

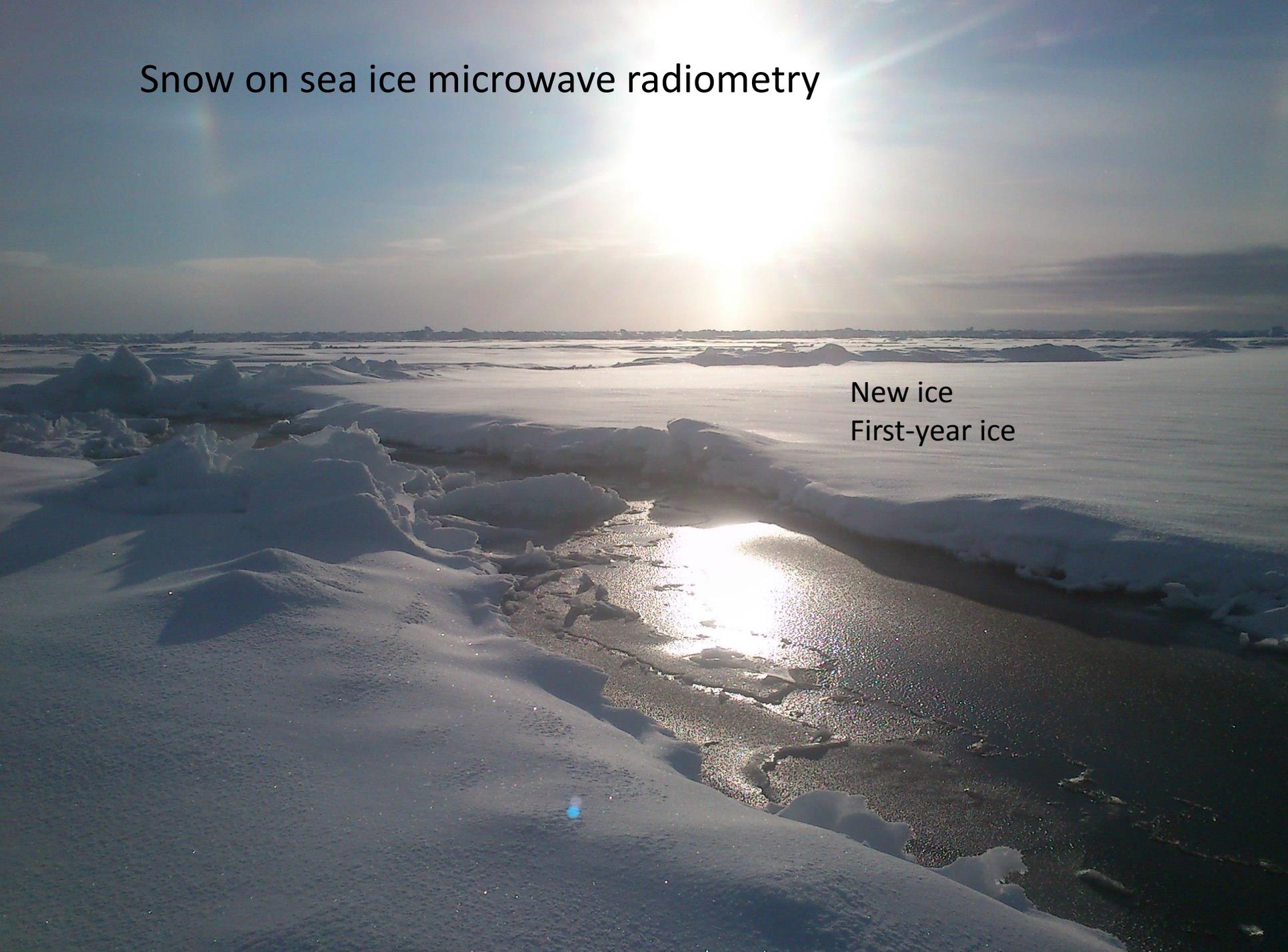
# Snow on sea ice retrieval using microwave radiometer data

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Danish Meteorological Institute  
Observatoire de Paris

# We know something about snow

- The temperature gradient within the snow is approximately linear and also in the ice it is linear. The temperature gradient in the snow is much larger than in the ice because the thermal conductivity in the snow ( $0.3 \text{ W/mK}$ ) is much lower than for ice ( $2.1 \text{ W/mK}$ ) by a factor of about 7. The snowpack metamorphosis is a function of the temperature gradient, the snow density and snow grain-size.
- After snowfall where the new snow deposit density is determined by temperature and wind speed compaction sets in driven by the load of the snowpack itself and recrystallisation processes, temperature and density of the snowpack.
- The thermal conductivity depend on the snow density and snow temperature and in ice on the ice salinity and temperature.
- Snow is a layered medium: In microwave radiometry it is very important to capture the layering in the snow in order to simulate realistic signatures.

