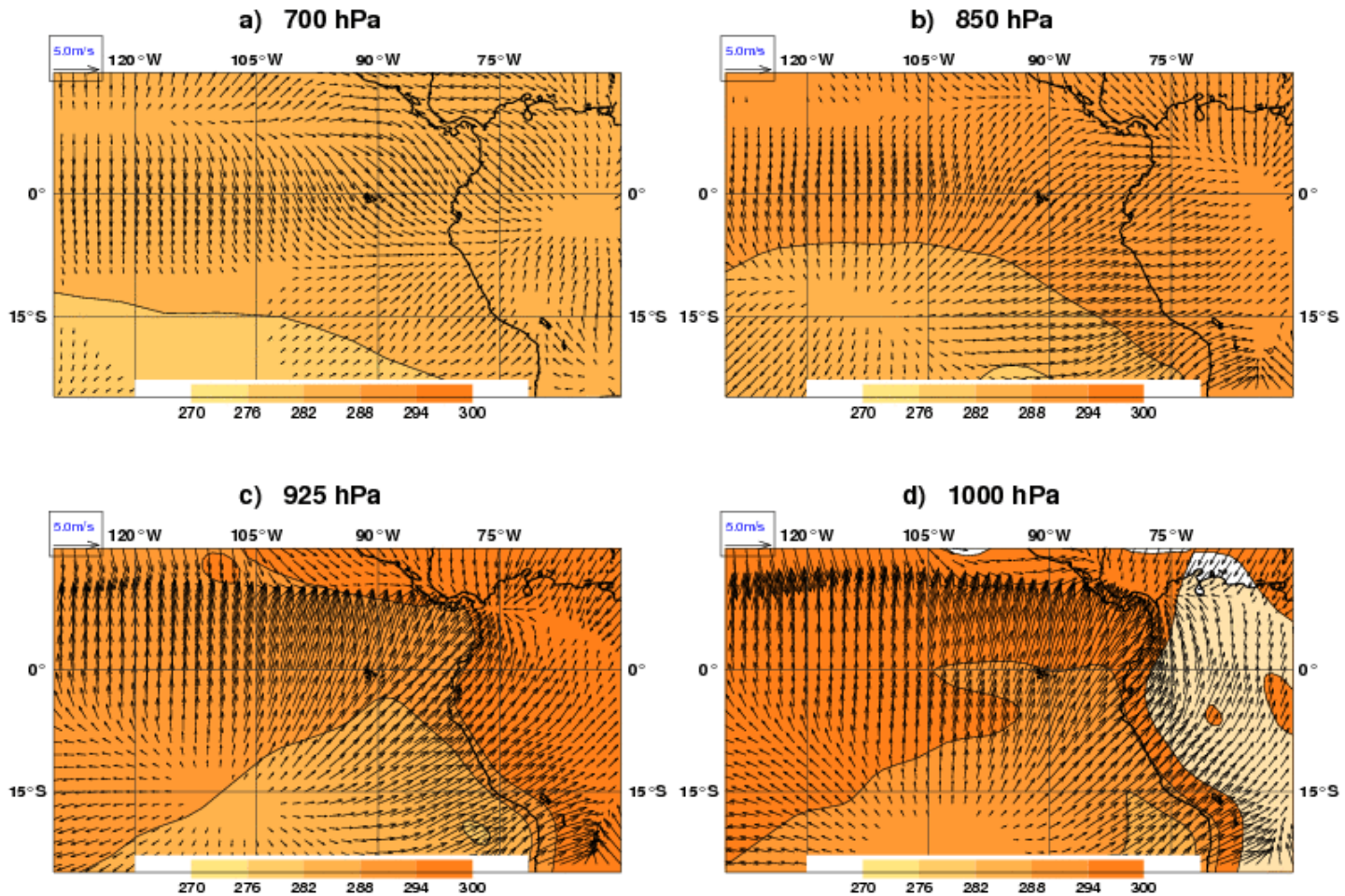
(b) 700 hPa StDev T



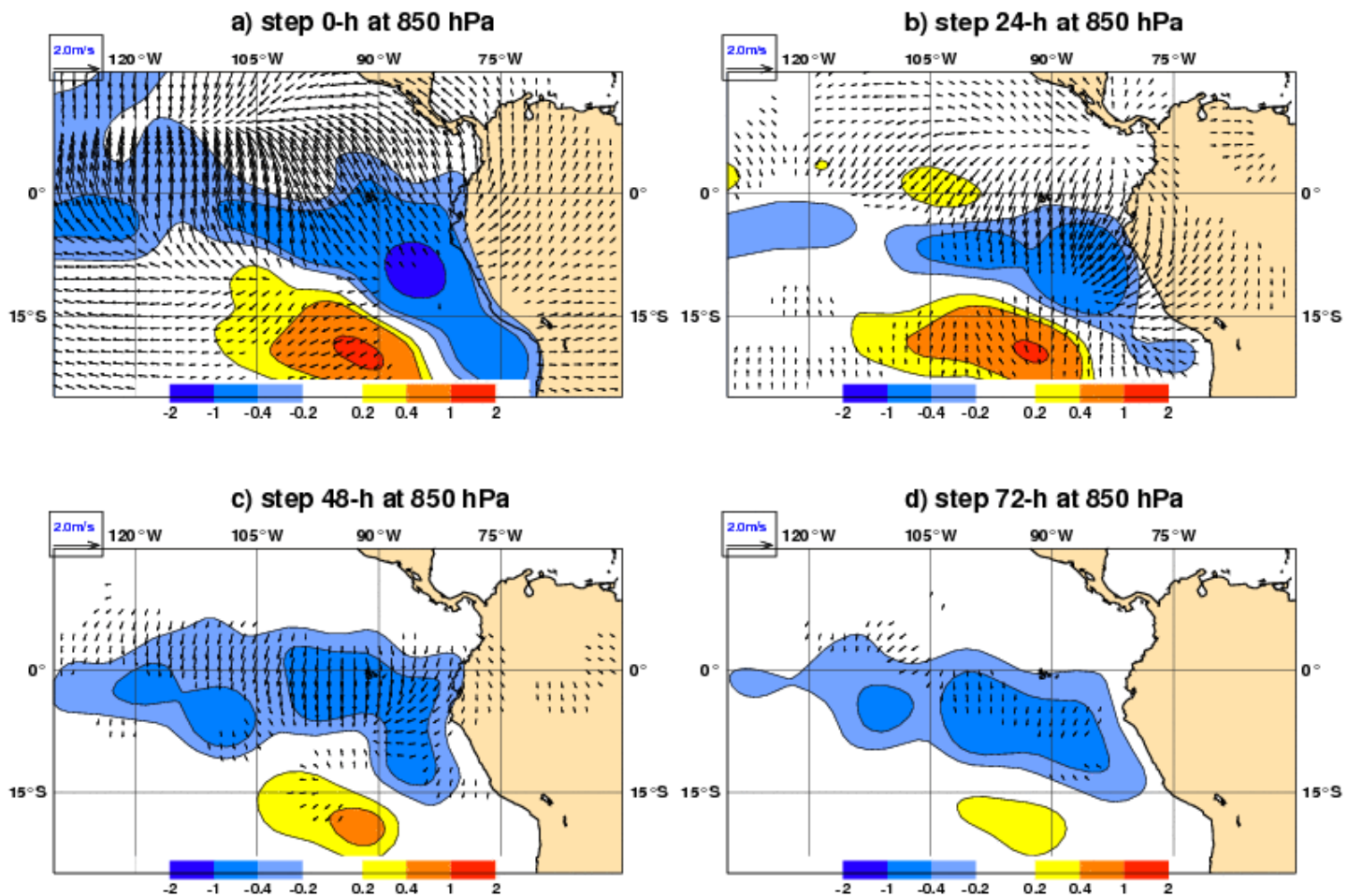
Day+5 Forecast errors: EPS ensemble mean vs. high-resolution