# Snow on sea ice microwave radiometry



New ice  
First-year ice



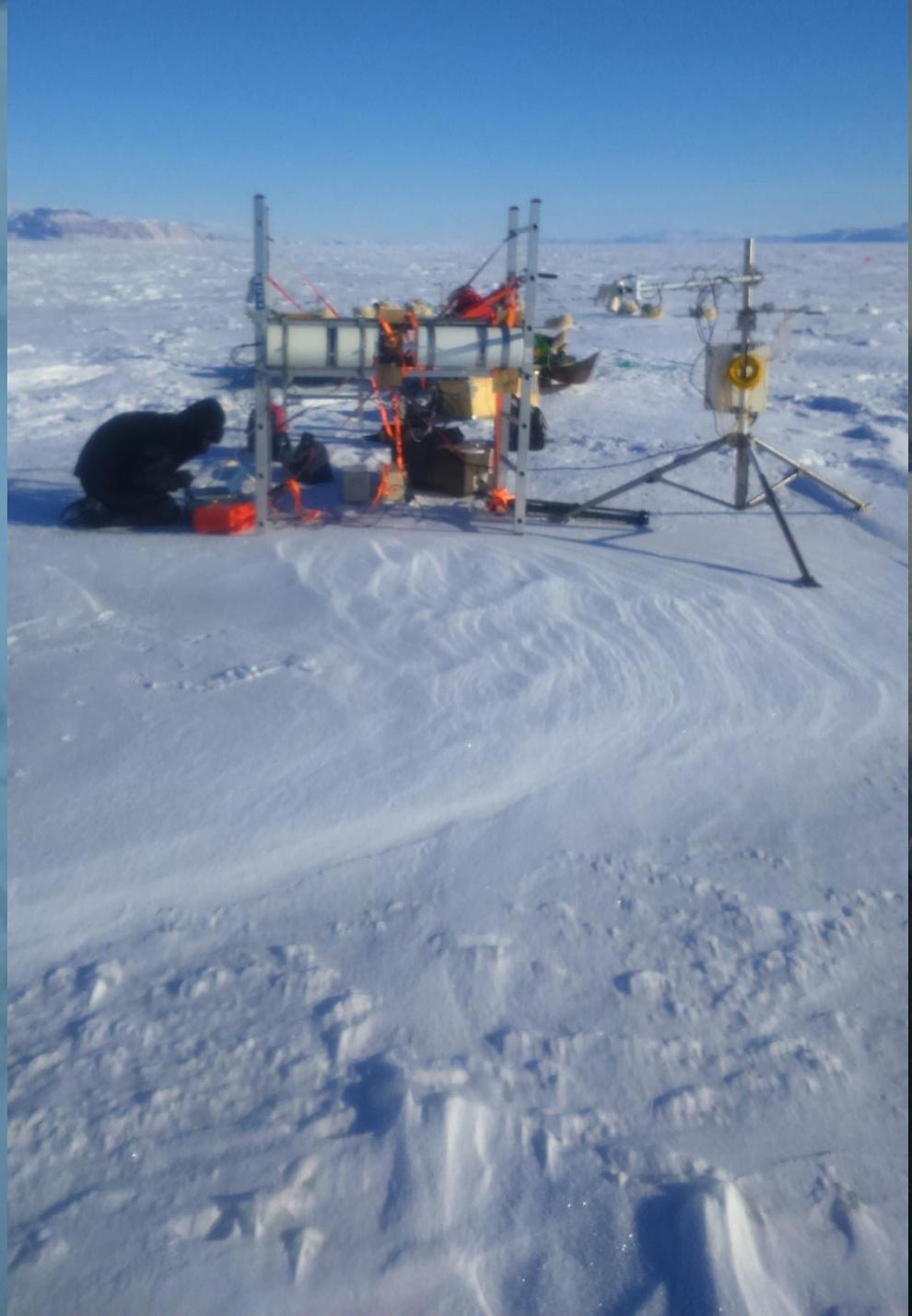
First-year ice  
Ridges/ deformed ice



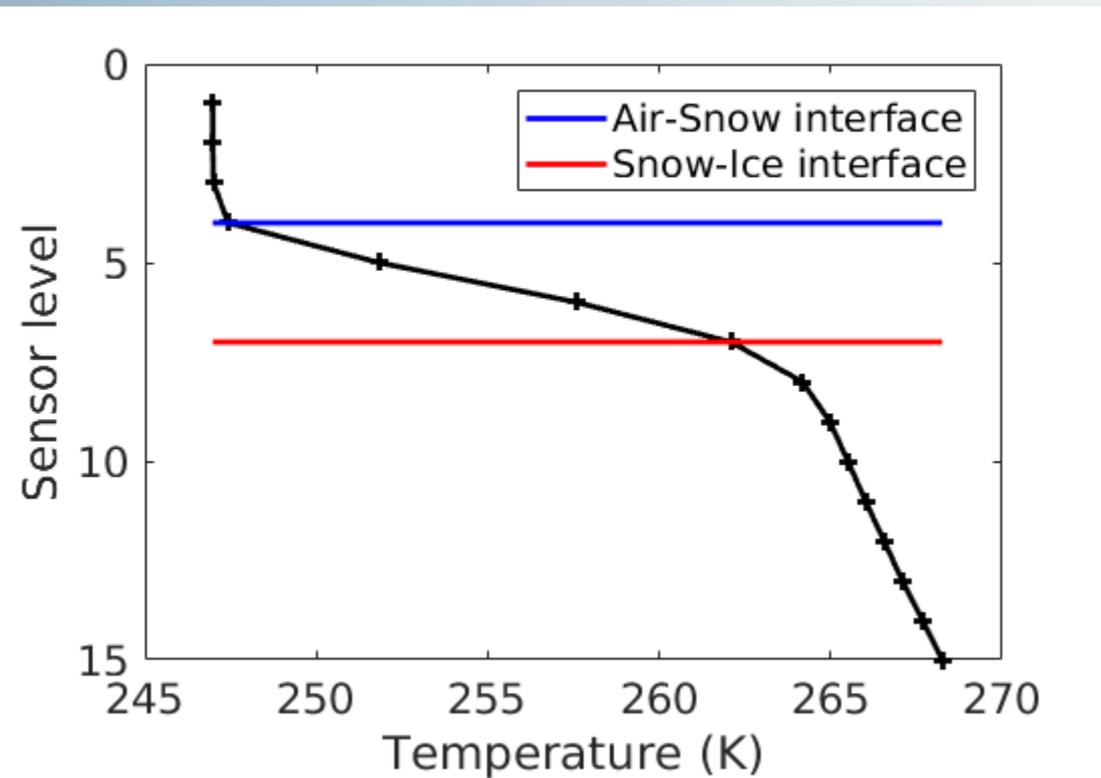
Multiyear ice  
Melt-ponds



Refrozen meltponds



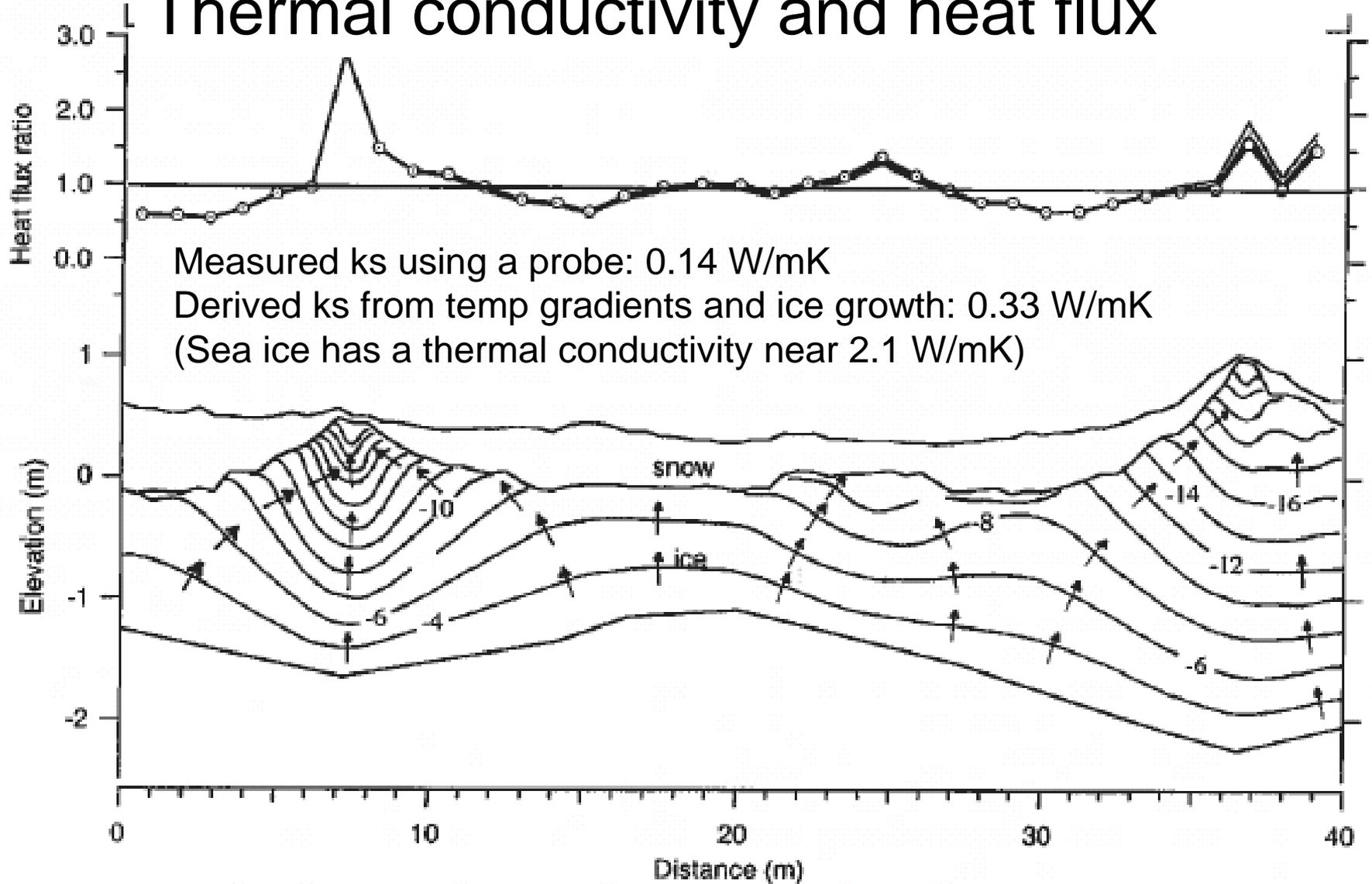
# The thermodynamics of sea ice



- **$k_i$** : thermal conductivity of ice f(salinity, temperature, porosity)
- **$k_s$** : thermal conductivity of snow f(density)
- **$T_w$** : water temperature (271.35K)
- **$T_{si}$** : snow ice tinterface temperature (measured by satellite?)
- **$T_a$** : snow surface temperature (measured by satellite)
- **$h_i$** : ice thickness
- **$h_s$** : snow depth

$$k_i \left( \frac{T_w - T_{si}}{h_i} \right) = k_s \left( \frac{T_{si} - T_a}{h_s} \right)$$

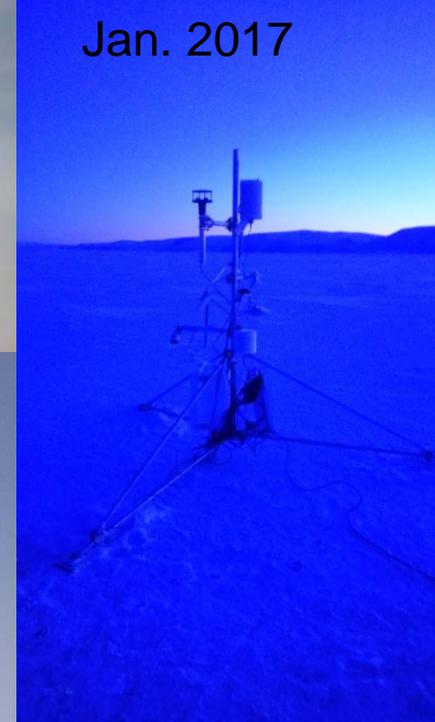
# Thermal conductivity and heat flux



Sturm, M. D. K. Perovich, J. Holmgren. 2002. Thermal conductivity and heat transfer through the snow on the ice of the Beaufort Sea. JGR 107(C21), 8043, doi:10.1029/2000JC000409

# Weather station in North West Greenland

Jan. 2017



Snow pinger

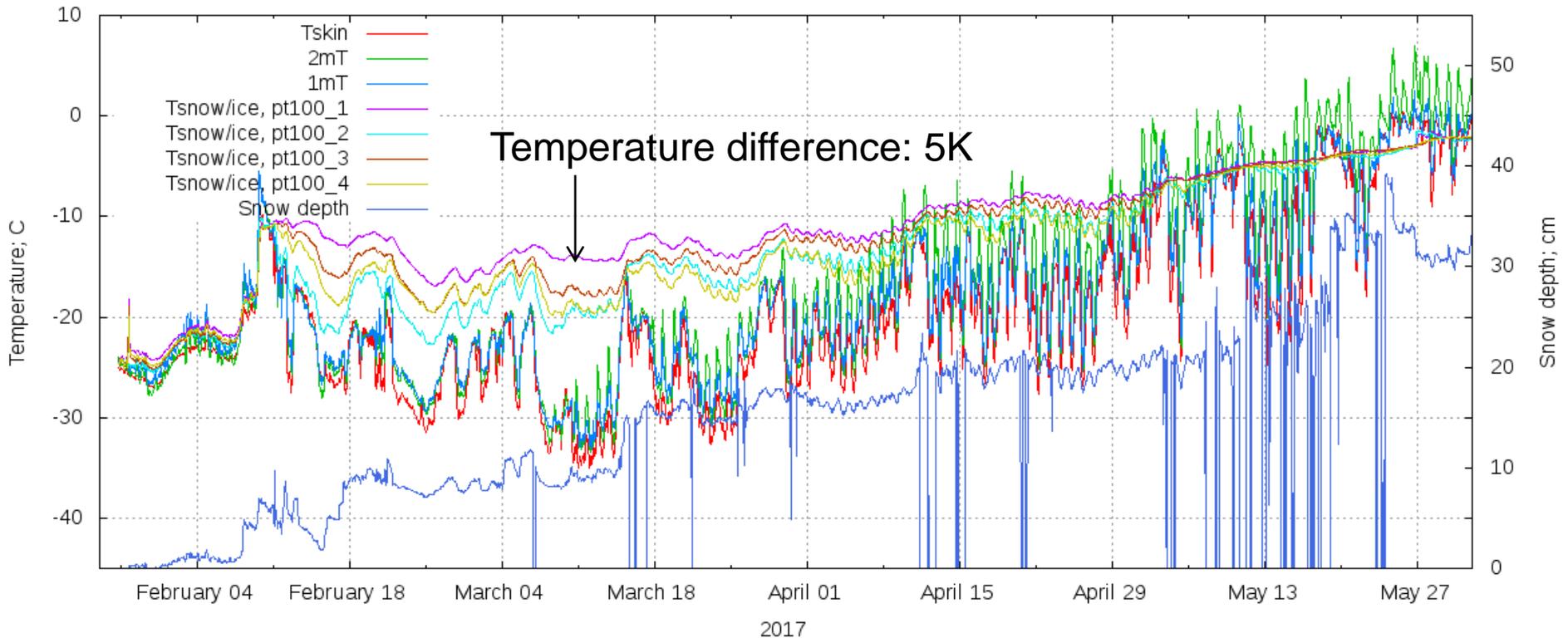


thermistors



June 2017

# Temperatures and snow depth from the weather station 2017



# How to measure snow?



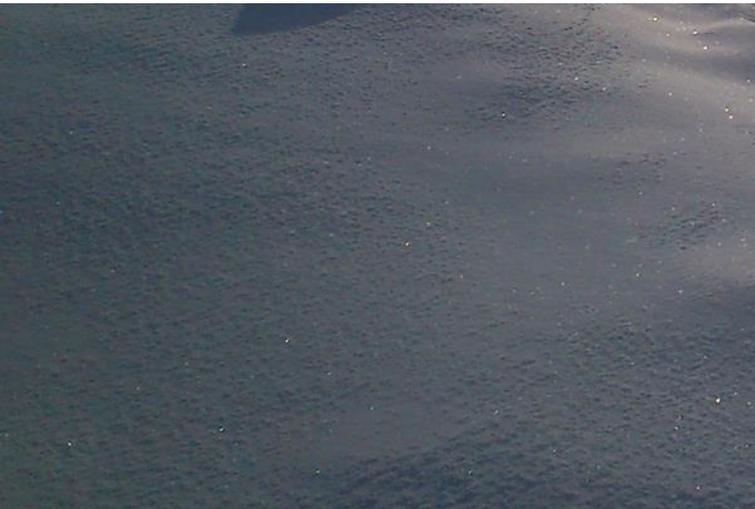
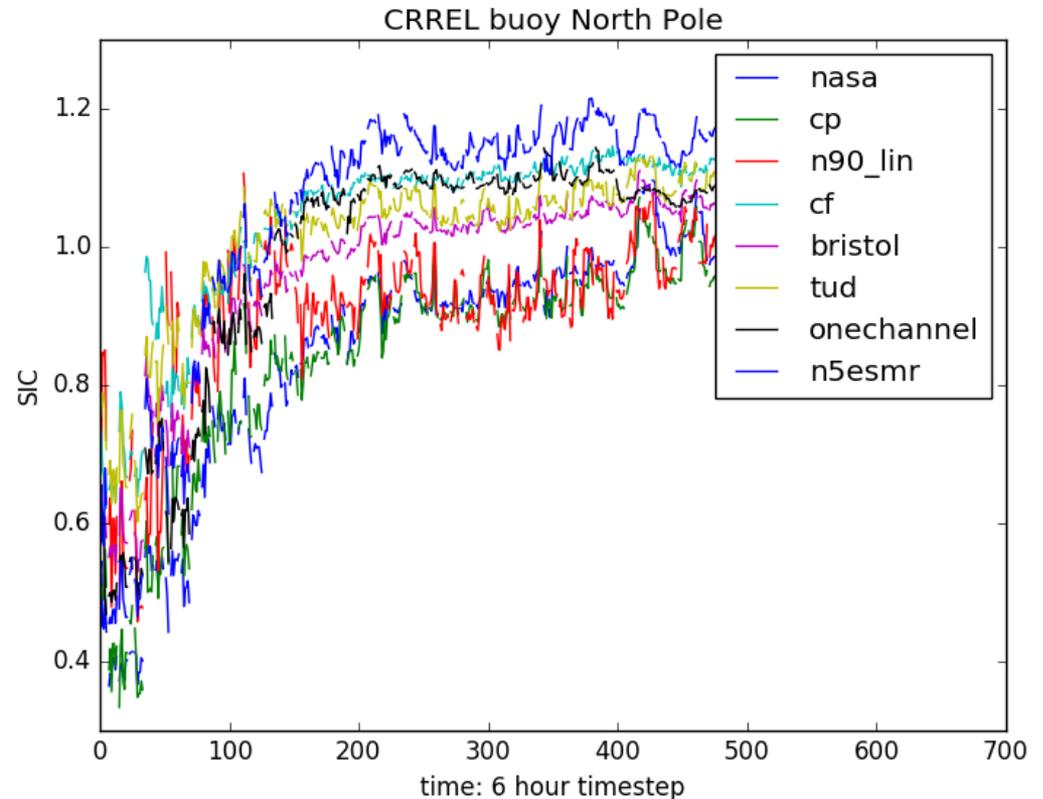
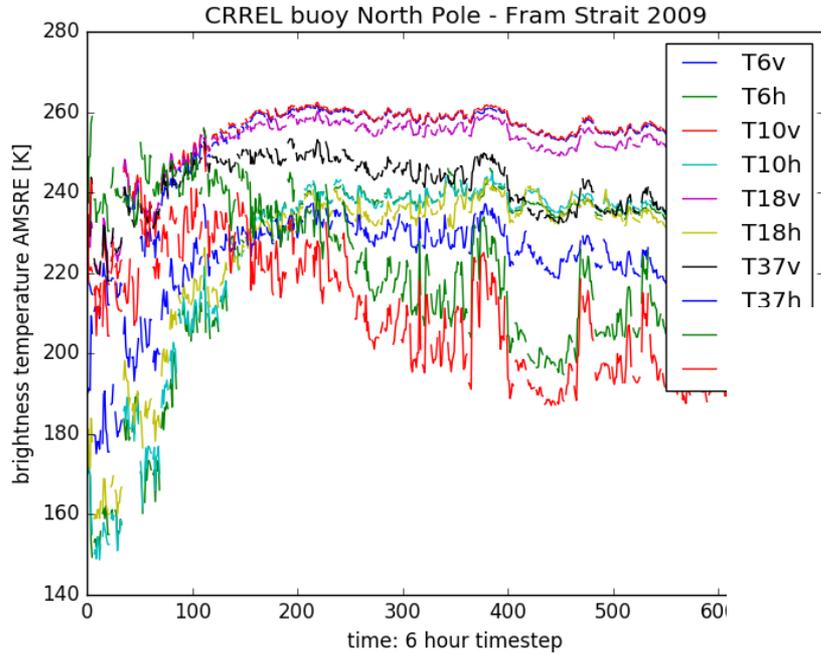
- Ice mass balance buoys: point measurements, interfaces derived from temperature profile.
- Snow radar: NASA Operation Ice Bridge (OIB), long transects covering different ice types, data in the Arctic are primarily from March and April
- Campaigns: measuring along snow lines, can include detailed description of density, grain size, temperature and salinity.

One of the most comprehensive and also publicly available datasets where microwave data is combined with snow data is the ESA CCI Round Robin Data Package: <http://www.seaice.dk/sicci2/>



# The round robin data package for snow and sea ice (concentration)

[www.seaice.dk/sicci2/](http://www.seaice.dk/sicci2/)

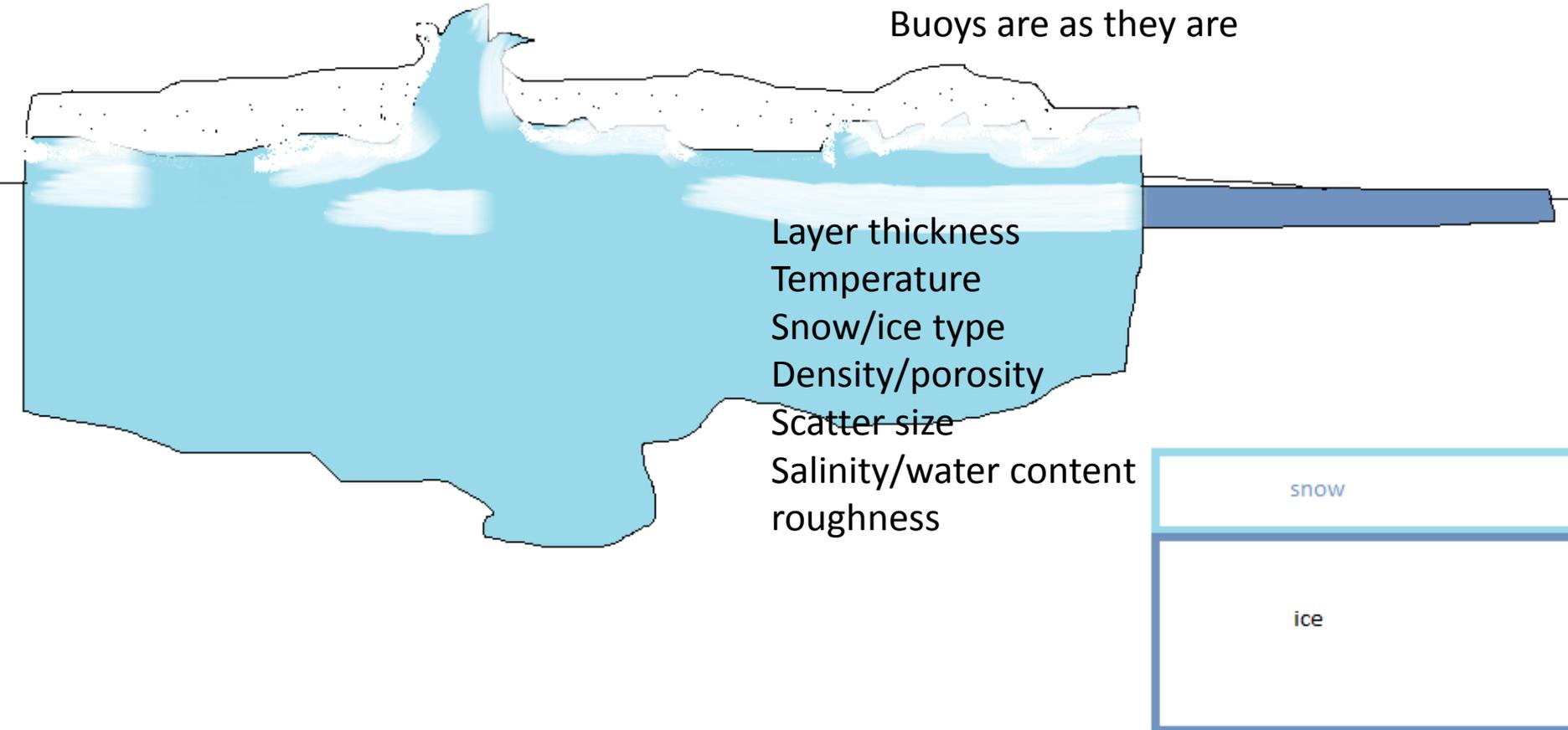


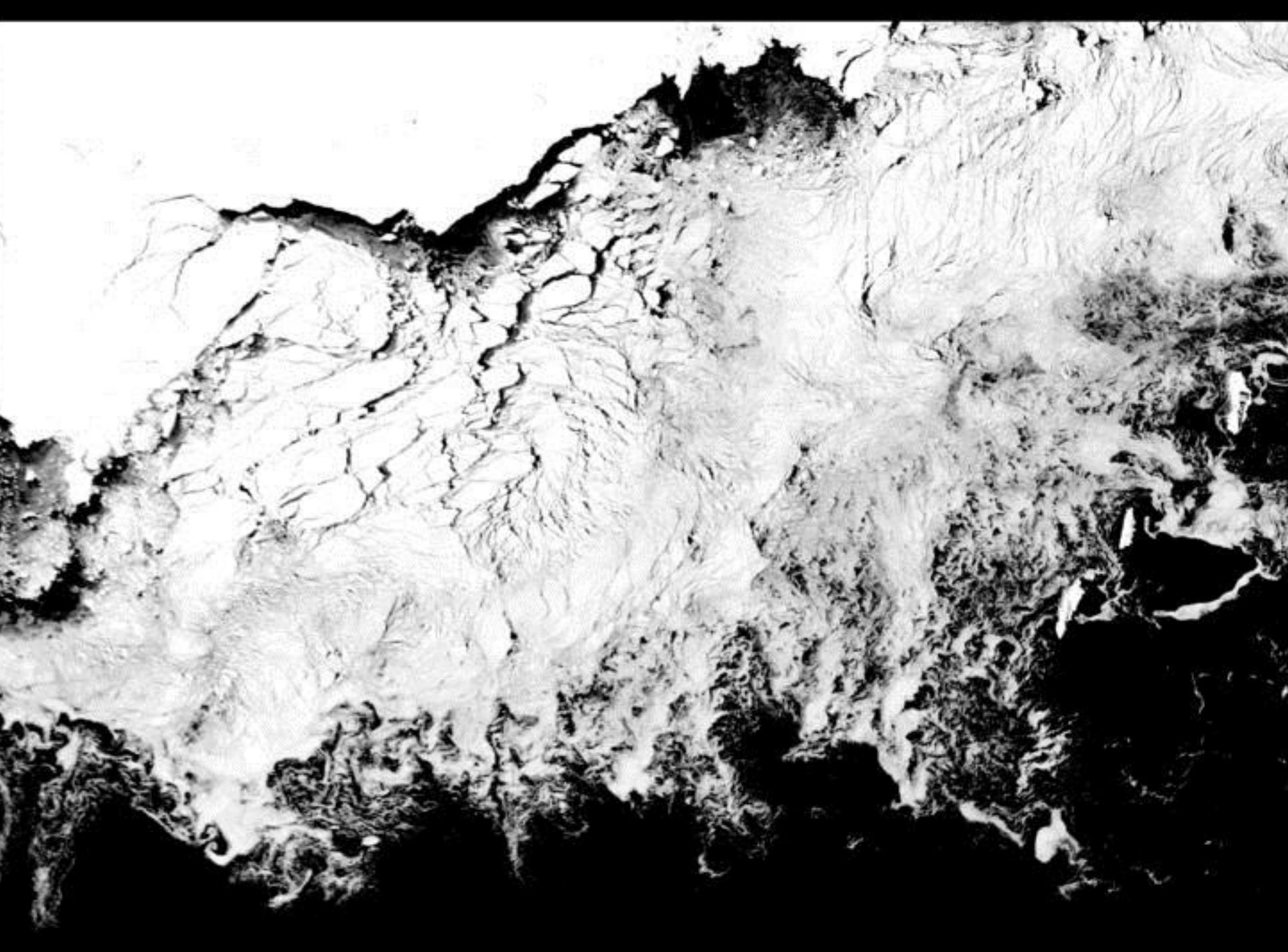
# Point vs spatially averaged footprints

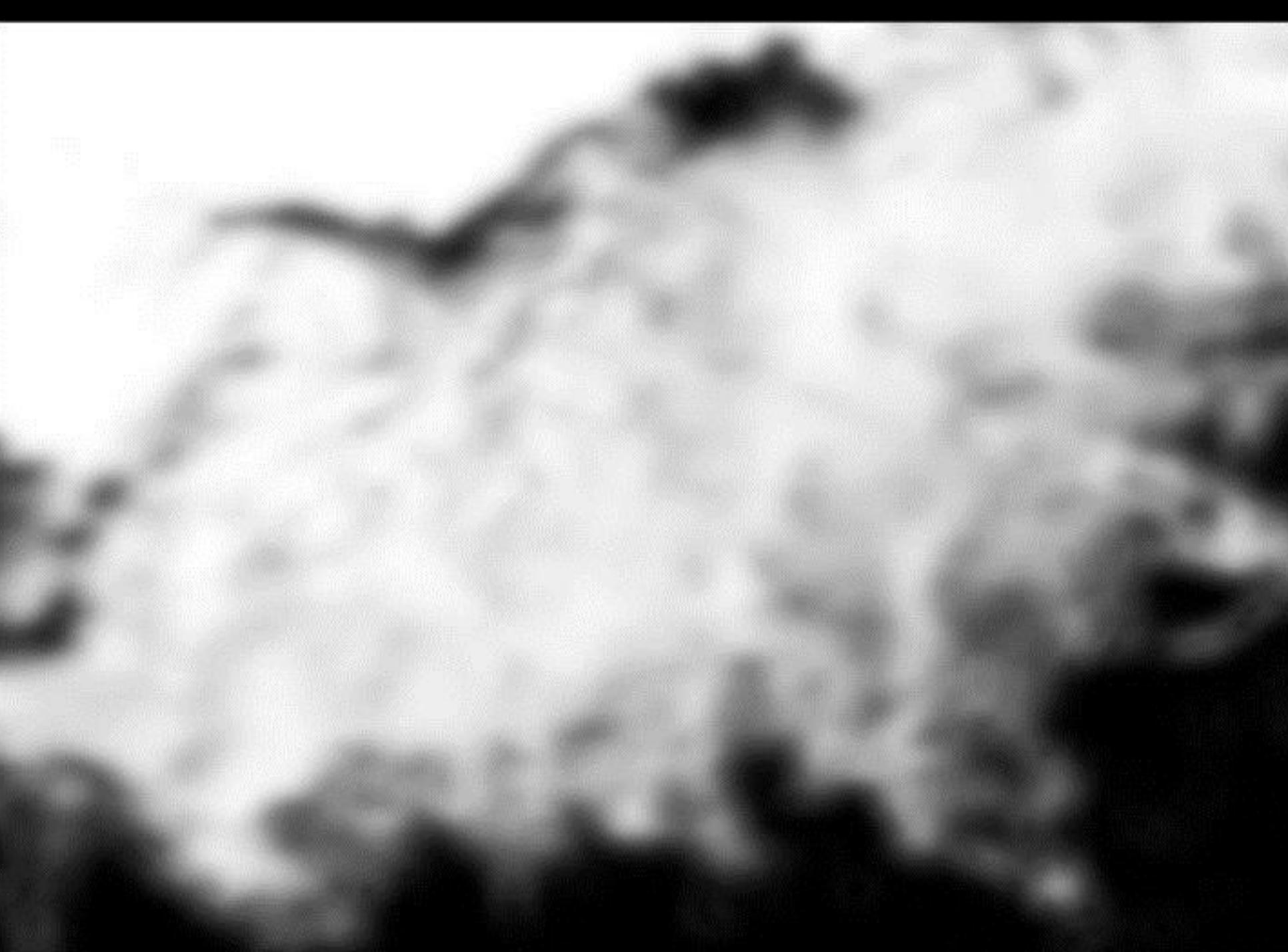
Sea ice concentration  
Spatial distribution of snow