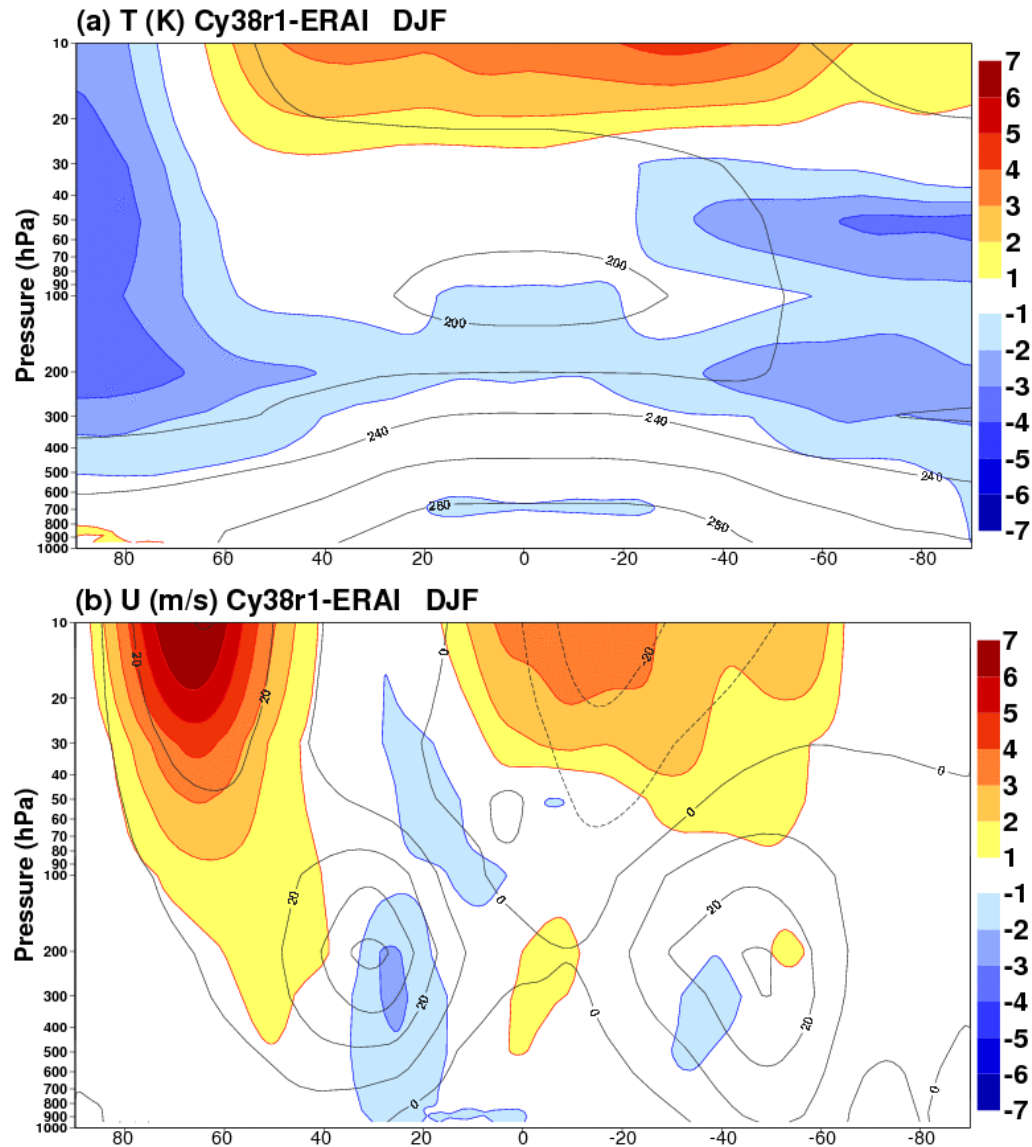
East Pacific: Mean wind & T in ON 2011



Analysis difference: control - GOES13 AMV denial

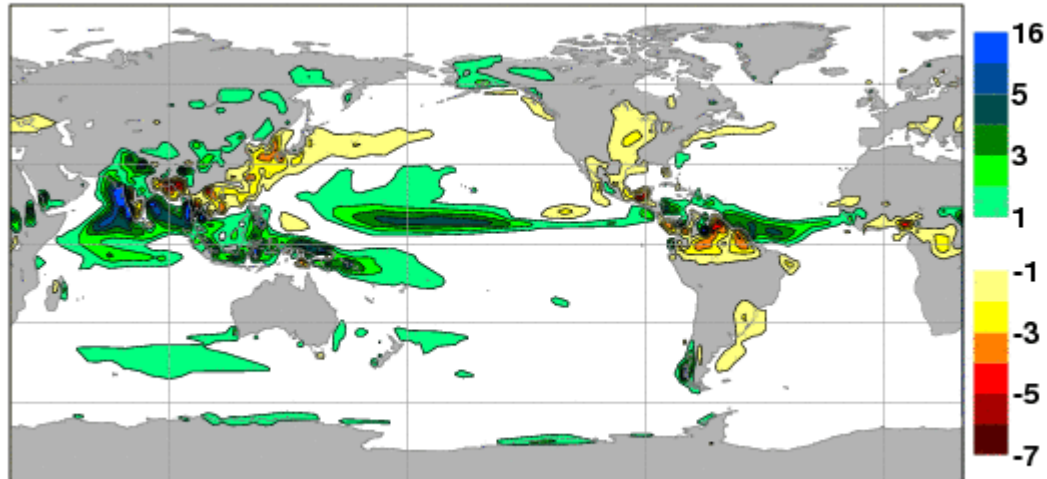


Zonal mean T and U errors in DJF *uncoupled*

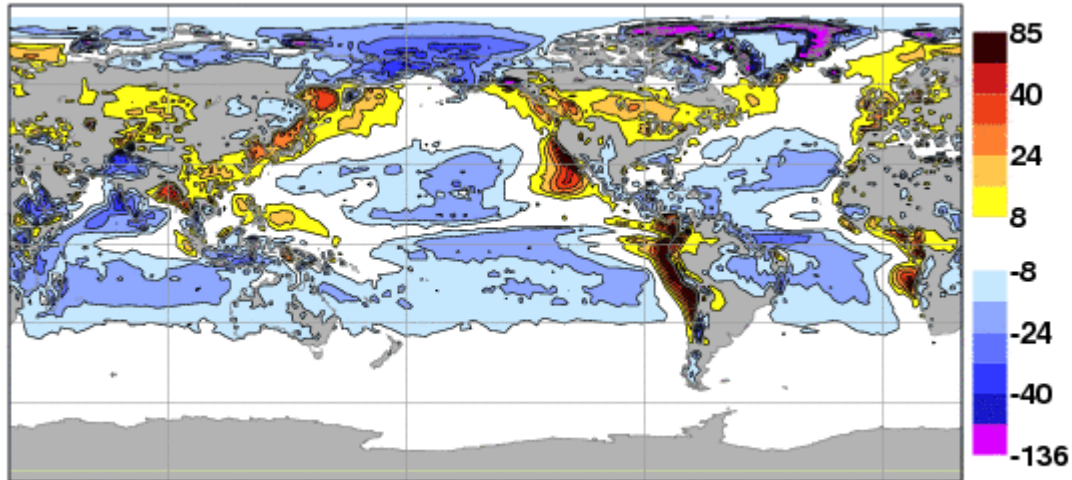


JJA Precip and SWnet errors *uncoupled*

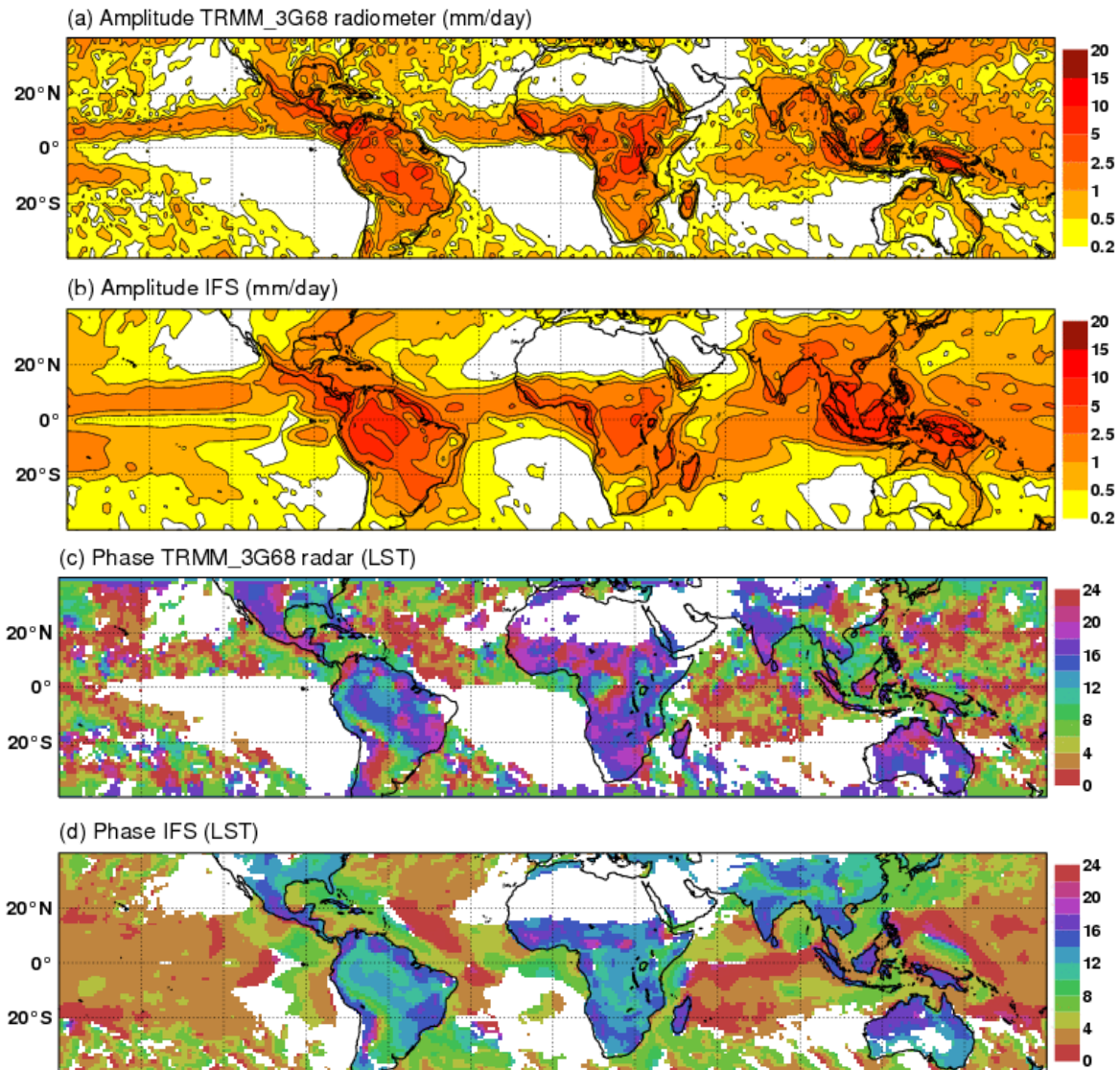
(a) JJA Precipitation difference with GPCP (mm/day)



(b) JJA top atmosphere SWrad difference with CERES

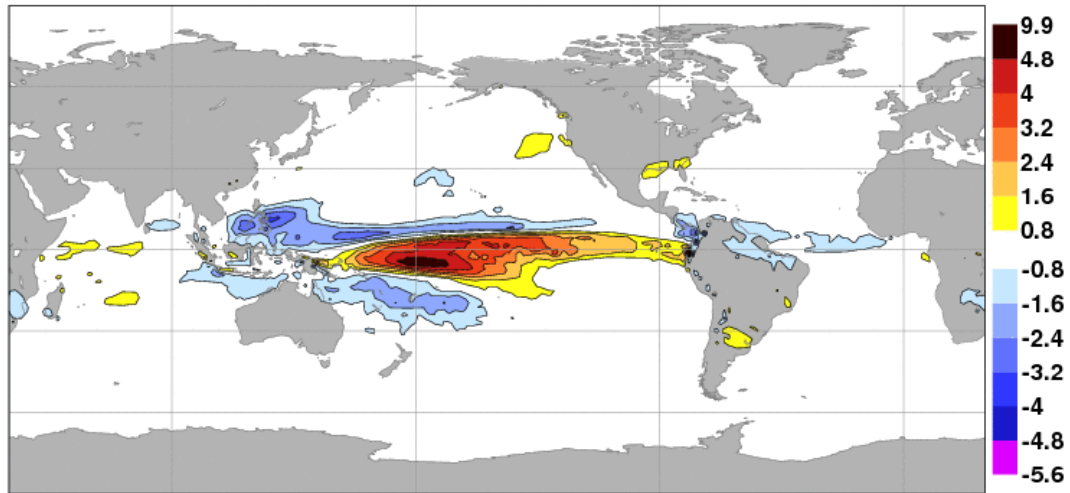


Diurnal cycle of Precipitation vs. TRMM

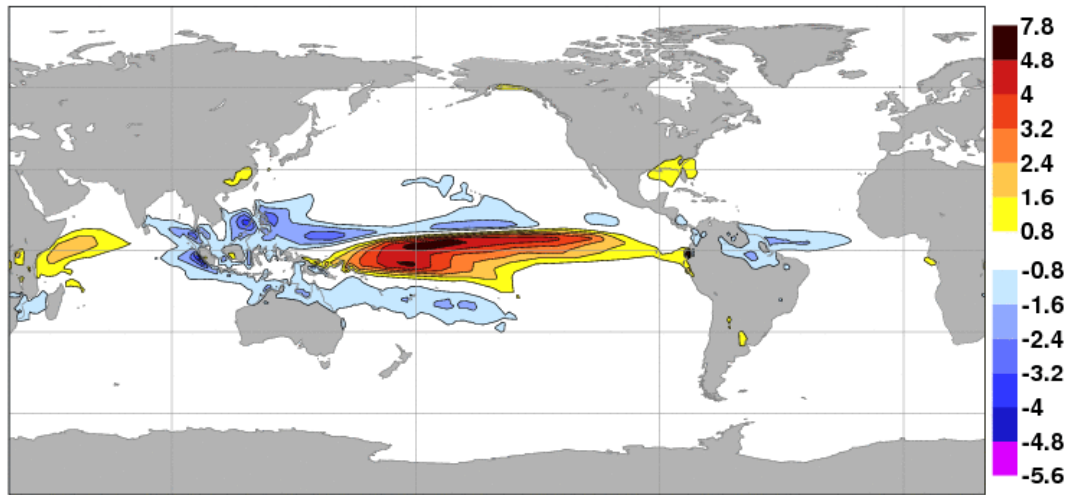


Covariance Nino3.4 SST & Prec. for DJF

(a) ERAI Covariance Nino3.4SST-Precip DJF

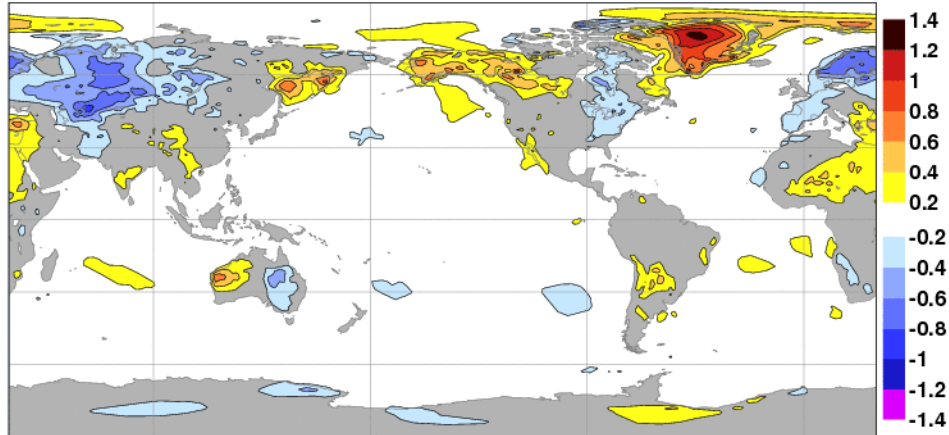


(b) Cy38r1 Covariance Nino3.4SST-Precip DJF

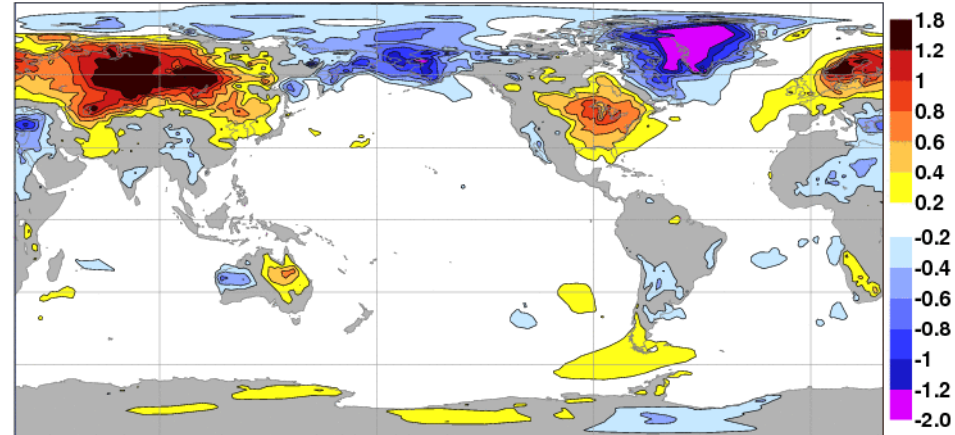


Teleconn. U10hPa Tropics & 2T for DJF

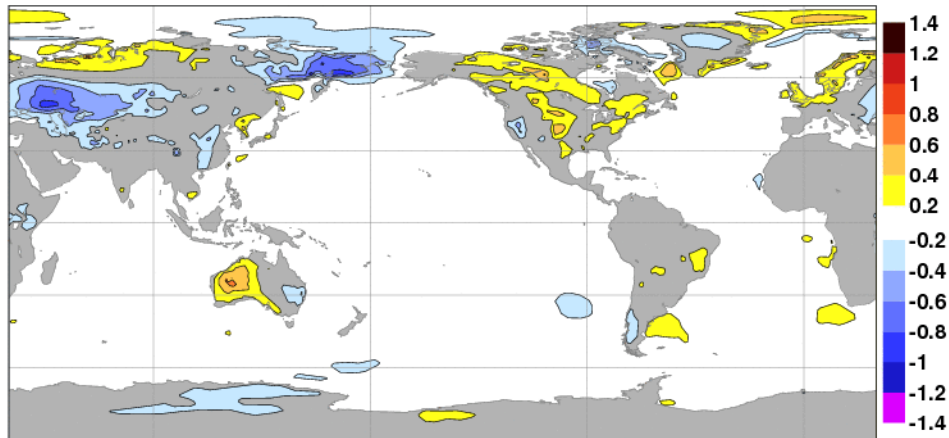
(a) ERAI Teleconnection -U10hPa-2T, 42 cases DJF