## RRDP

Satellite data are on a 50 km grid  
OIB are averaged on 50km sections  
Buoys are as they are



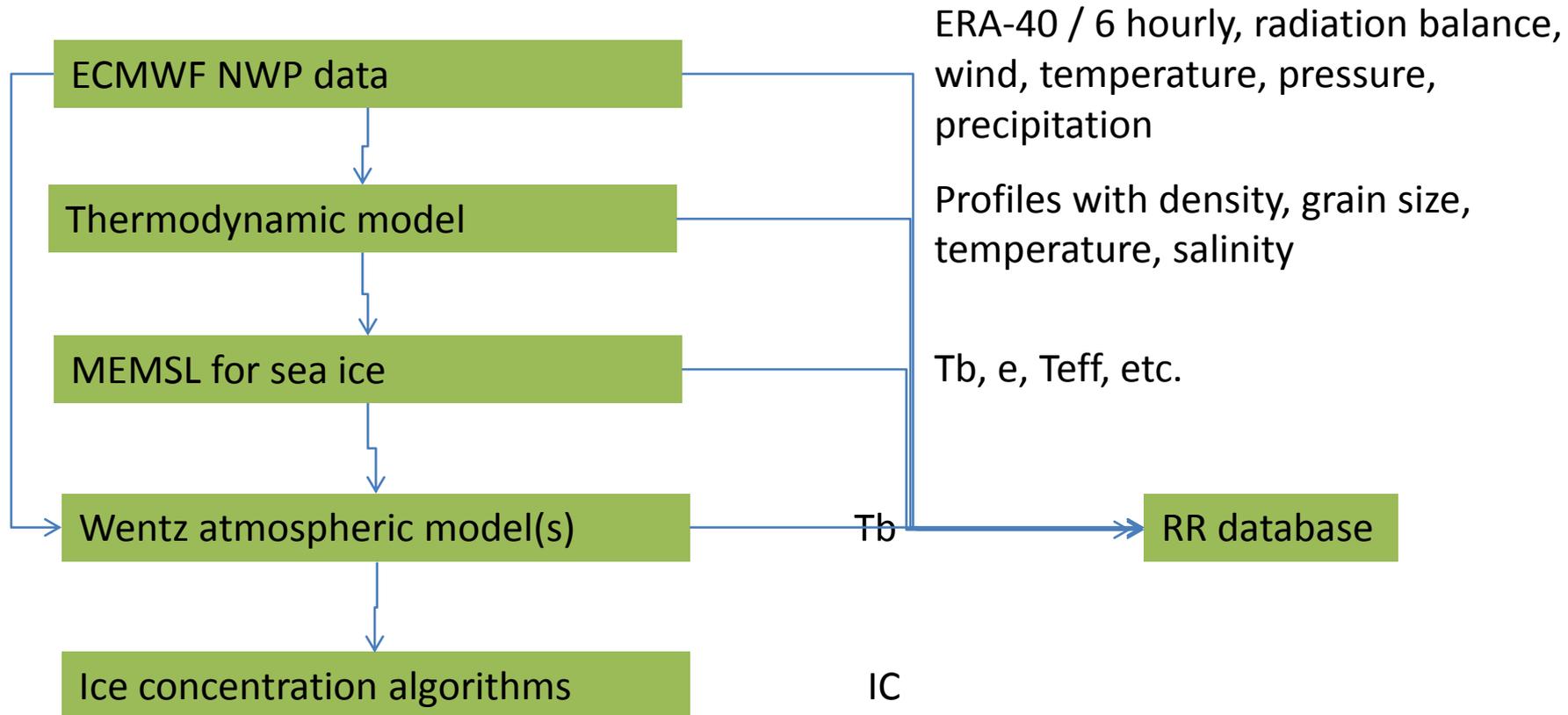




# Microwave radiometry



# The simulated data



The system is described in:

Tonboe, R. T. The simulated sea ice thermal microwave emission at window and sounding frequencies. *Tellus 62A*, 333-344, 2010.

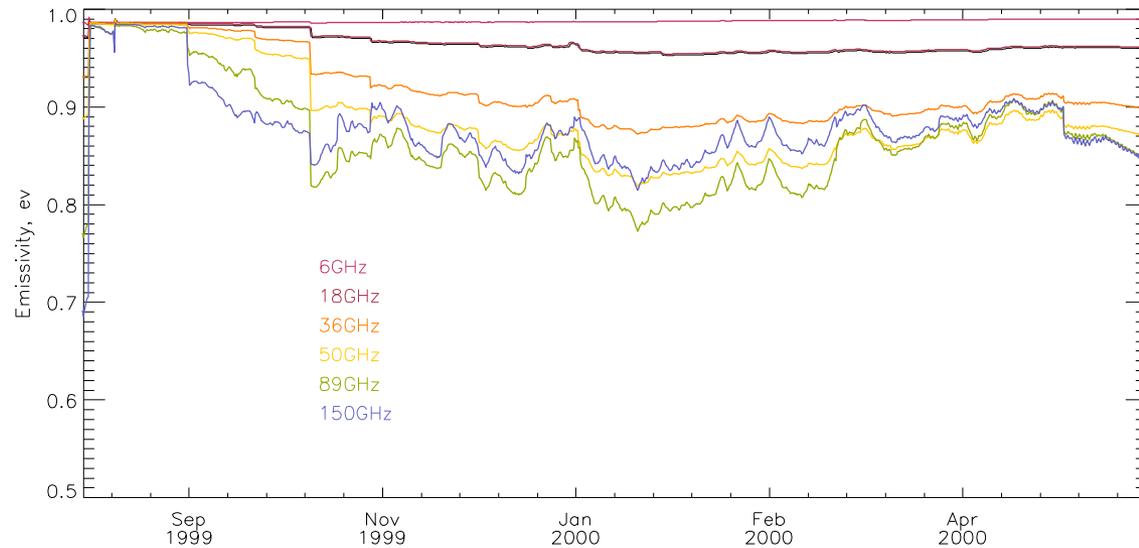
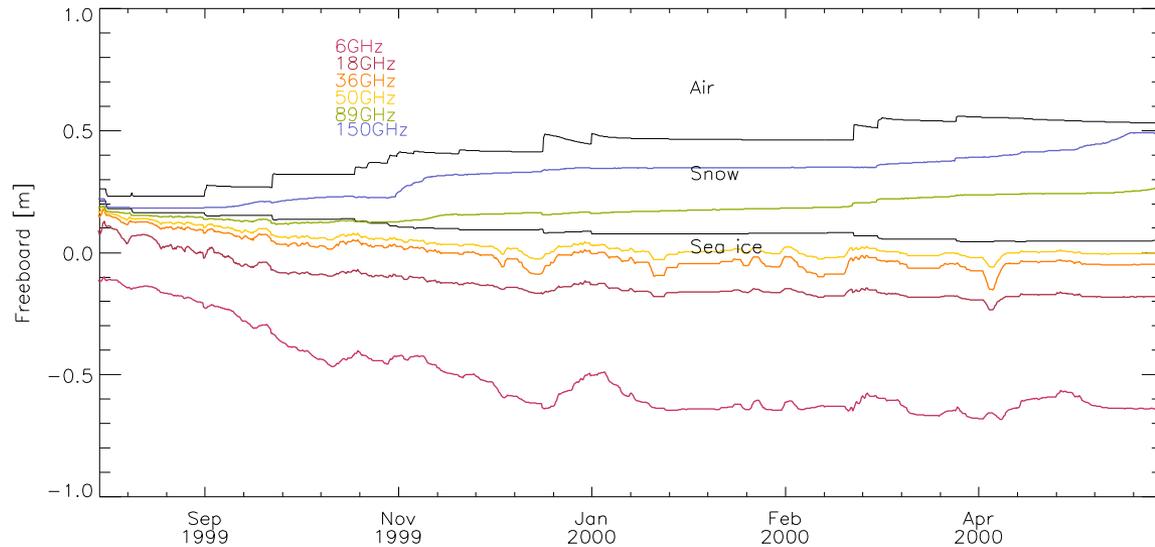
Tonboe, R. T., G. Dybkjær, J. L. Høyer. Simulations of the snow covered sea ice surface temperature and microwave effective temperature. *Tellus 63A*, 1028-1037, 2011.



## Energy and mass balance

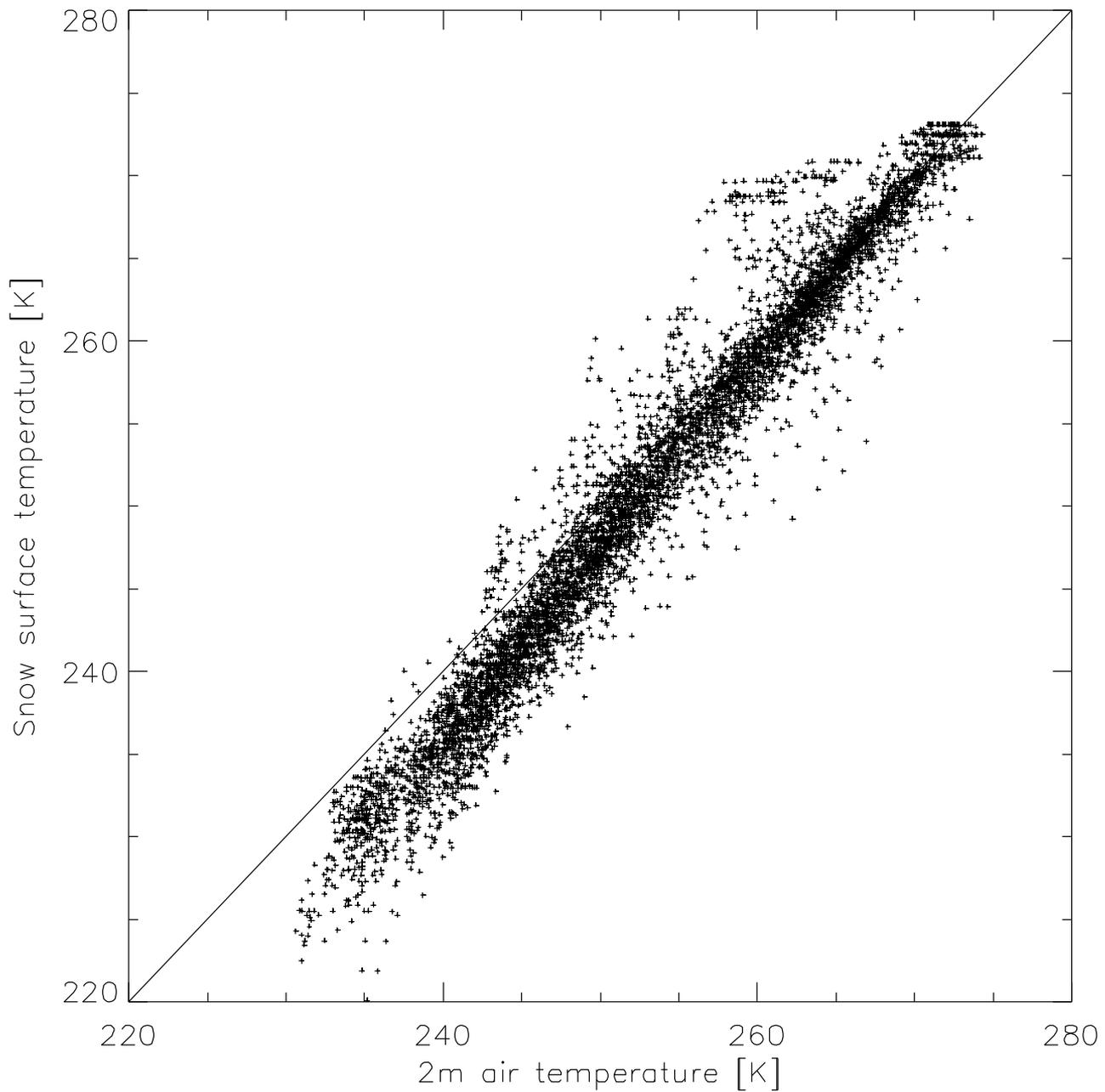
- Air temperature
- Wind speed
- Radiation balance
- Snow accumulation
- Ice growth