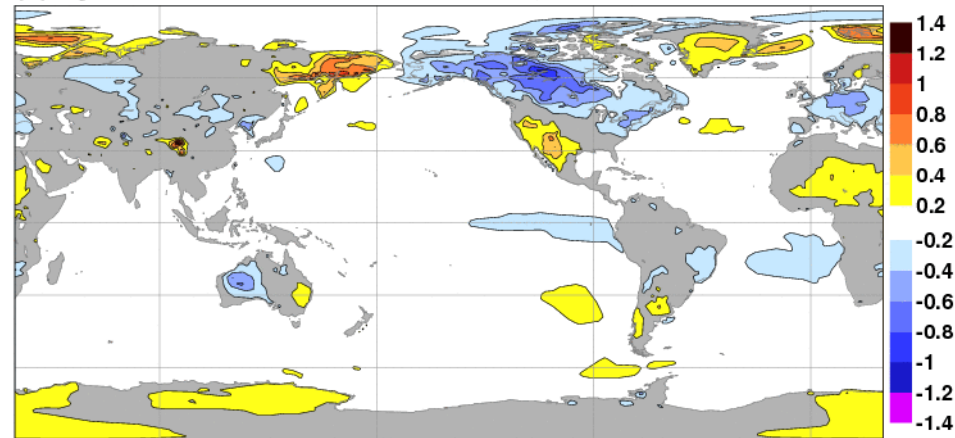
(c) ERAI Teleconnection +U10 hPa-2T, 37 cases DJF



(b) Teleconnection -U10 10hPa-2T, 105 cases DJF

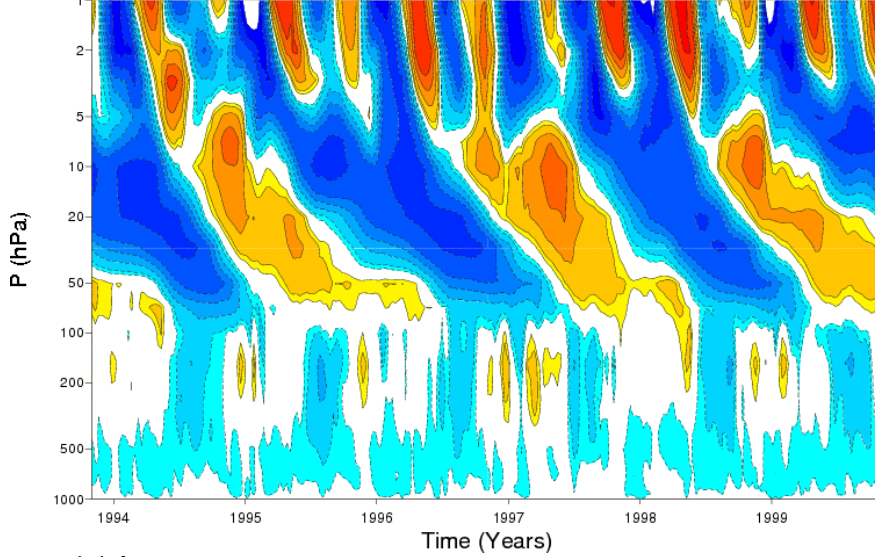


(d) Cy38r1 Teleconnection +U10 hPa-2T, 100 cases DJF

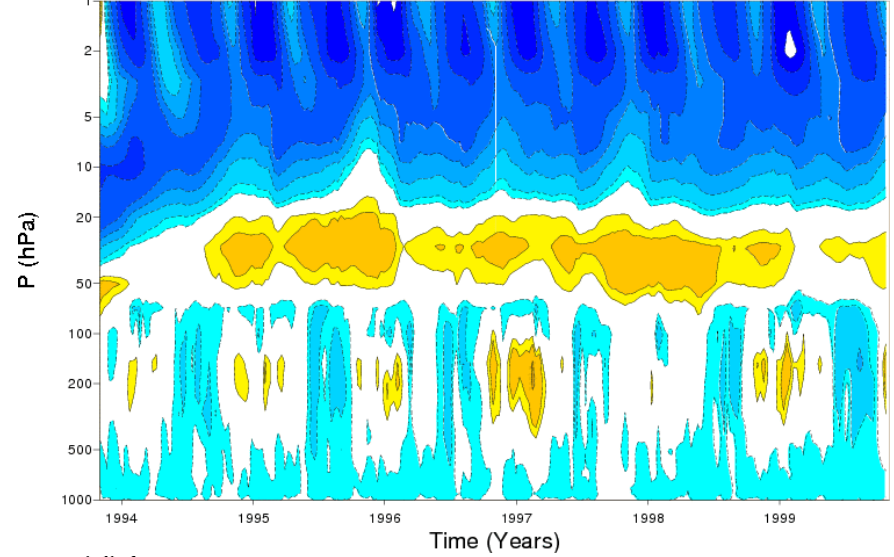


The QBO

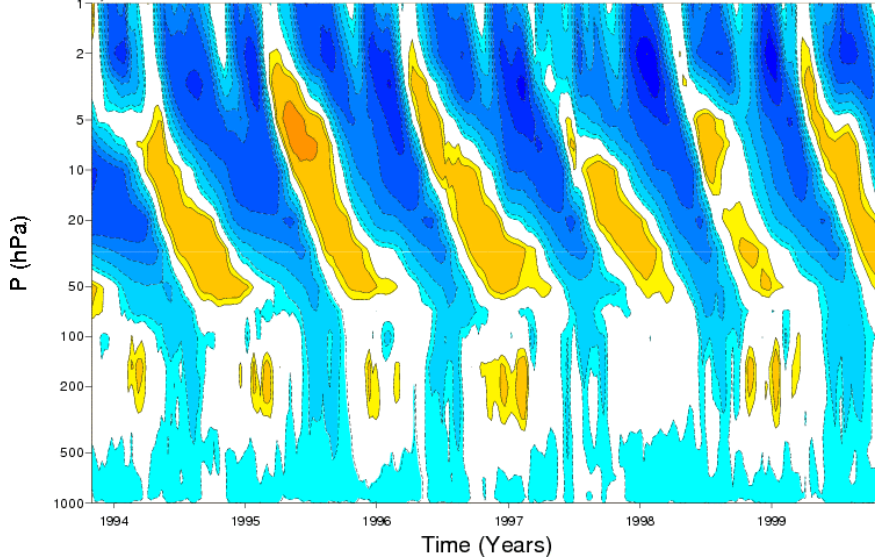
(a) ERAI



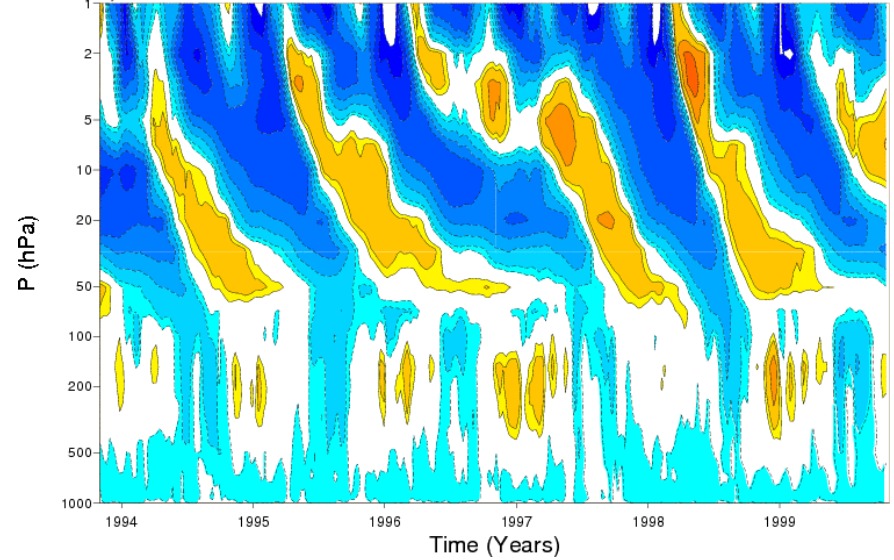
(b) L91noGWD



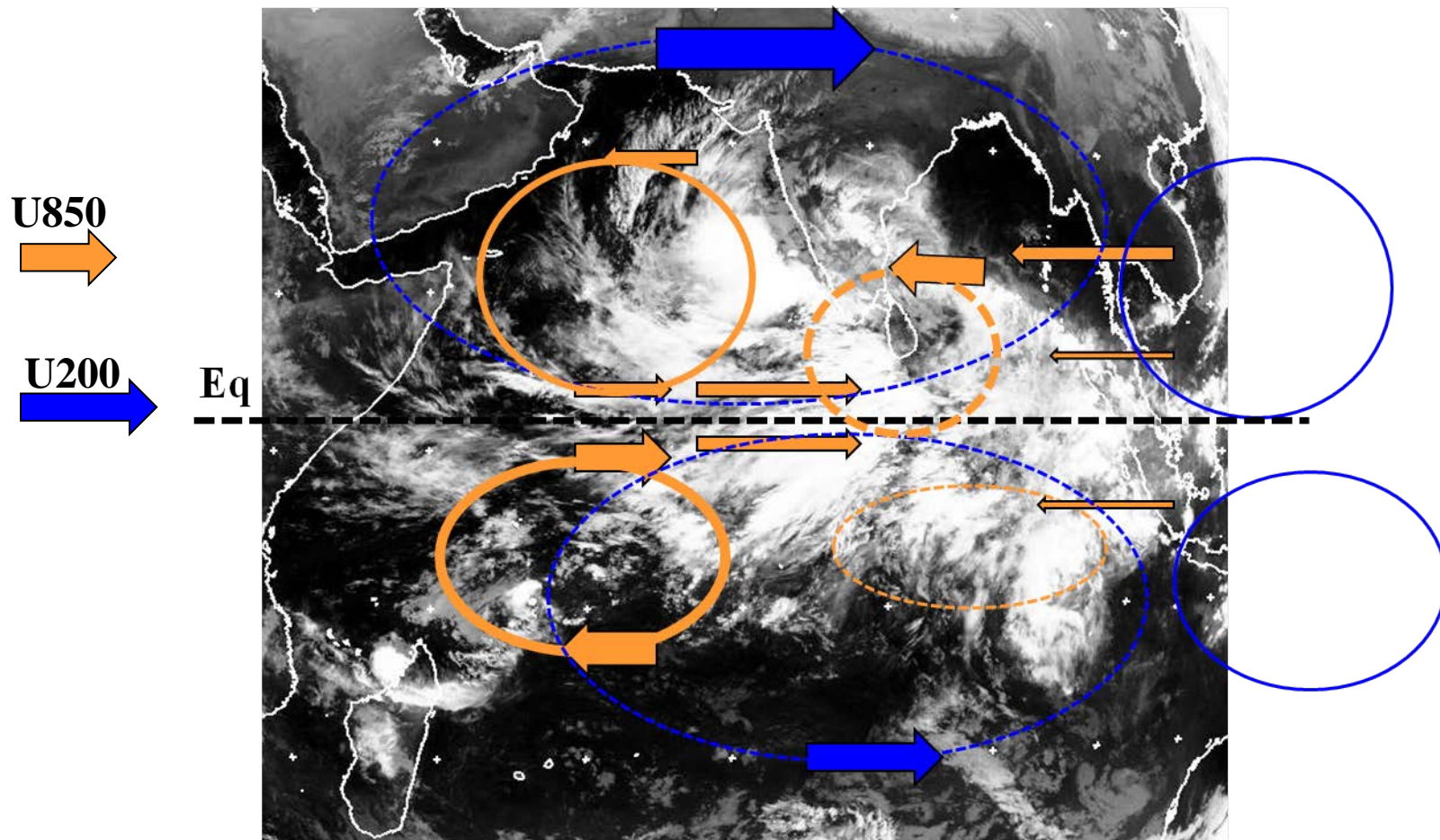
(c) L91



(d) L137

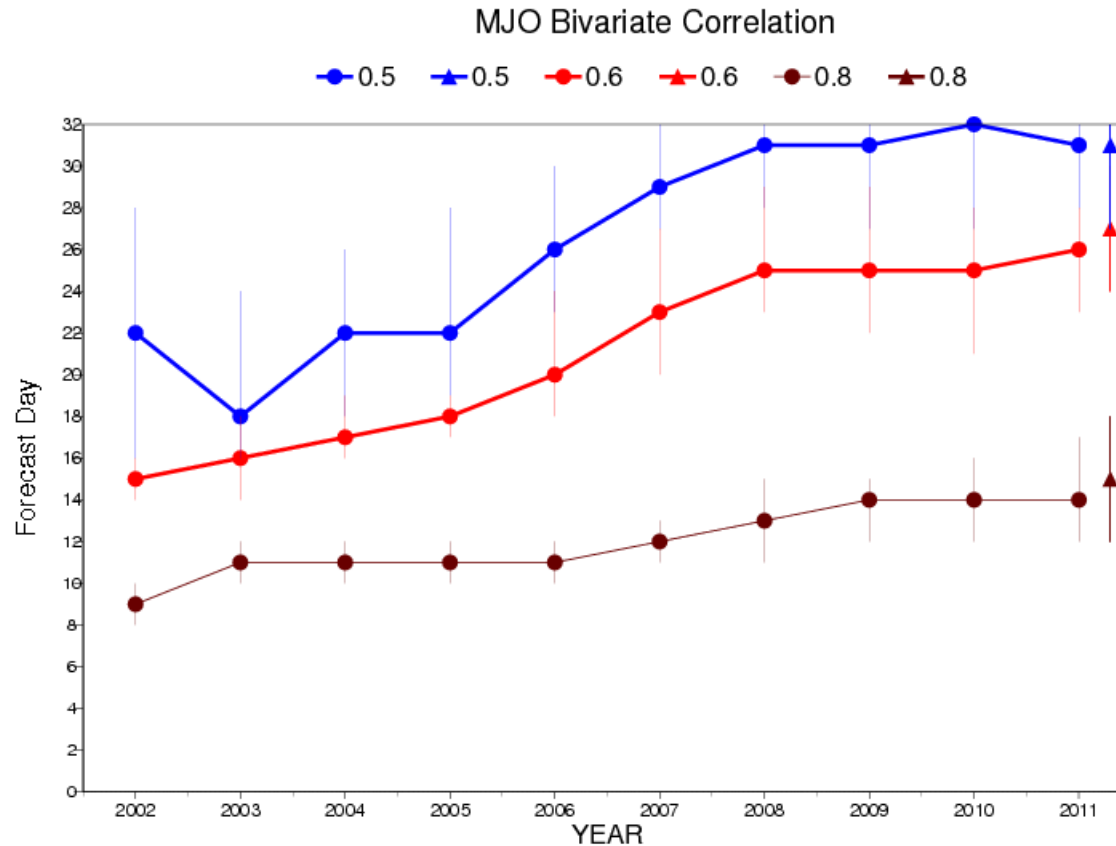


The MJO



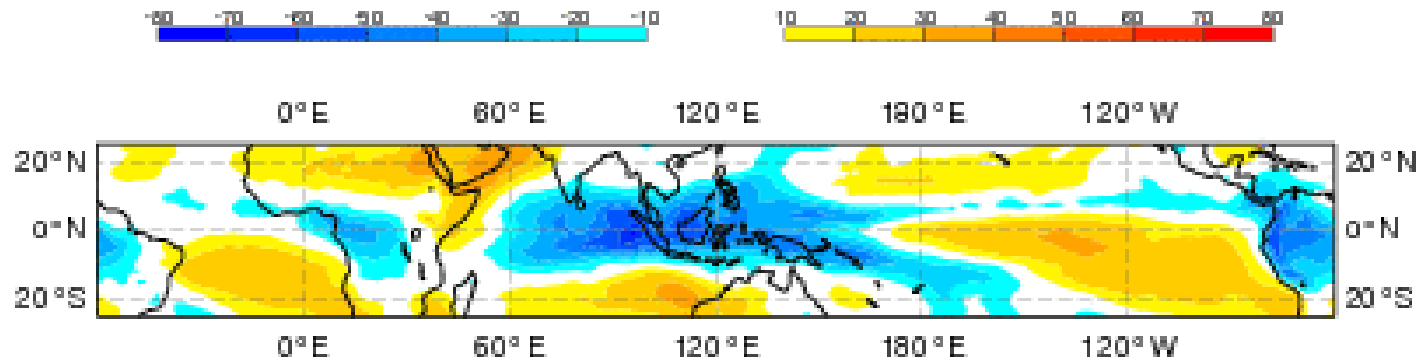
24 November 2011: Meteosat 7 + IFS Analysis

Progress in MJO prediction

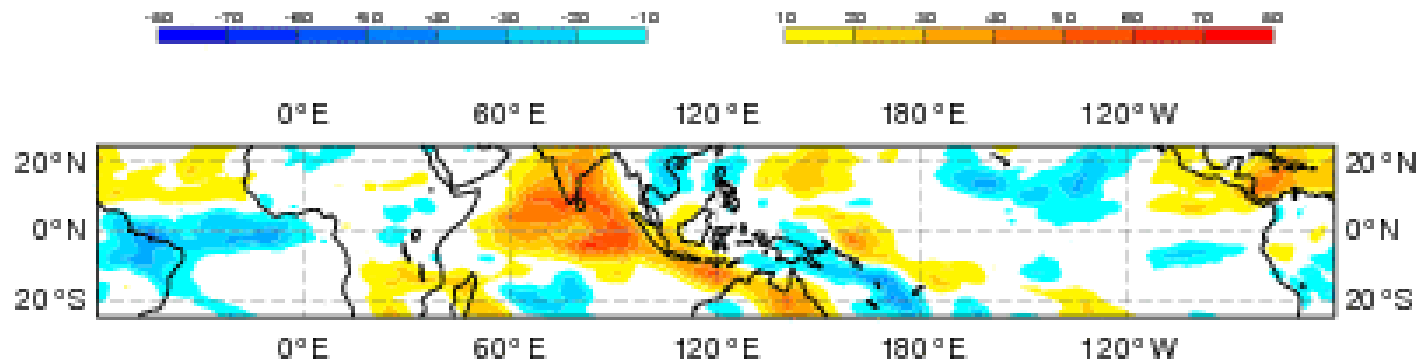


YOTC: OLR anomalies

(a) MJO Phase 2/3 36 h OLR anomaly

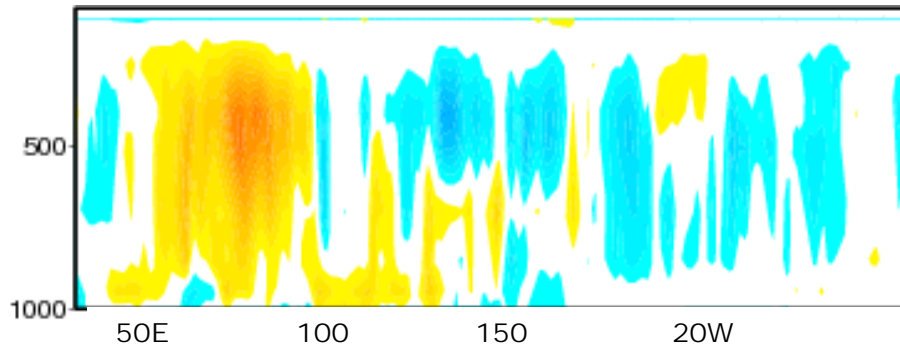


(b) MJO Phase 6/7-2/3 36 h OLR

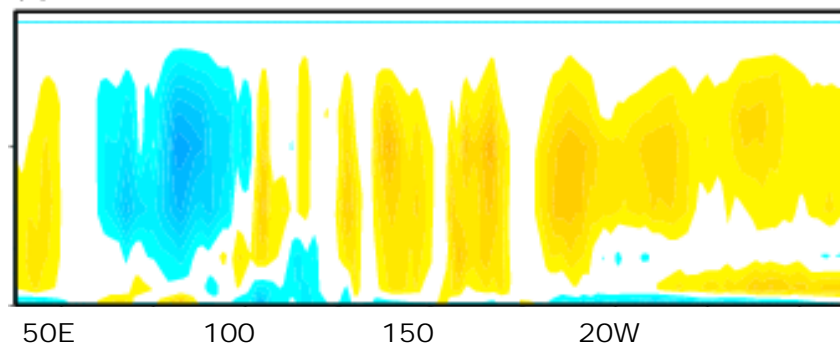


YOTC: Difference in T-tendency: Phase 6/7-Phase 2/3

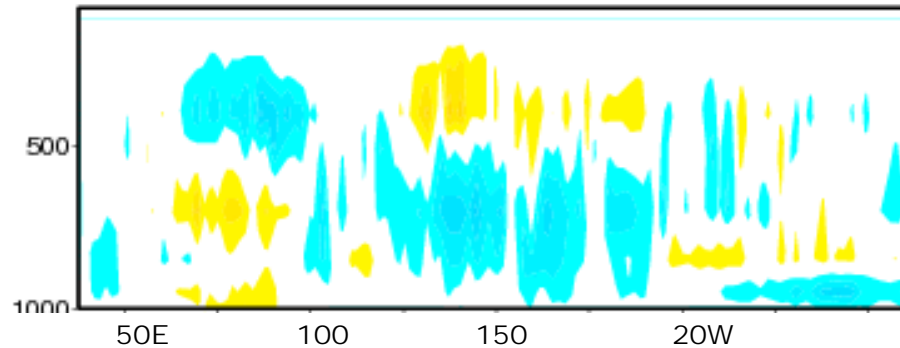
Dynamics (K/day)



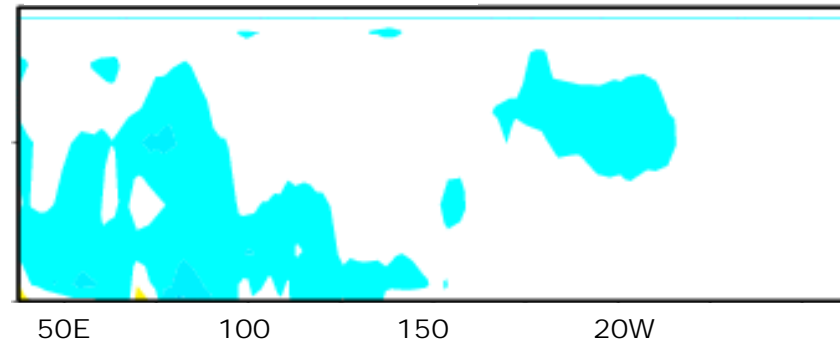
Conv (K/day)



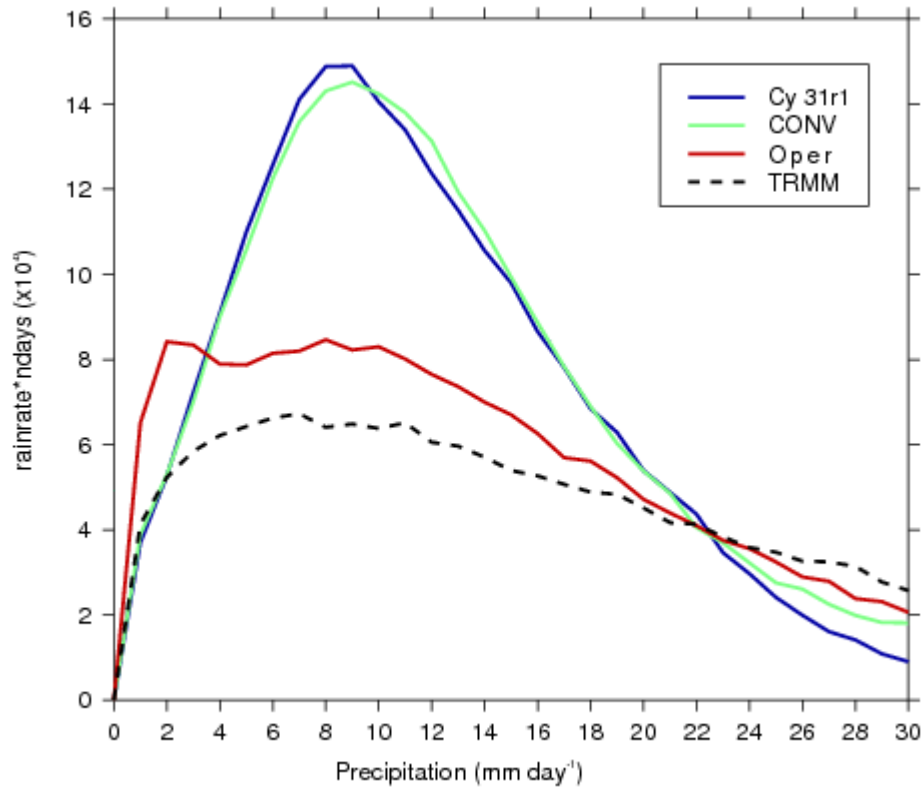
Cloud (K/day)