# The microwave emissivity (6-150GHz)





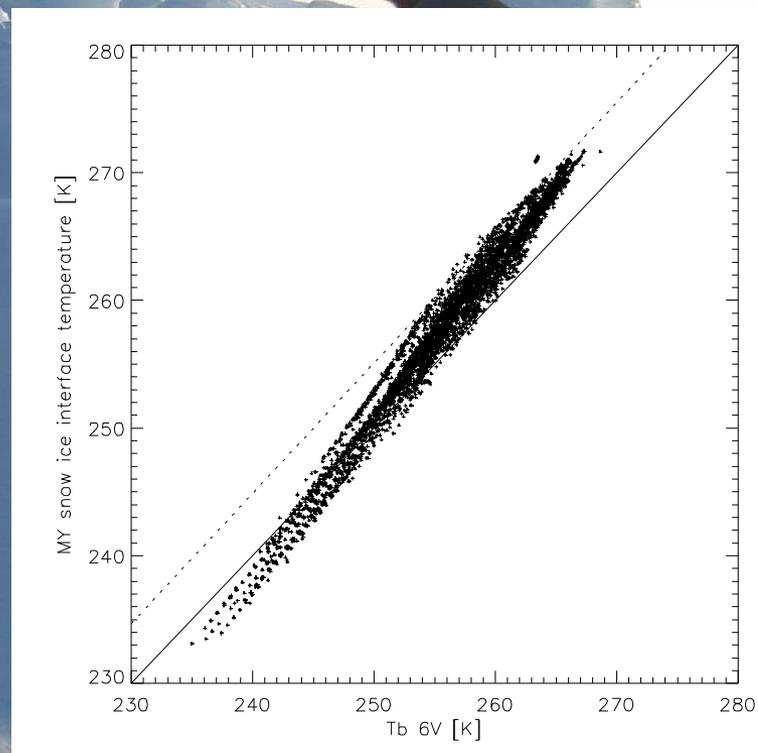
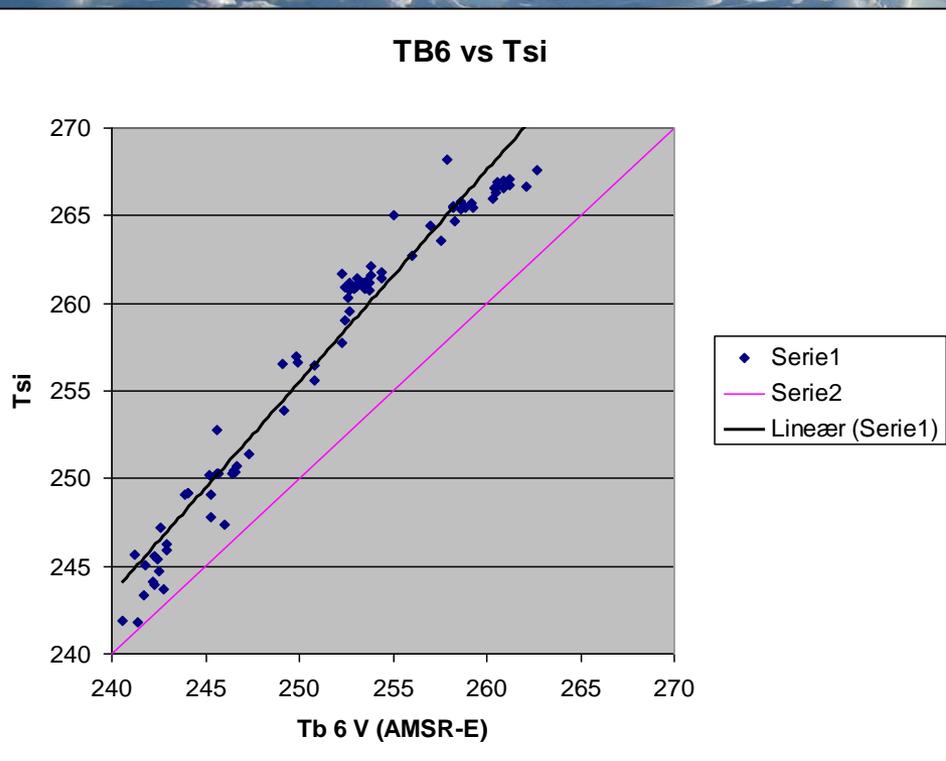
# Multiyear ice air temperature surface temperature



# Penetration depth is a function of (ice) temperature and *vice versa*

Measurements collocated with  
satellite observations (from Phil  
Hwang)

6 simulated multiyear ice  
profiles



Physical temperature vs. effective temperature?

# Snow depth algorithms using microwave radiometer data

## Proxy for scattering, $h_s = a_1 + a_2 \text{GRV}(\text{ice})$

Markus, T. and D. J. Cavalieri. 1998. Snow depth distribution over sea ice in the Southern Ocean from satellite passive microwave data. In *Antarctic Sea Ice: Physical Processes, Interactions and Variability*, 19-39. Washington, DC: American Geophysical Union.

## Thermodynamical effect of the snow cover

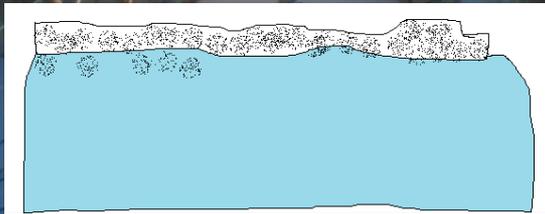
Maass et al. Snow thickness retrieval over thick Arctic sea ice using SMOS satellite data. *The Cryosphere* 7, 1971-1989, 2013.

Retrieval of snow depth is either retrieved directly ( $SD=f(\text{GR}...)$ ) or by inversion ( $T_b=f(SD...)$ ). Sometimes a translation between the detailed input to the model and the bulk parameters to retrieve.

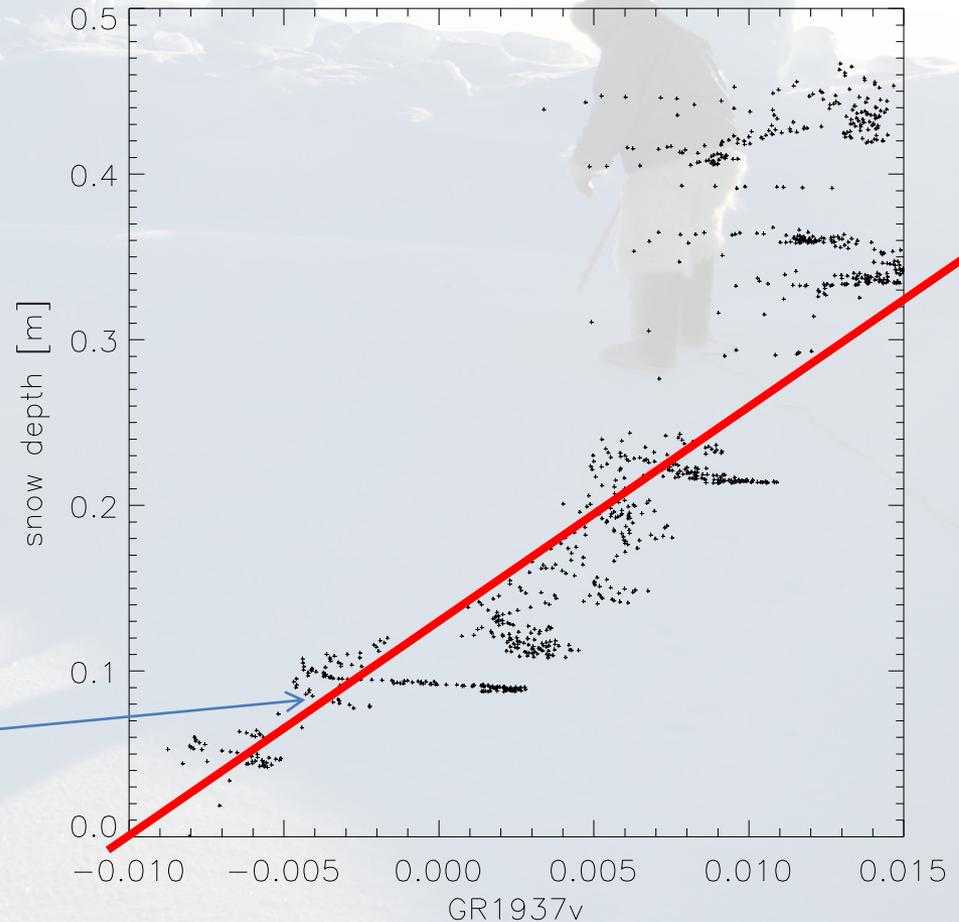
# The snow thickness algorithm -an empirical regression model



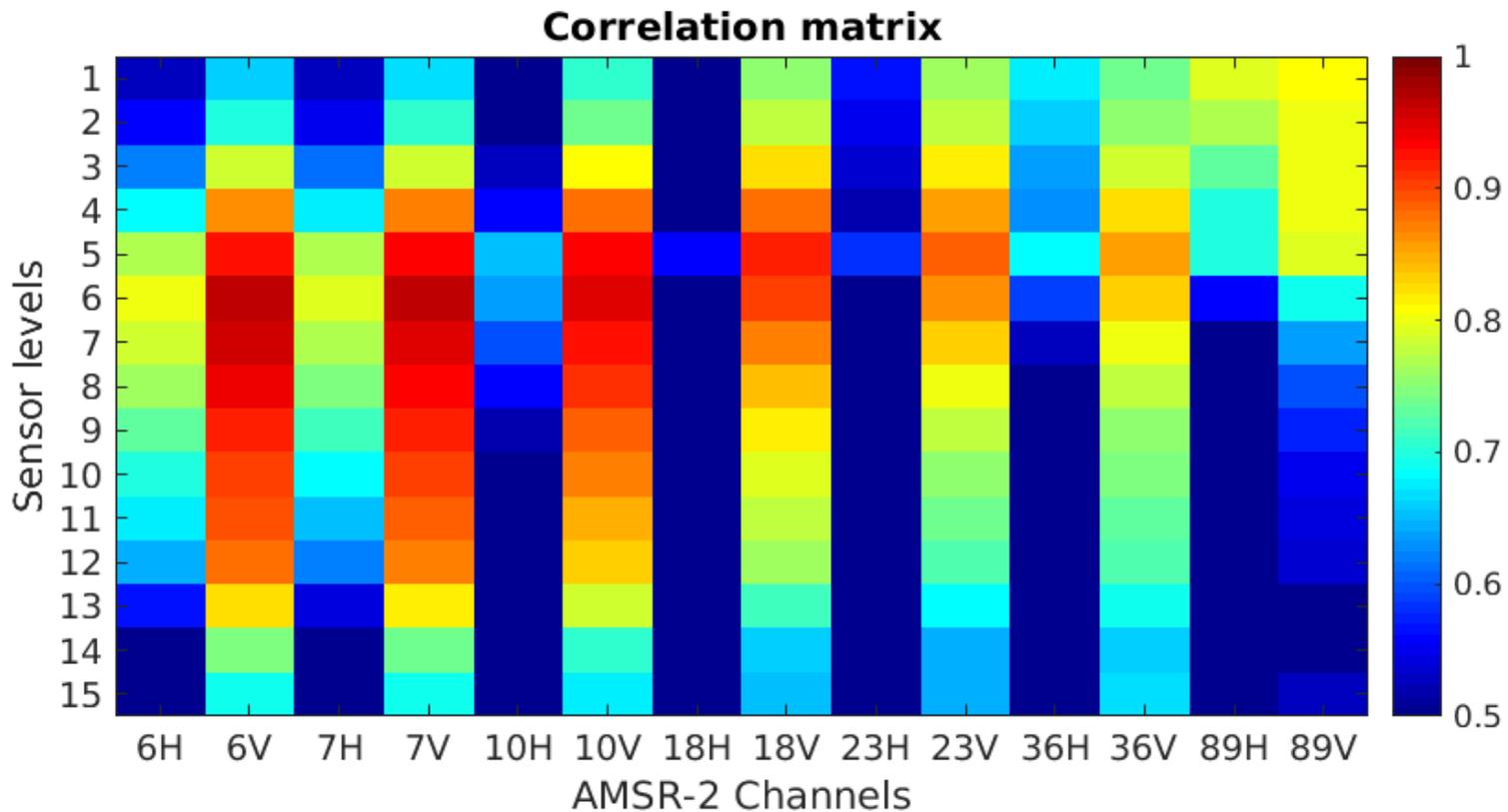
Depth hoar above ice surface  
*Lincoln Sea 2006*



The line slope and offset  
are based on a particular  
dataset

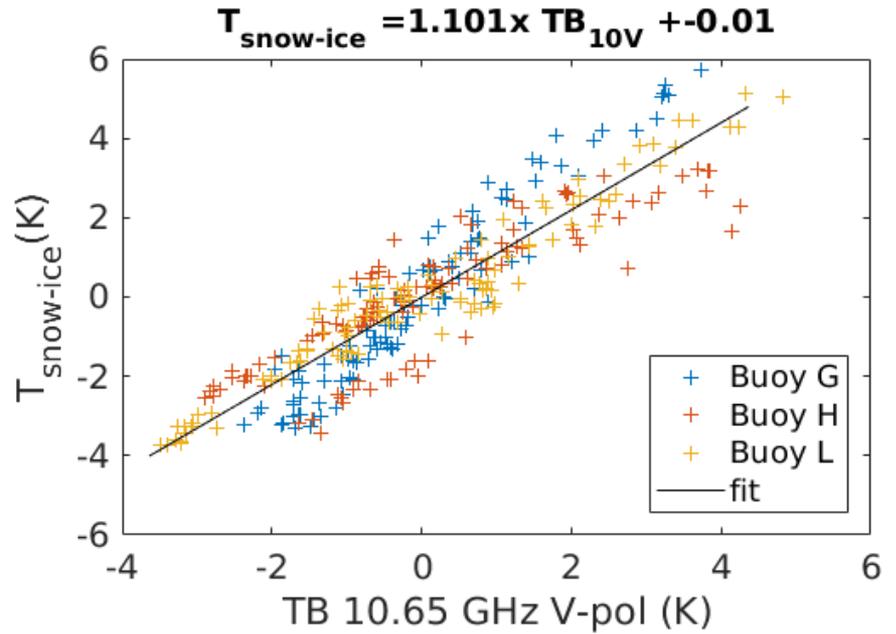
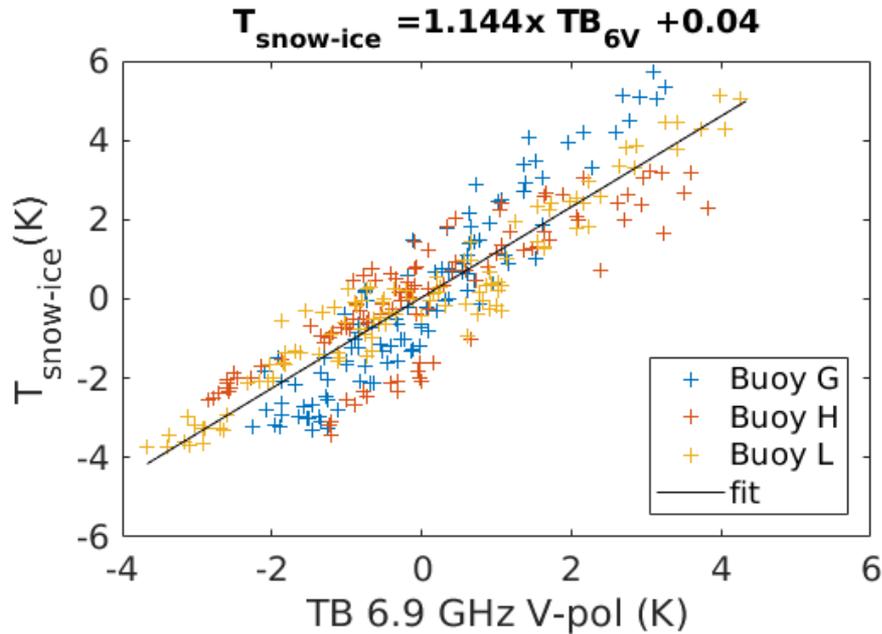


# CRREL buoy 2012J



Top story on <http://www.osi-saf.org/>

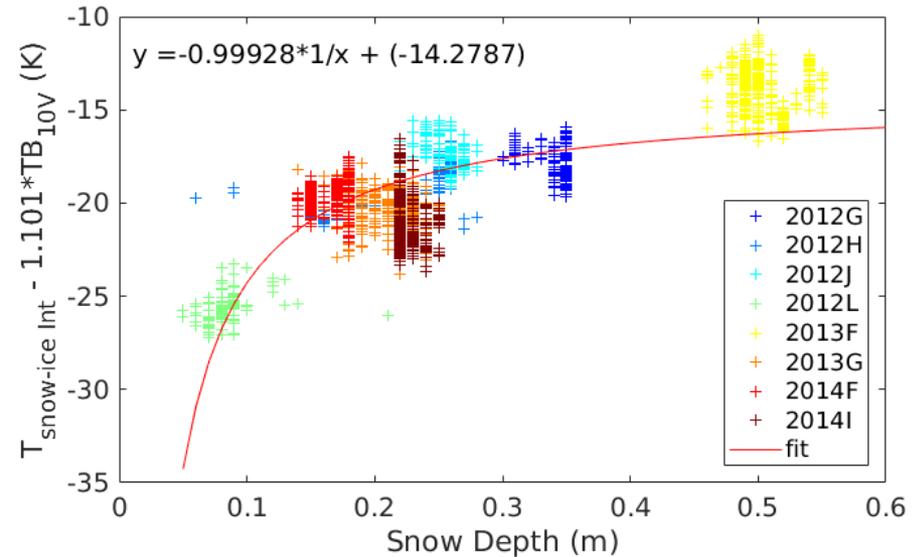
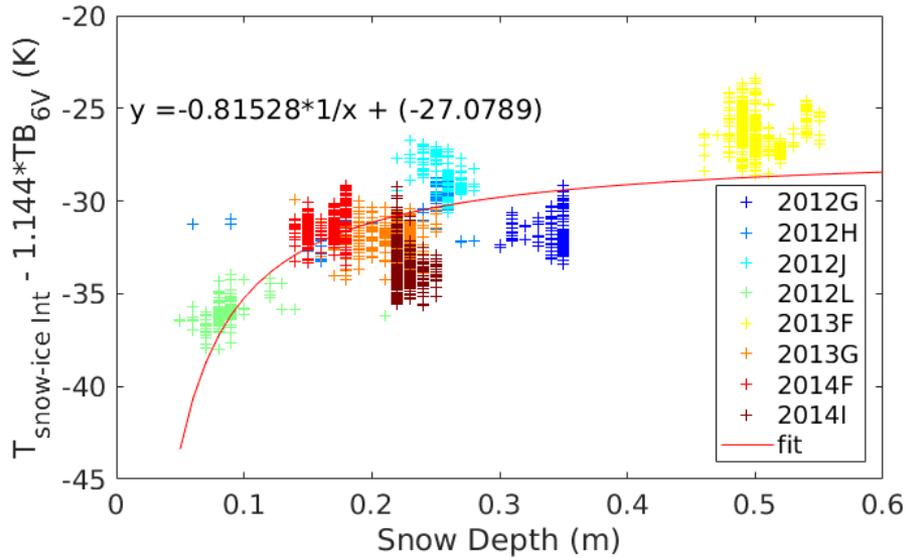
# The snow ice interface temperature



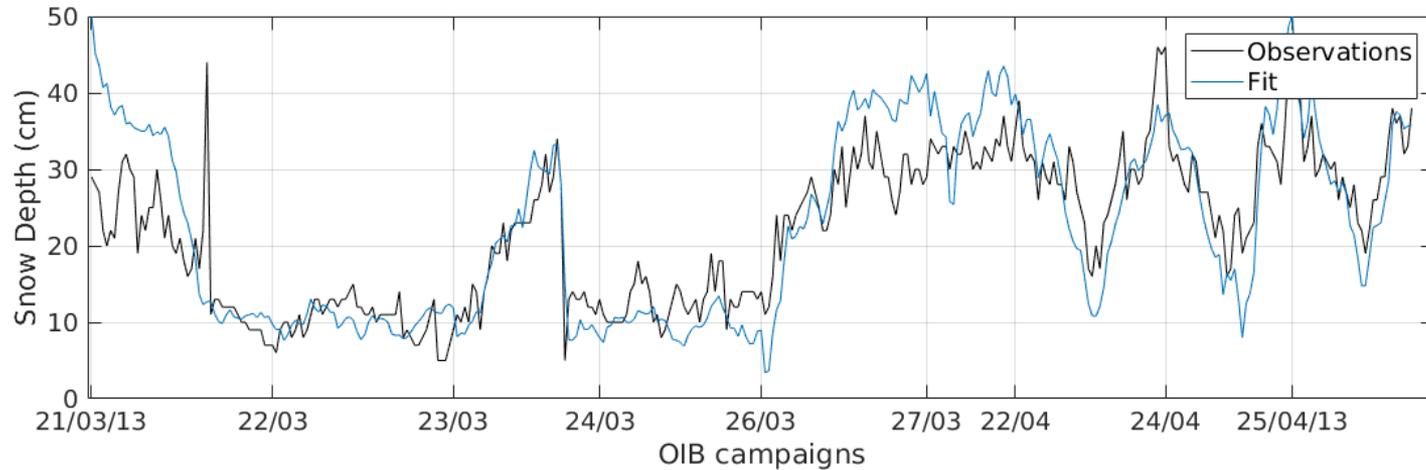
$$\frac{dT}{dz} = \frac{T_a - T_{si}}{hs}$$

$$T_{si} = T_a/2 + C/hs$$

# The snow depth dependence



$$hs = 1.7701 + 0.017462 \times T6v - 0.02801 \times T18v + 0.0040926 \times T37v$$



# Different ways to retrieve snow depth

## Simple regression models

Snow\_depth = f(GR1836v), e.g. Markus and Cavalieri, 1998

Snow\_depth = f(GR0610v, GR0618v, .....) using linear regression and the RRDP data.

Snow depth = f(Tb6v, Tb18v...)

## Brightness temperature models based on regression combined with OE

```
function [T6vsim,T6hsim,T10vsim,T10hsim,T18vsim,T18hsim,T36vsim,T36hsim,T89vsim,T89hsim]=reg_mod13(IST,SD,IT)
```

```
%the mar+april 2013 model
```

```
T6vsim= 151.981535394 + 0.39827296166 .* IST+ 23.3600203008 .* SD -3.03183834111 .* IT;
```

```
T6hsim= 55.2623240539 + 0.687577210357 .* IST+ 12.9621301692 .* SD -1.66486943272 .* IT;
```

```
T10vsim= 145.878105173 + 0.435432823207 .* IST+ 0.743658800361 .* SD -4.20200228328 .* IT;
```

```
T10hsim= 45
```

```
T18vsim= 13
```

## Brightness temperature model based on physics combined with OE

High resolution 1-D emission model - a sea ice version of MEMLS in combination with optimal estimation.