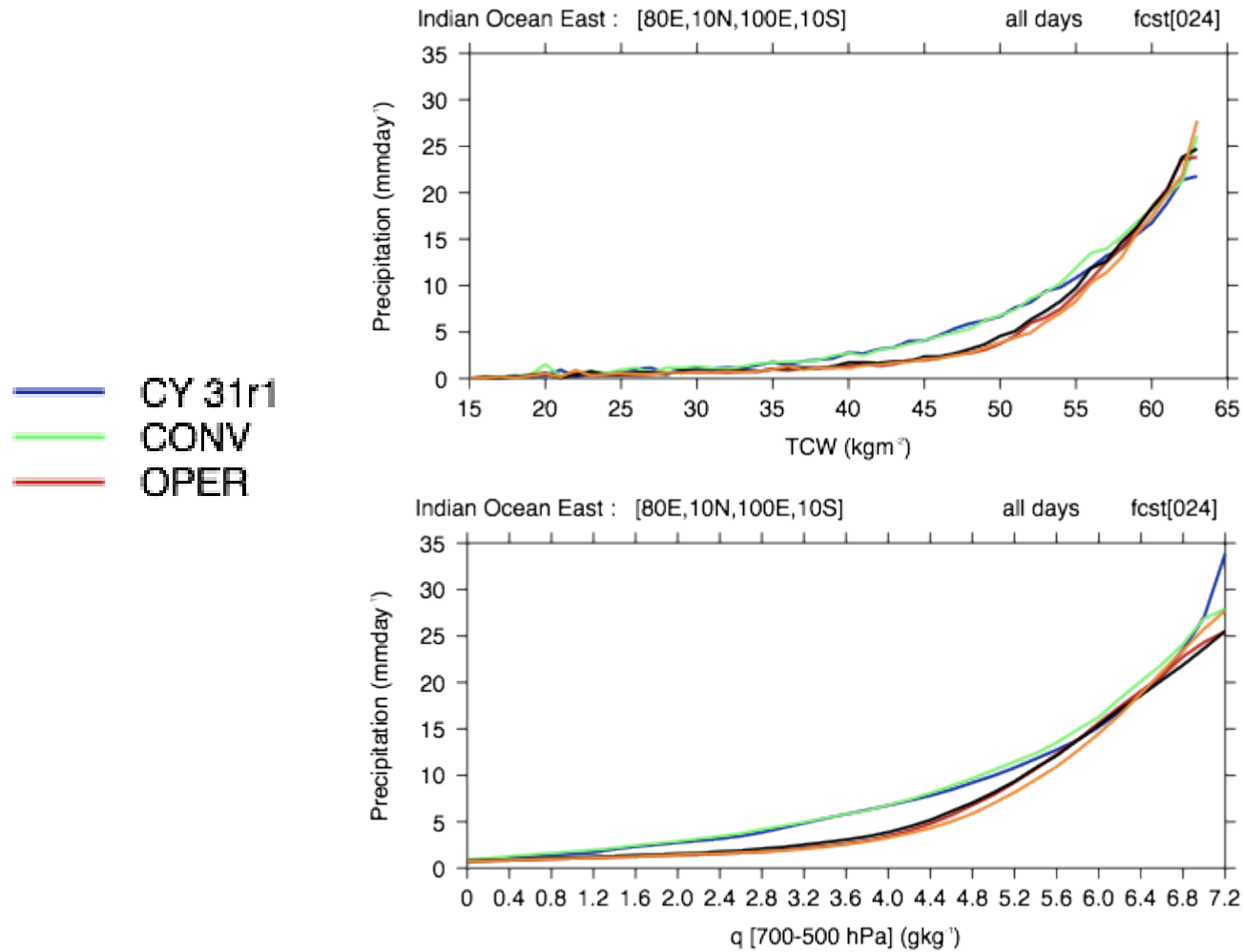
Radiation (K/day)



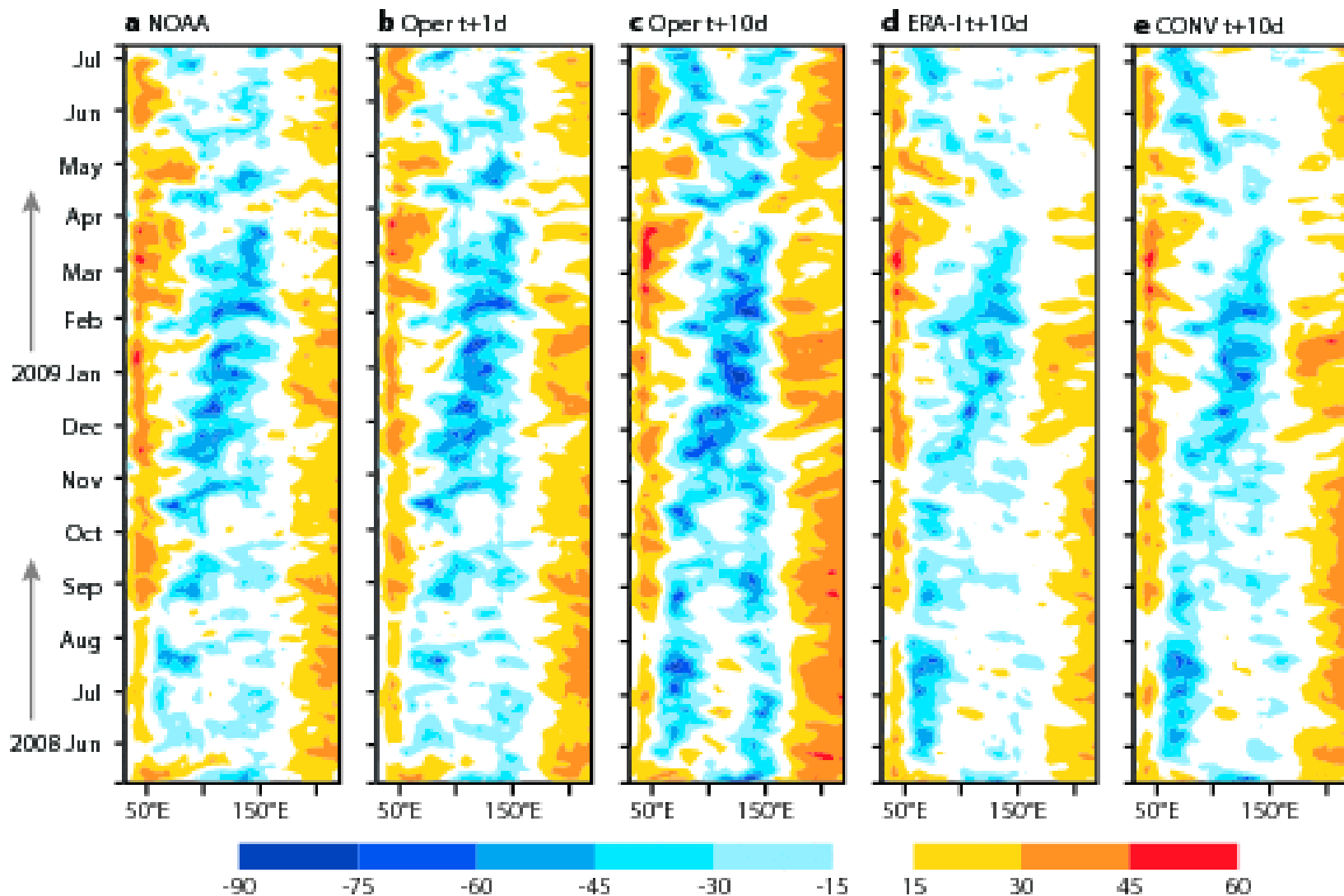
YOTC: Pdf of 24h Precipitation vs TRMM



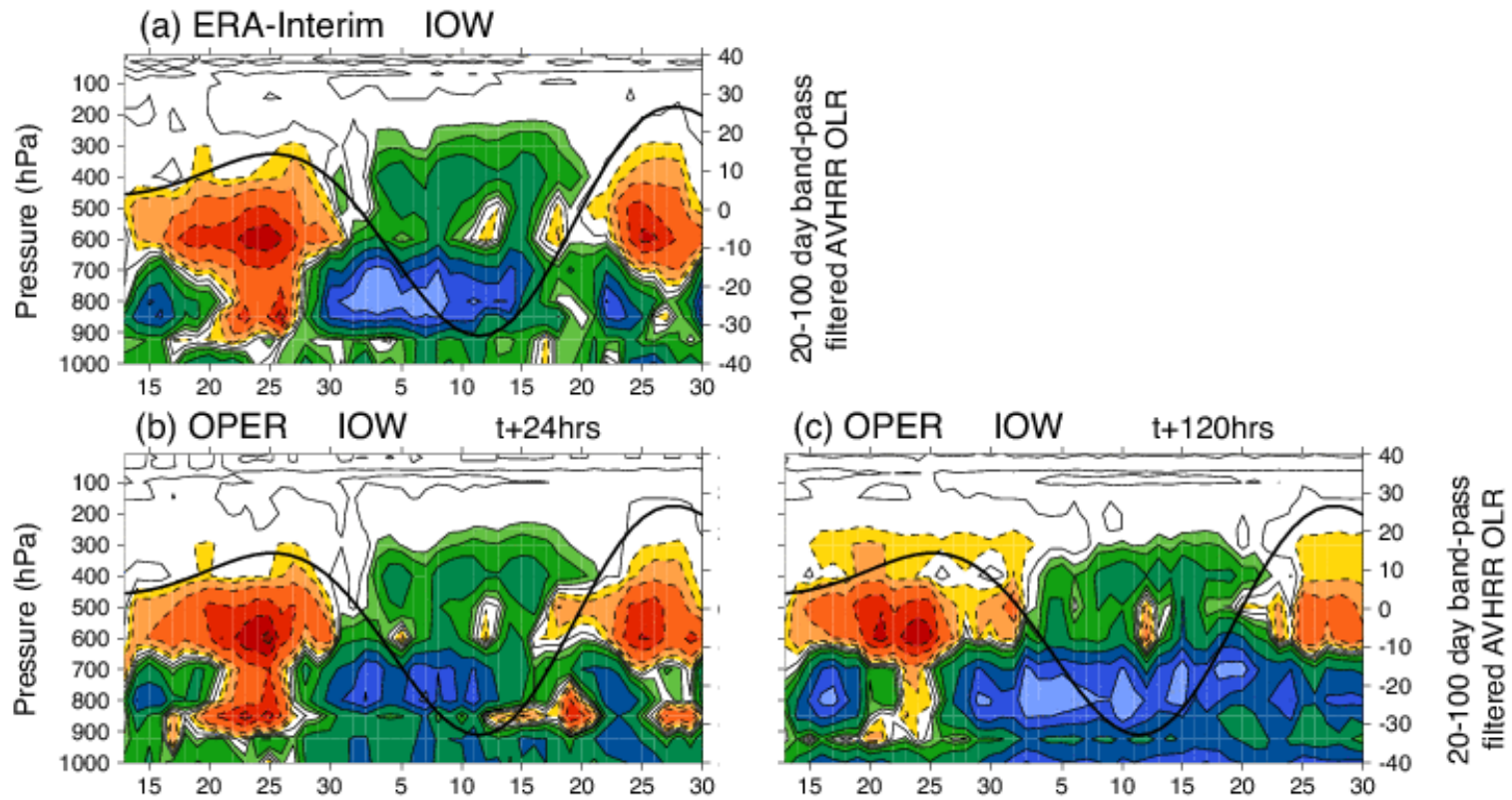
Indian Ocean: Precip vs TCW & RH



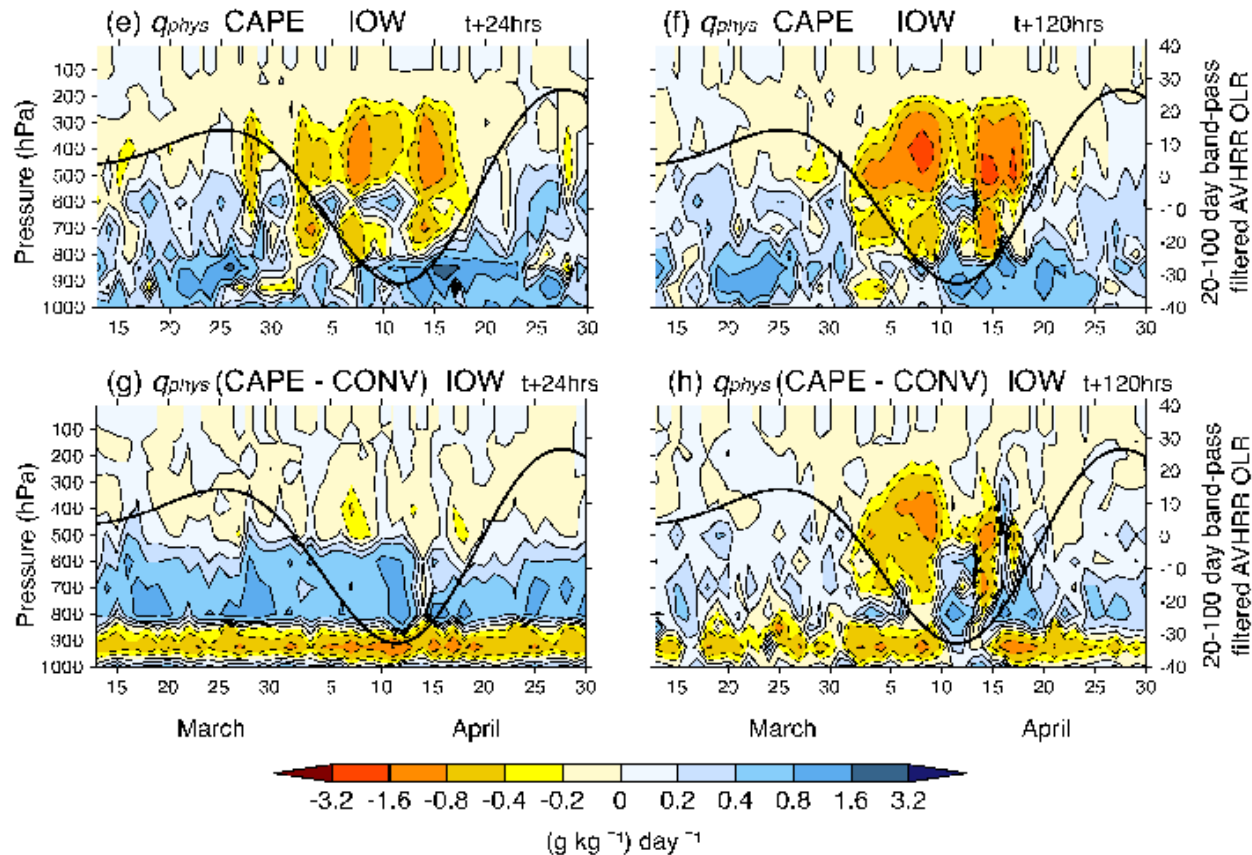
YOTC: Hovmoeller of the OLR anomaly



YOTC: MJO event of Indian Ocean time series OLR and q-anomalies

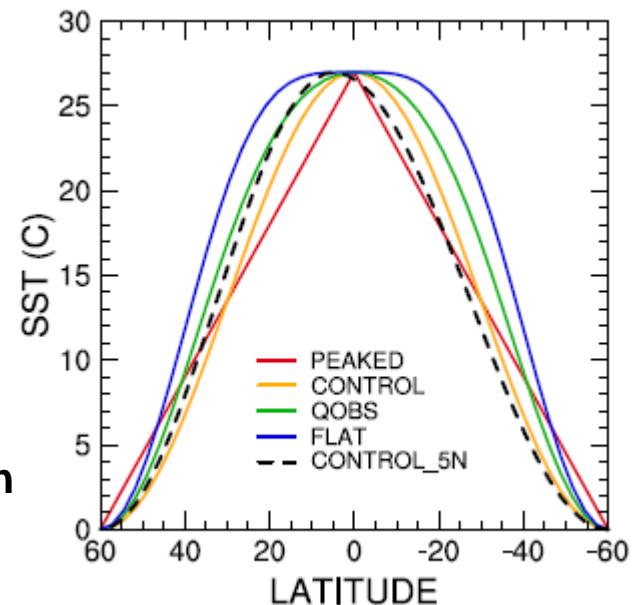


YOTC: Evolution of Indian Ocean OLR anomaly & dq/dt physics



Towards high resolution of global convection: on the AQUA Planet

- prescribed SST distribution
- Perpetual: Sun fixed over the Equator
- prescribed trace gaz concentrations
- **Run 4-member ensemble at T159 for 1-year**
- **start from a balanced state=6-month integration**



see: “The Aqua-Planet Experiment (APE): Control SST simulation” and
“Response to changed meridional SST profile” , **Blackburn et al. 2012, *J. Met. Soc. Japan***

Obs - real Earth - AQUA planet

TCW

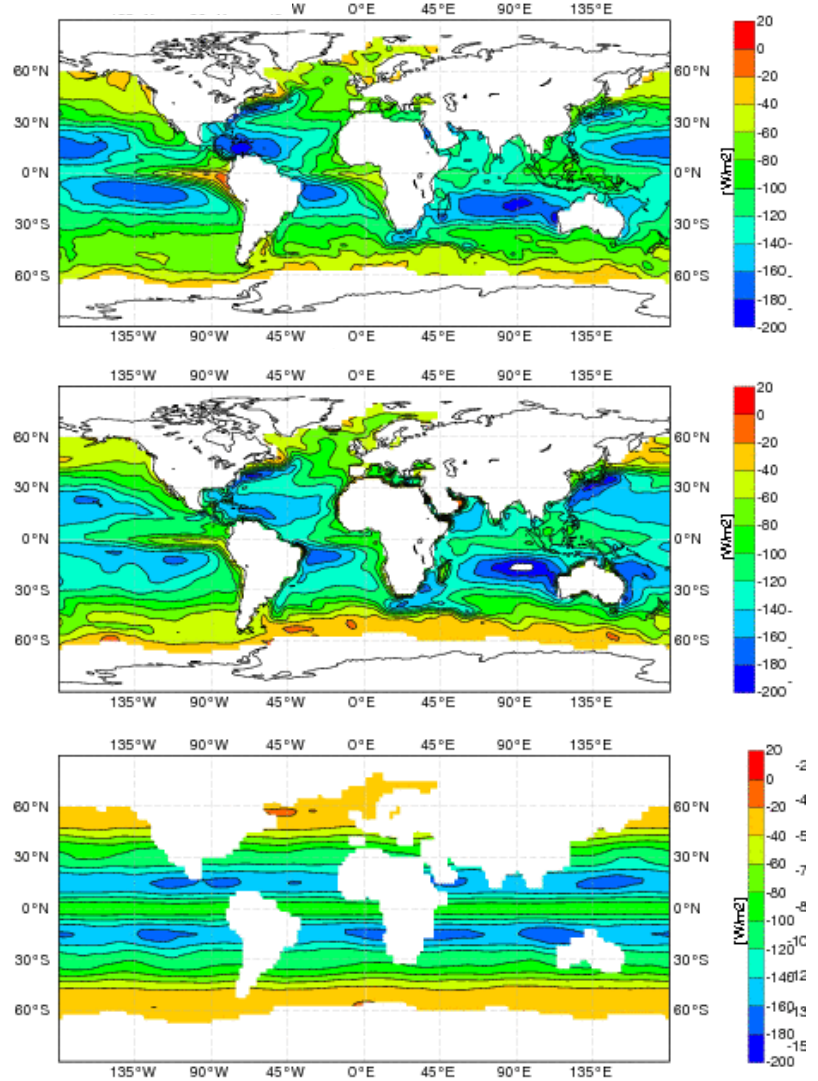
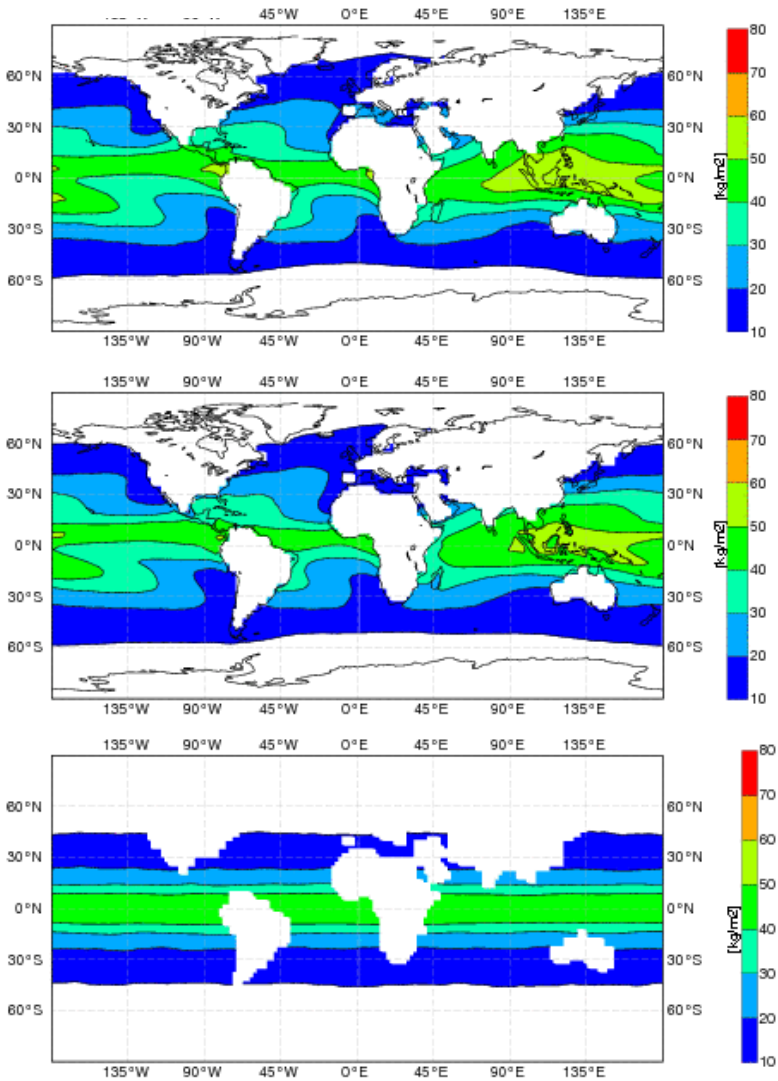
Lat. Heat

SSMI

HOAPS

Real Earth

AQUA



Obs - real Earth - AQUA planet

Precip

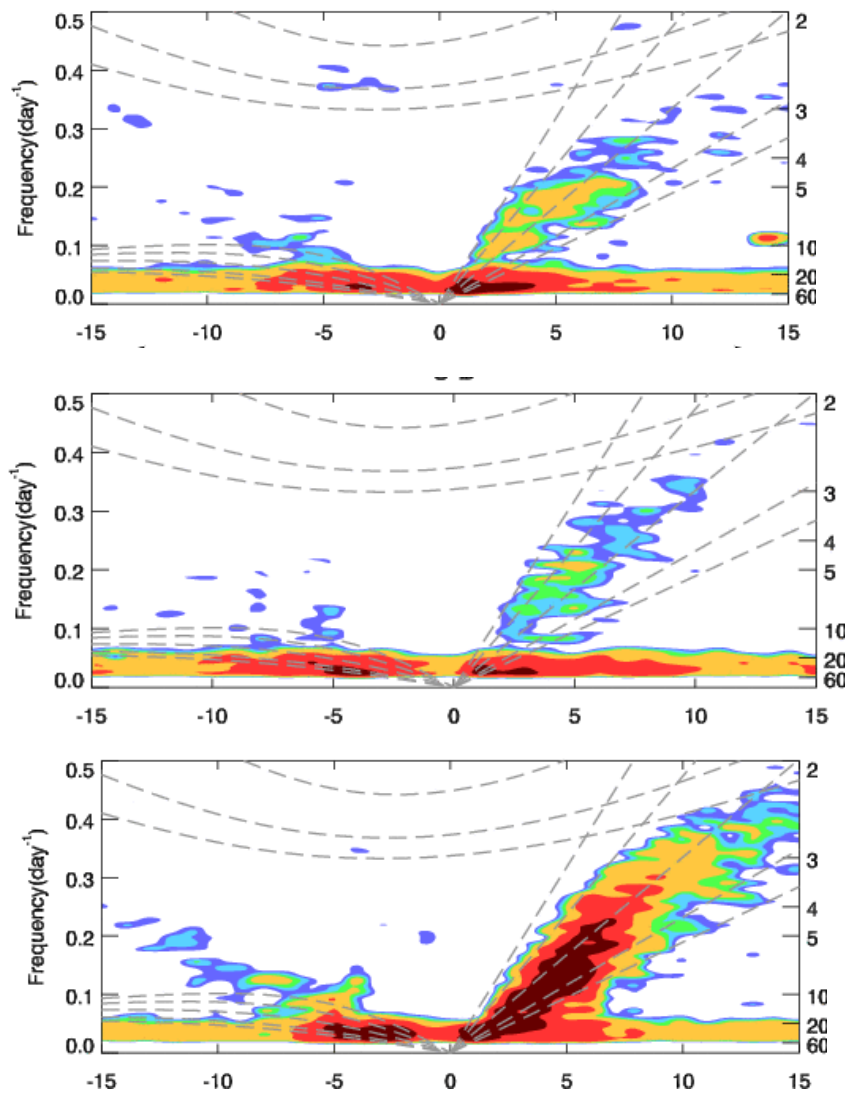
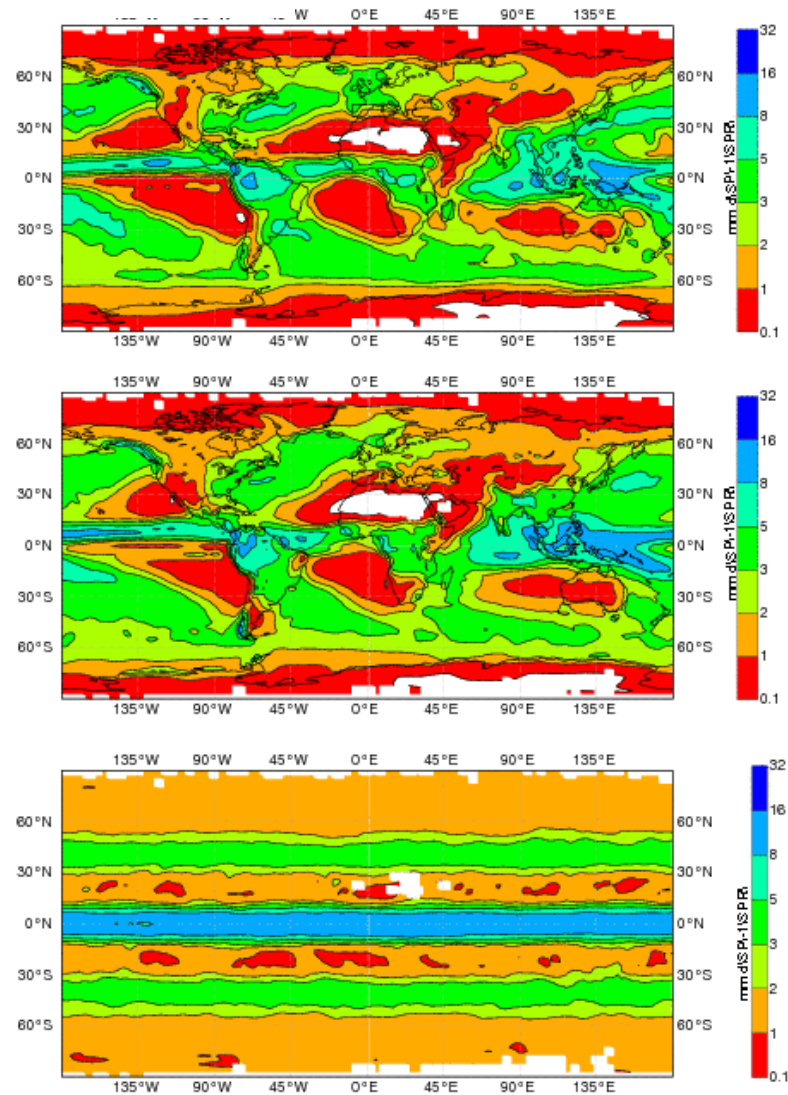
OLR Spectra