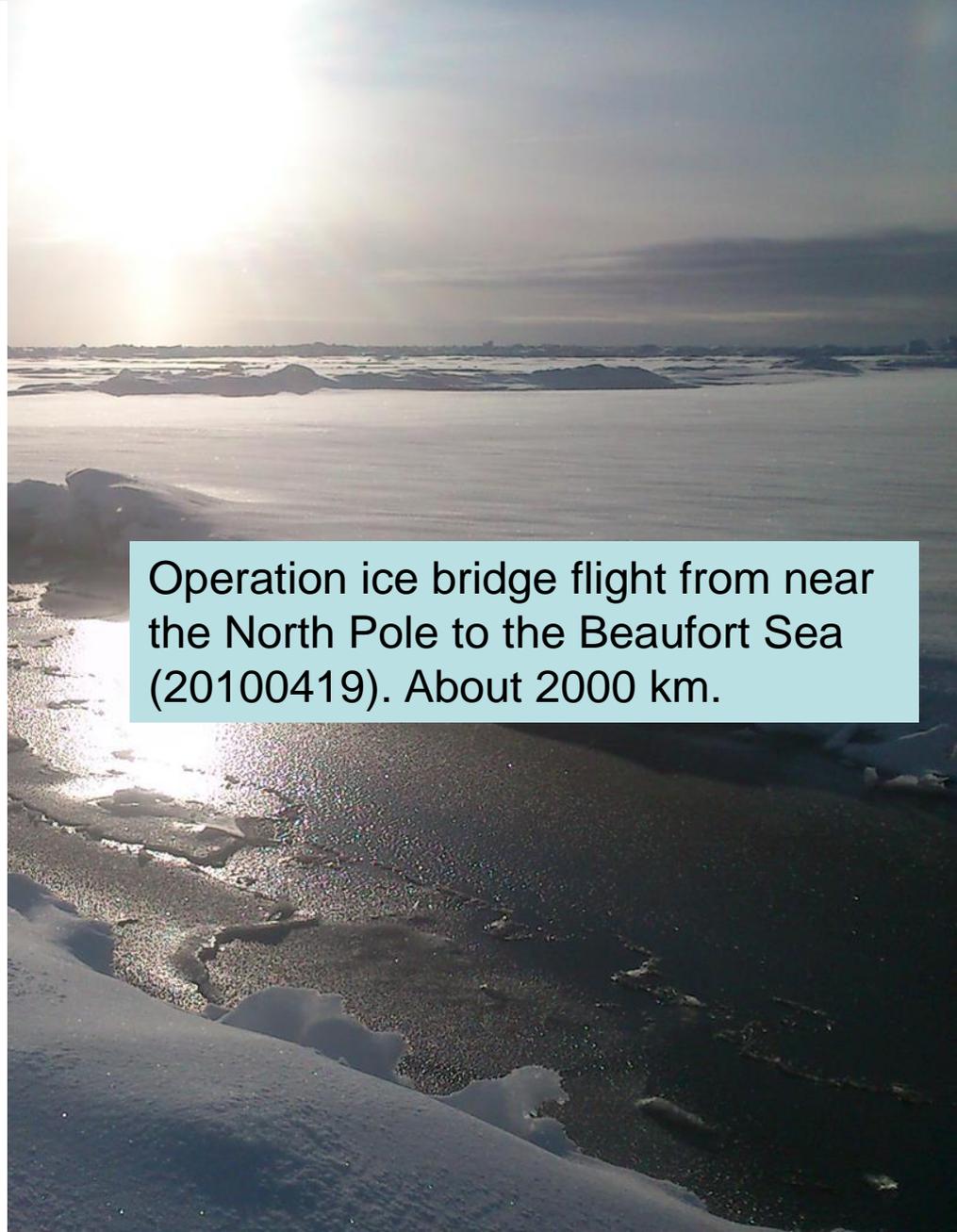
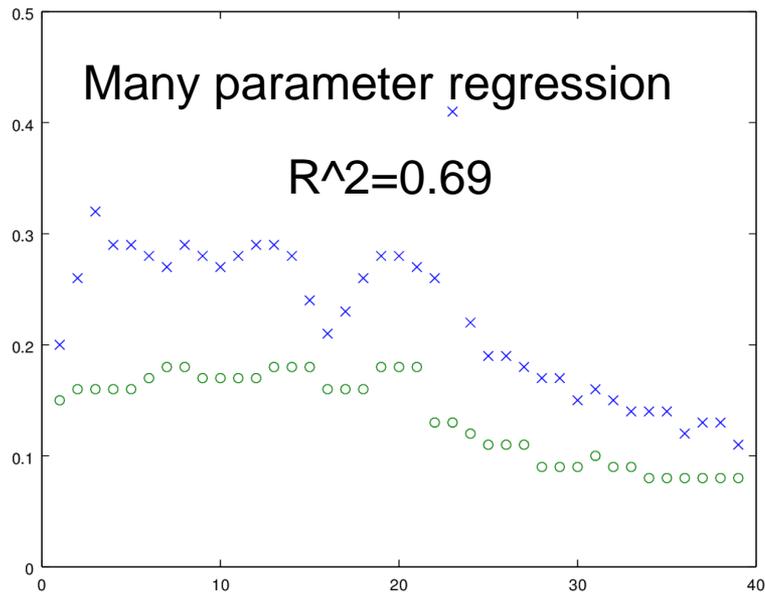
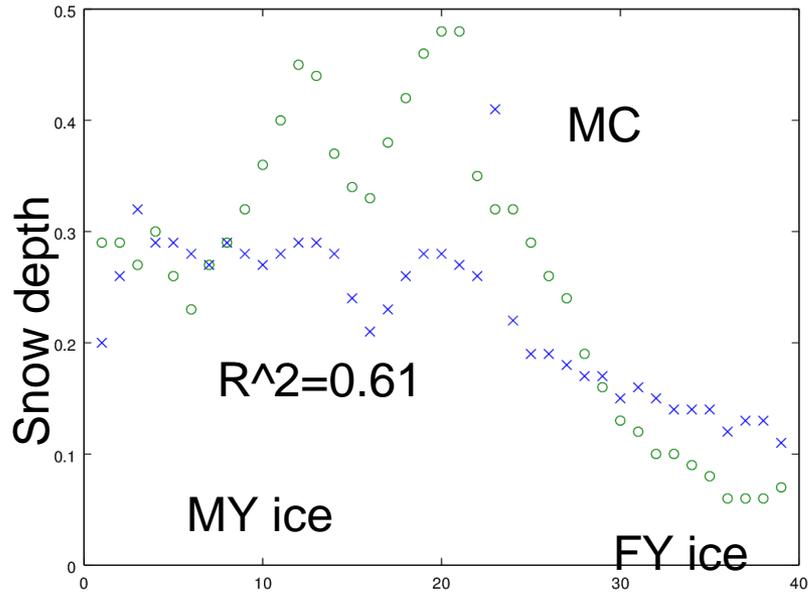
```
T36hsim= 131.862853412 + 0.429489868157 .* IST -214.352191714 .* SD -3.03524993537 .* IT;
```

```
T89vsim= 2.52567864217 + 0.902202528995 .* IST -180.427137566 .* SD+ 1.90480465092 .* IT;
```

```
T89hsim= 31.1206976877 + 0.743826485118 .* IST -184.806381816 .* SD+ 3.19723383624 .* IT;
```

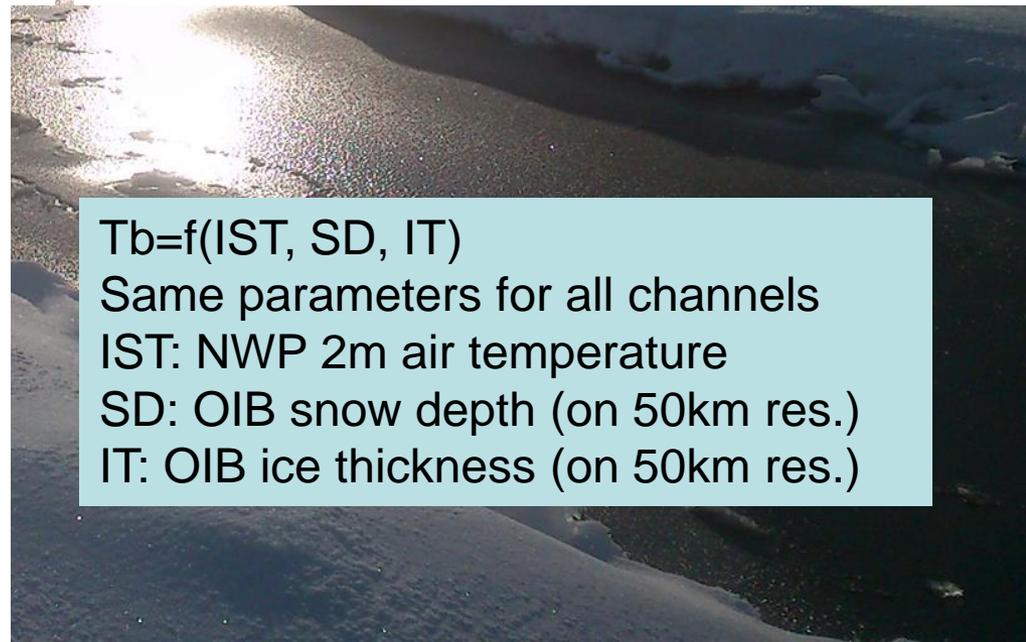
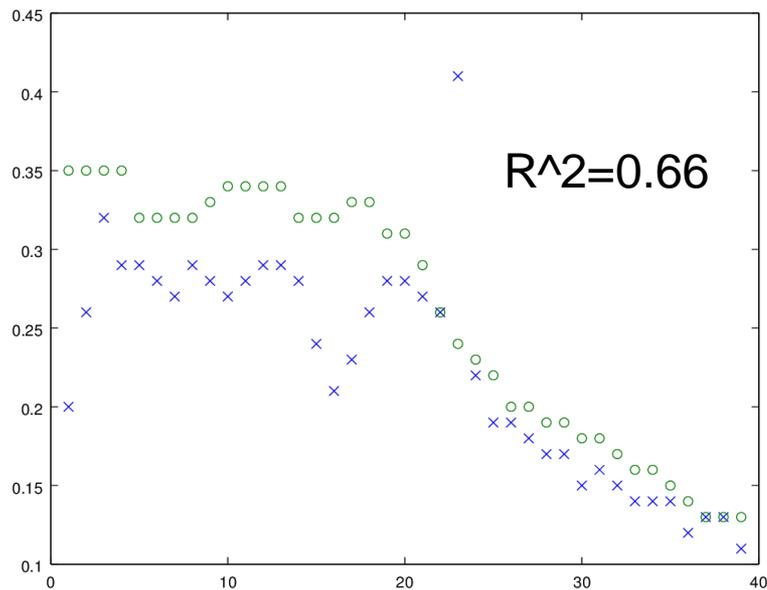
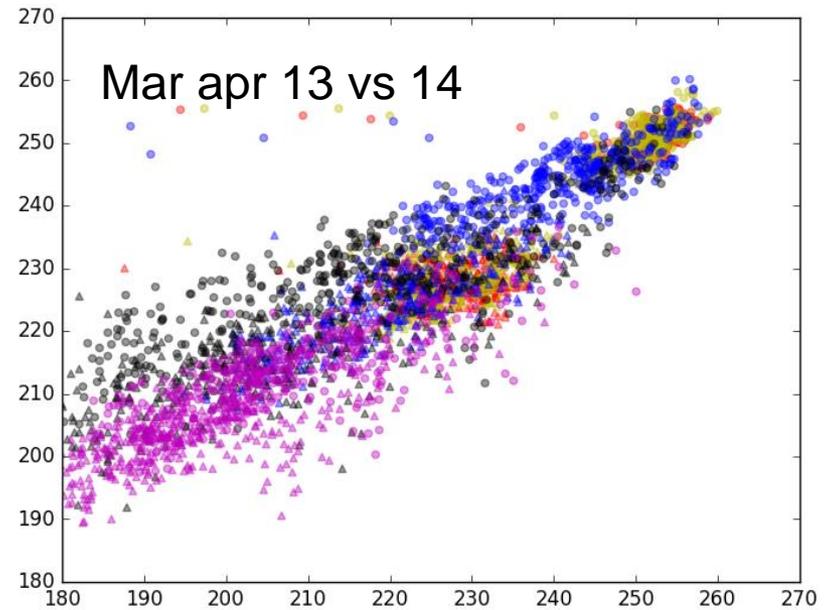
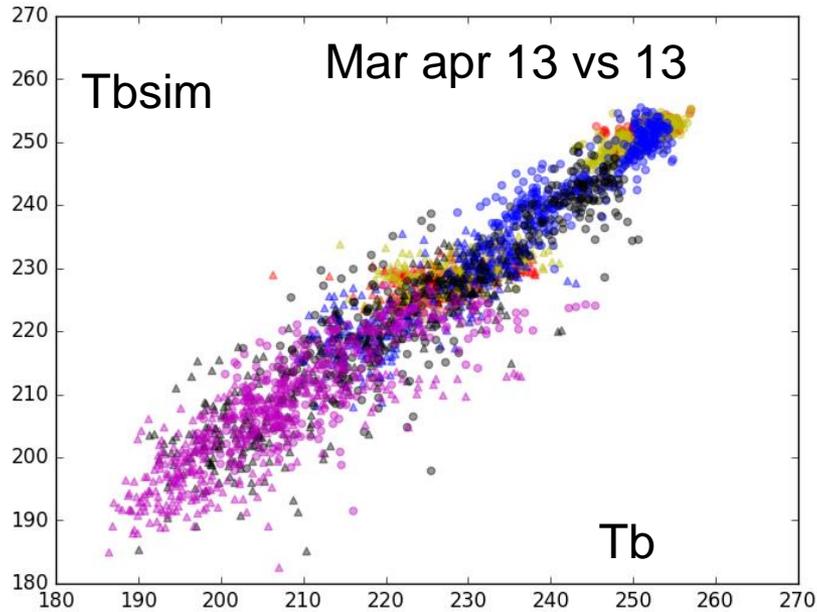
Optimal estimation aka statistical estimation theory needs first guess and covariance matrices

# Simple regression model case



Operation ice bridge flight from near the North Pole to the Beaufort Sea (20100419). About 2000 km.