GPC2.2

NOAA

Real Earth

AQUA

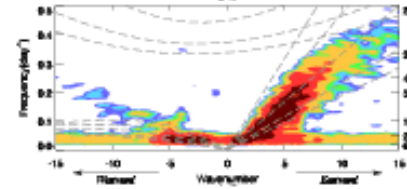
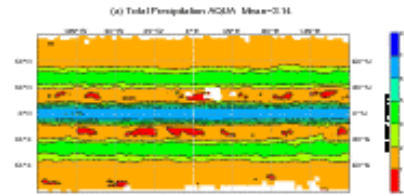


AQUA planet: resolution +/- conv

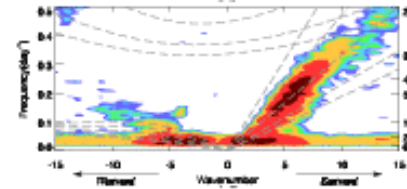
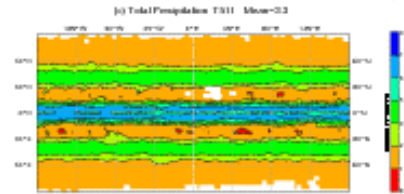
Precip

OLR Spectra

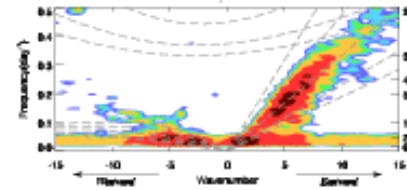
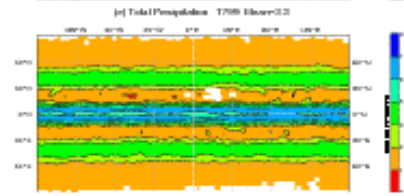
T159



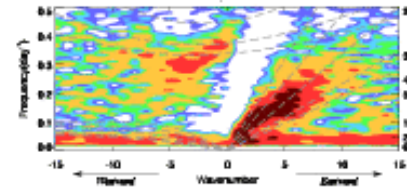
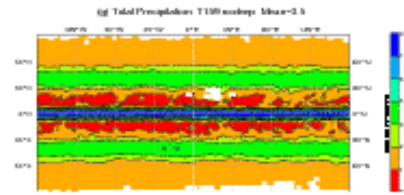
T511



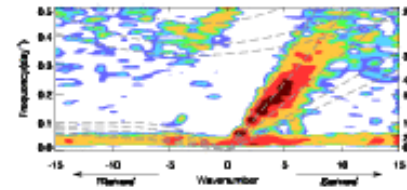
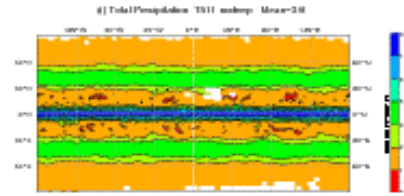
T799



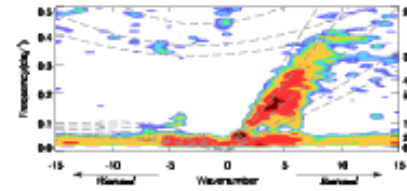
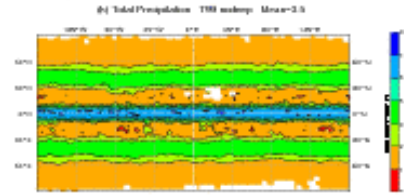
**T159
no deep**



**T511
no deep**



**T799
no deep**



The small Planet: horizontal length-scale and time scale

$$R_a' = R_a / \gamma_r$$
$$f' = f \gamma_t \Leftrightarrow t' = t / \gamma_t$$

$$Ro = \frac{U}{R_a f}$$



But

$$\frac{\partial u}{\partial x'} + \frac{\partial v}{\partial y'} = \gamma_r \left(\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} \right) = -\frac{\partial w}{\partial z} = -\frac{\partial \omega}{\partial p} \quad ?$$

also thanks to N. Wedi and S. Malardel for earlier versions of small planet

The small Planet: vertical scaling

$$g' = g \gamma_g; \quad H = \frac{RT}{g} \rightarrow H' = H / \gamma_g$$

$$\gamma_r = \gamma_t = \gamma_g$$



$$Ri = g \frac{\delta T}{\bar{T}} \frac{H}{U^2} = N^2 \frac{H^2}{U^2}; \quad c = \sqrt{g H}$$



or Lamb
parameter

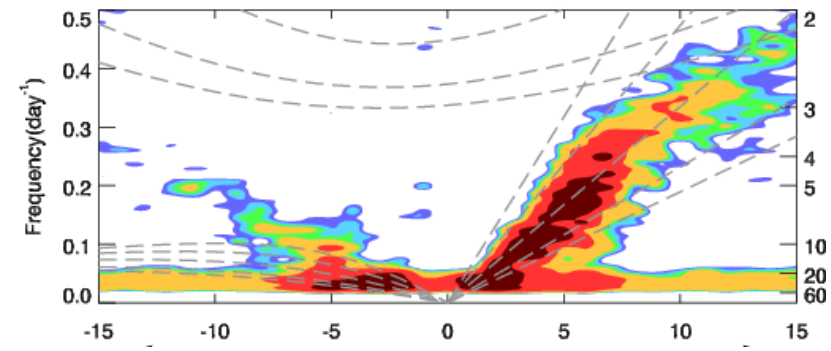
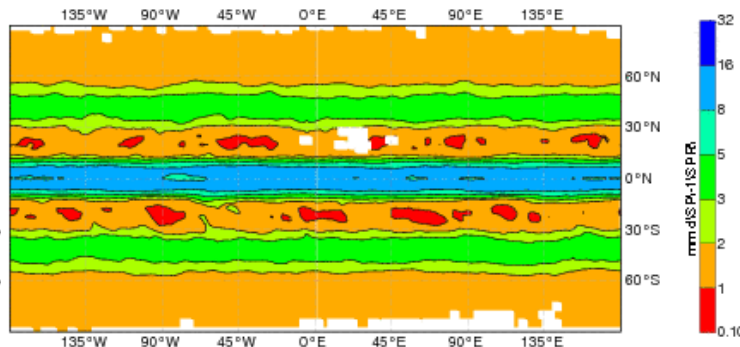
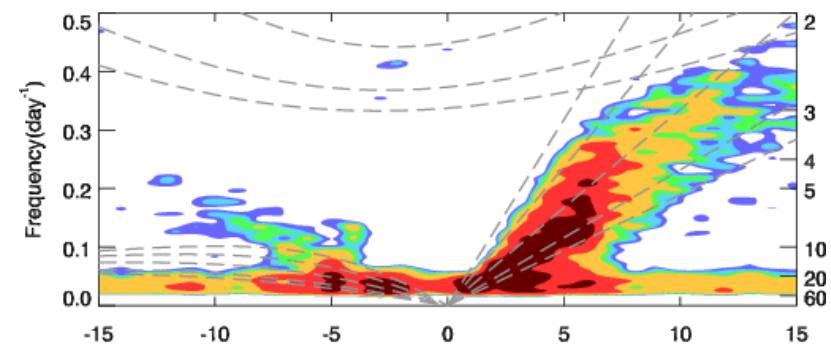
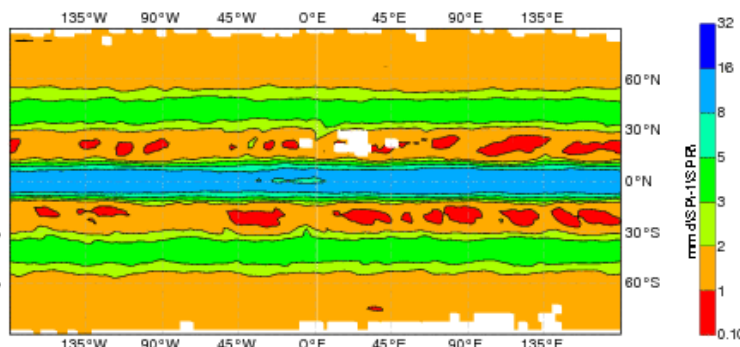
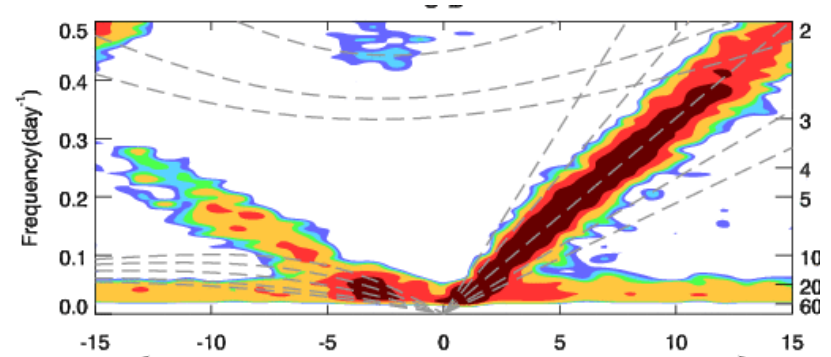
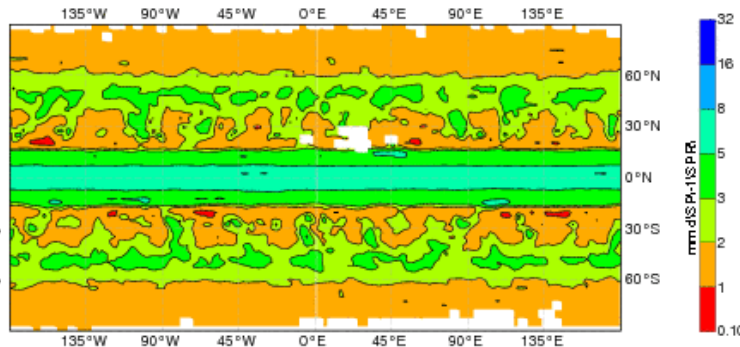
Physics scaling



AQUA planet: $f=t$ & g scaling

Precip

OLR Spectra



R/4
t/4

R/4
t/4
 g^*4
physics

R/8
t/8
 g^*4
physics

Conclusions

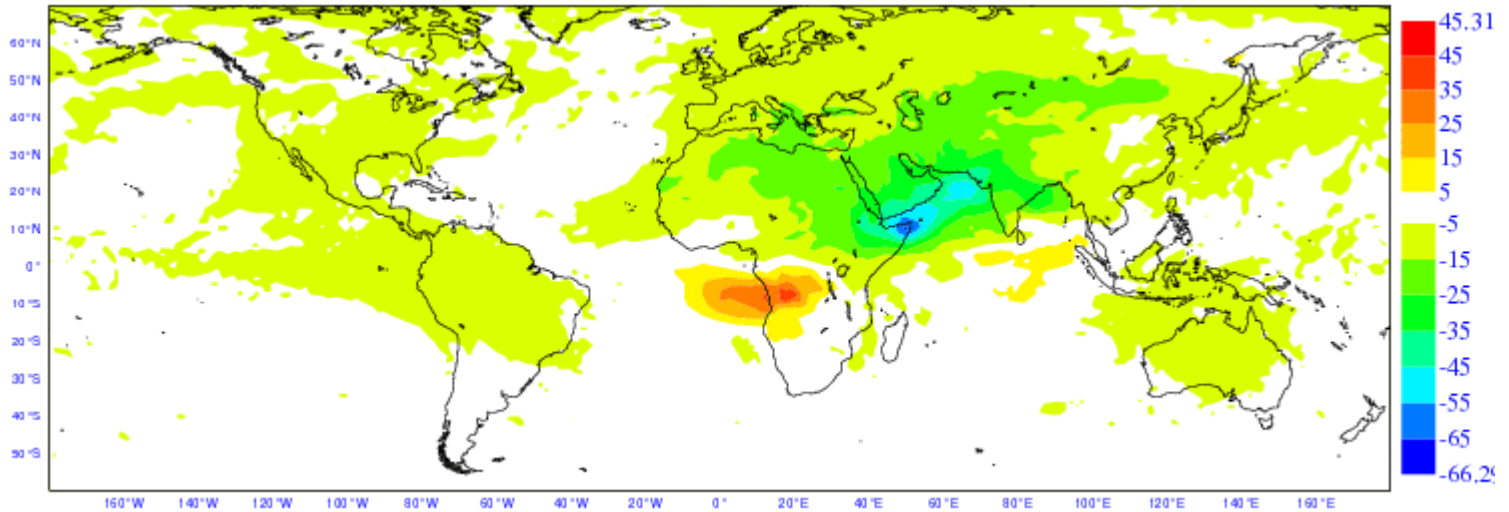
- Most important area for energy generation and conversion is the upper-tropical troposphere
- Convective heating must occur in the right phase of the large-scale wave, and as T variations are small must show the right sensitivity to mid-tropospheric moisture
- Main forecast errors concern spindown of Hadley cell (why?), and a too strong SE Asian Monsoon - moistening, easterly wind bias
- Largest low-level wind errors are in the East tropical Pacific where also analysis uncertainty is relatively large for 950-700 hPa winds
- with small Aqua planet nice forecast tool for global convection/waves=scalable Prototype for different planets

Any further

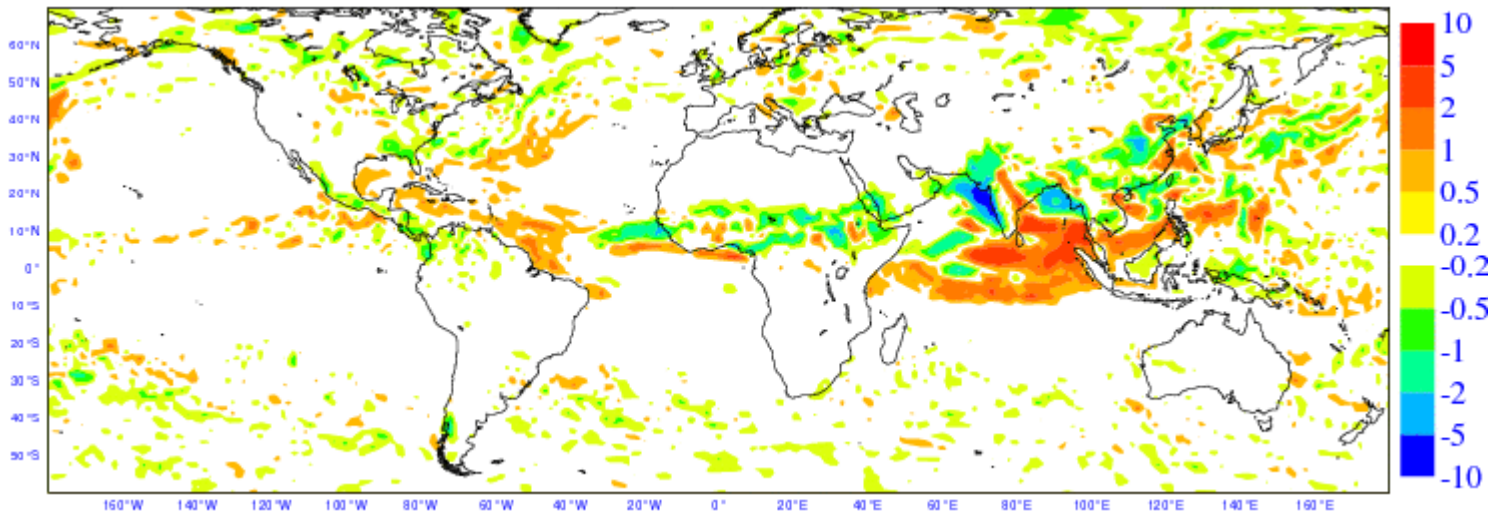
- Diurnal cycle will probably remain problem, address Monsoon heating, slightly overestimated shallow transport (trade Cu, dry PBL) though necessary for predictability, drizzle
- Major analysis impact expected from ADM-AEOLUS wind lidar
- and ???

Uncertainty in SW flux diverg. due to Aerosols.

Div
W/m²



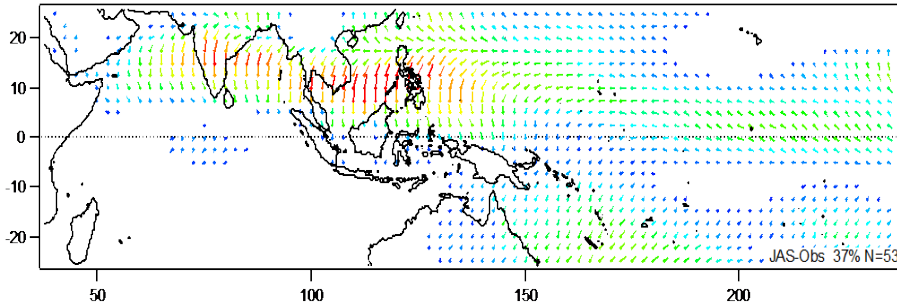
CP
mm/d



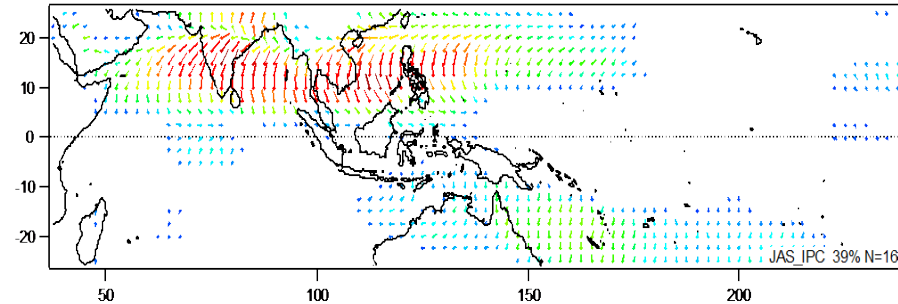
Nota: 10 W/m² ~ 0.35 mm/day

Intraseasonal variability by events

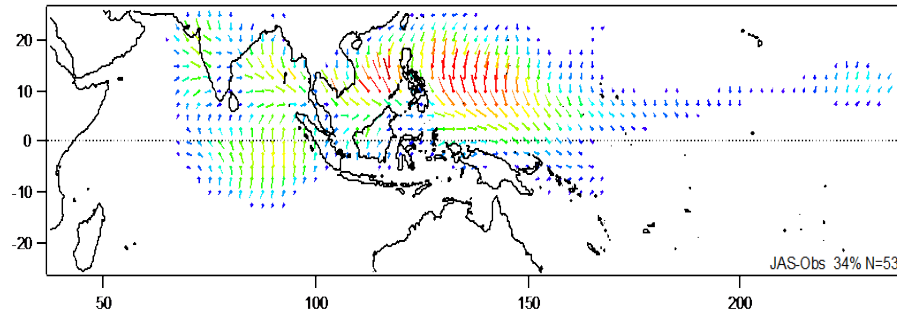
JAS Obs 30 years for U 850 hPa



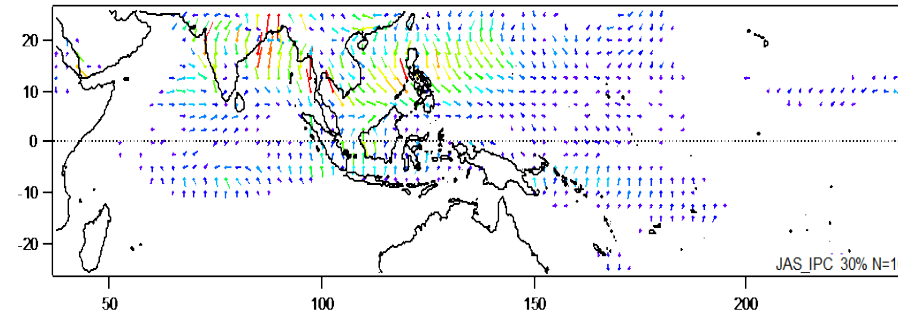
IFS 8x1-year for U 850 hPa



JAS Obs 30 years for Rain



IFS 8x1-year for Rain



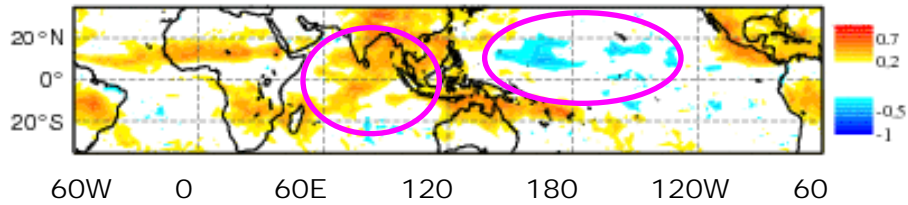
courtesy J.Ph. Duvel see also their (Clim Dyn. 2012)

Correlations with T' at 250 hPa for Phase 2/3 and forecast steps 12-36

Oper

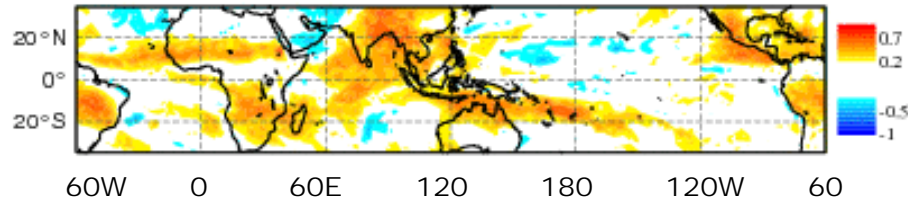
dT/dt_conv

MJOcomp Phase 2/3 250 hPa Corr T-dTconv Nos:103 rms=0.182 ops



Precip

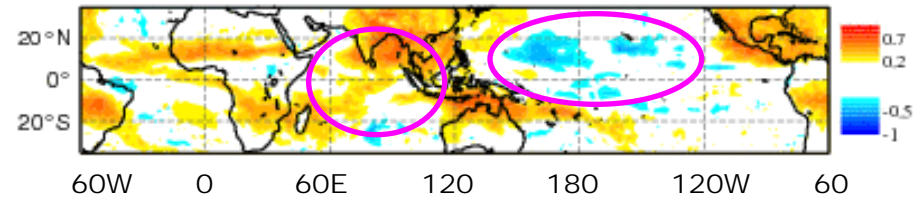
MJOcomp Phase 2/3 250 hPa Corr T-precip Nos:103 rms=0.206 ops



Old Conv

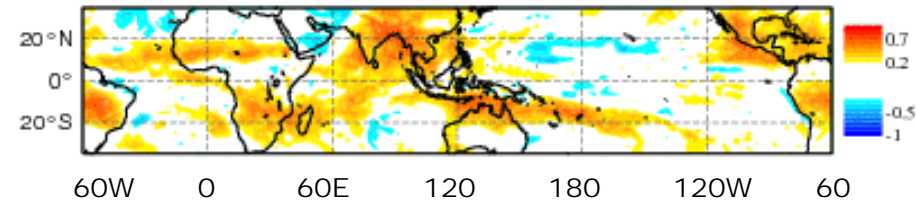
dT/dt_conv

MJOcomp Phase 2/3 250 hPa Corr T-dTconv Nos:103 rms=0.205 oldconv

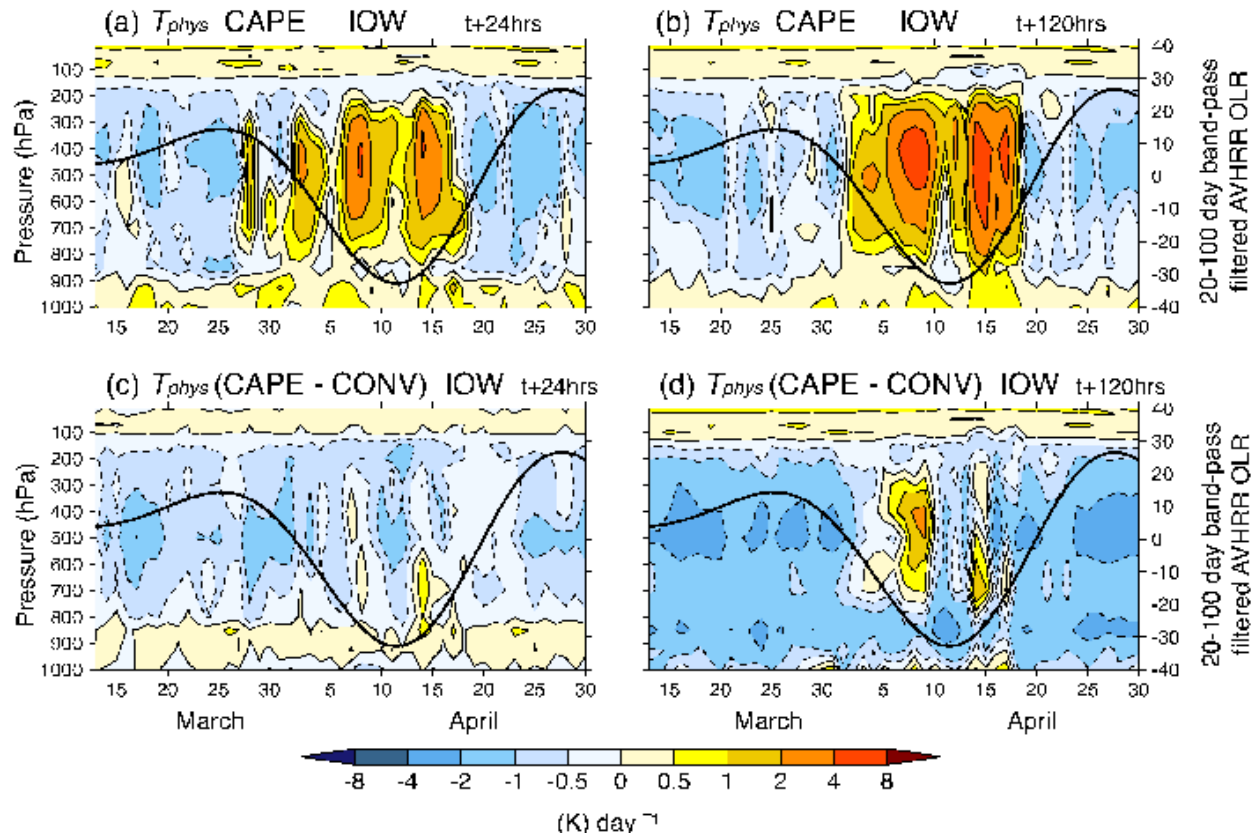


Precip

MJOcomp Phase 2/3 250 hPa Corr T-precip Nos:103 rms=0.209 oldconv



YOTC: Evolution of Indian Ocean OLR anomaly & dT/dt physics



YOTC: MJO forecasts April 2009: phase and amplitude as function of Fc lead time