# Brightness temperature model based on regression + OE



$$T_b = f(IST, SD, IT)$$

Same parameters for all channels

IST: NWP 2m air temperature

SD: OIB snow depth (on 50km res.)

IT: OIB ice thickness (on 50km res.)

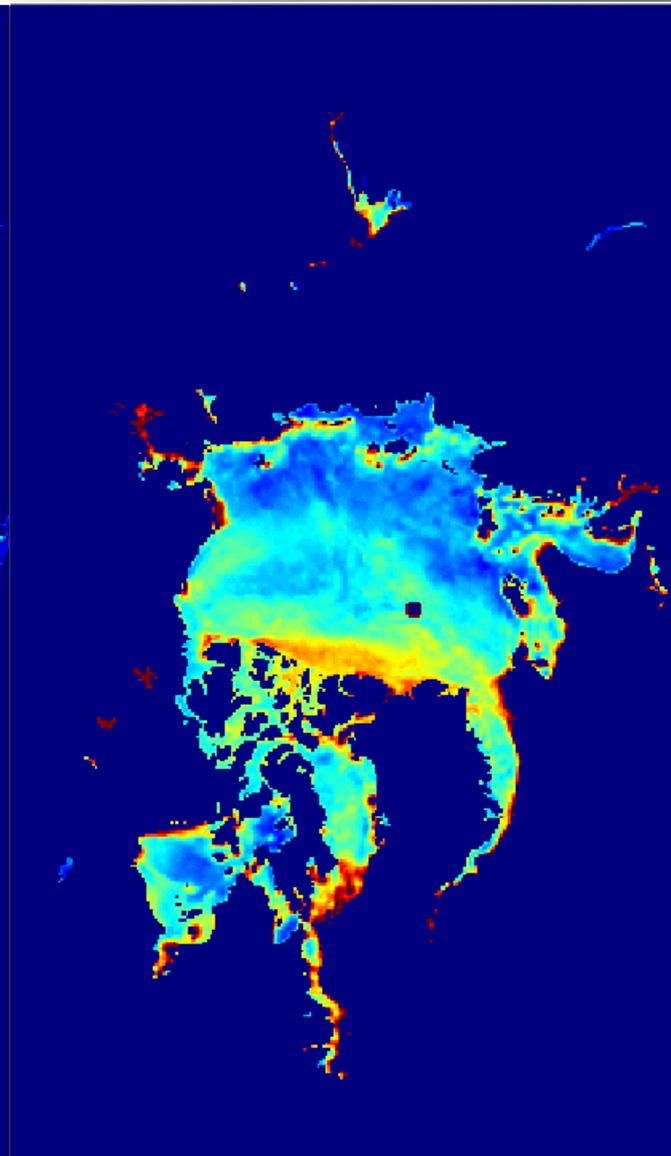
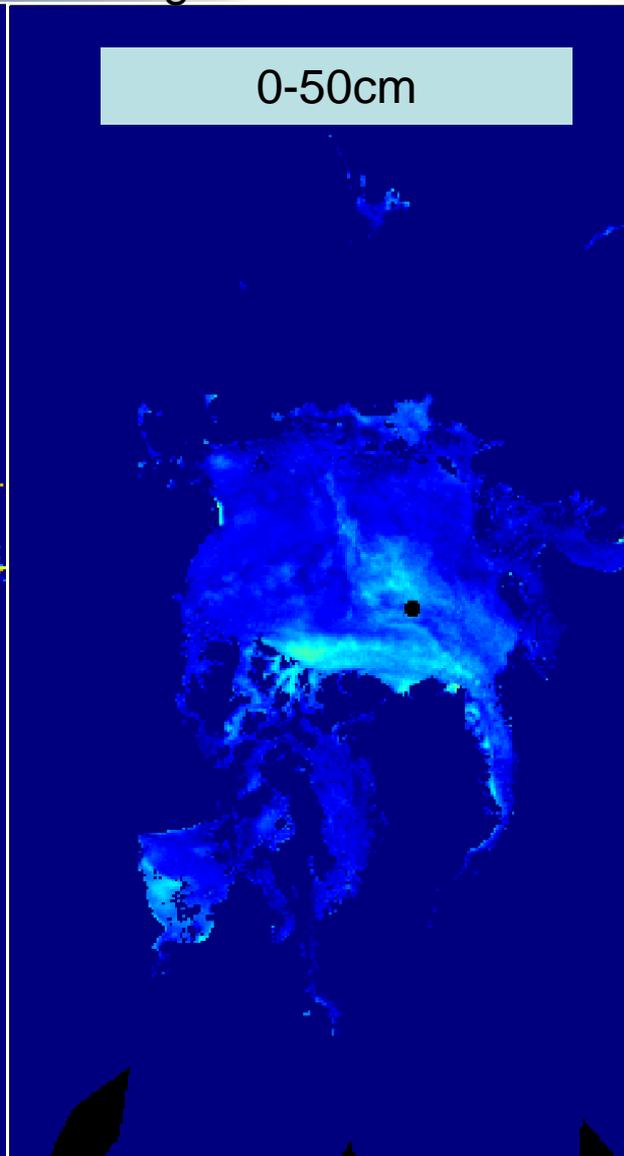
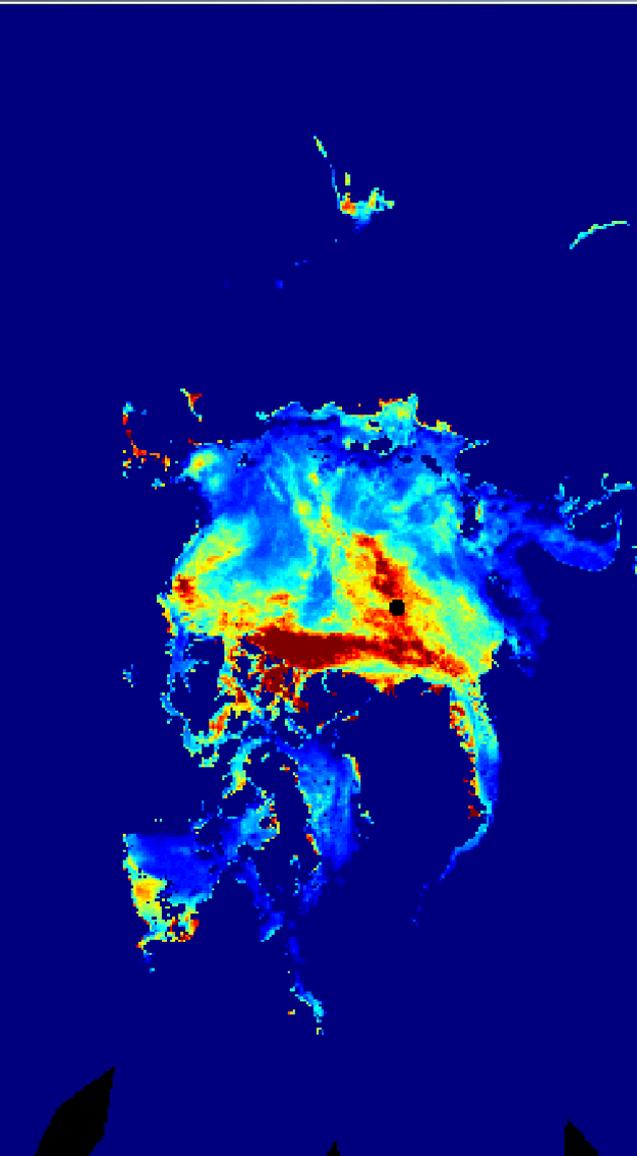
# The NRT AMSR2 daily L3 25km Tb and sea ice concentration polar grids

MC

Multi parameter  
regression

Tb reg. Model +  
OE

0-50cm



# Optimal estimation with multilayer sea ice emission model

## First guess

snow depth =  $1.77 + 0.017 \cdot T_{6v} - 0.028 \cdot T_{18v} + 0.0041 \cdot T_{37v}$  (Lise's equation)

$T_s = 1.53 \cdot T_{6v} - 136.6$  [220-272 K]

Salinity =  $8.5 + 100GR$  [0-7ppt]

Snow density =  $320 \text{ kg/m}^3$

Ice thickness =  $1 - 25GR$  [max 3.5 m]

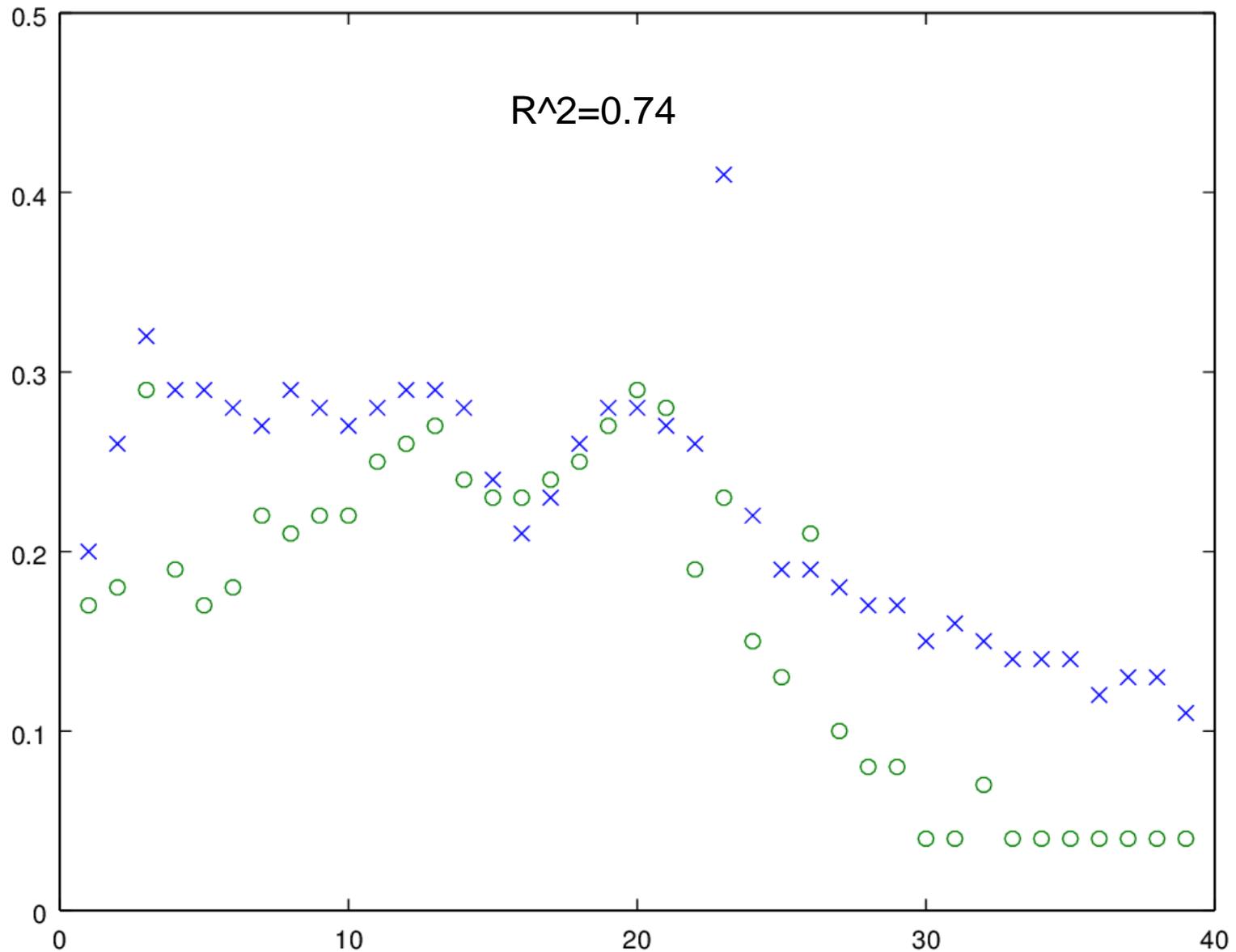
## Vertical distribution

Temperature distribution: linear profile

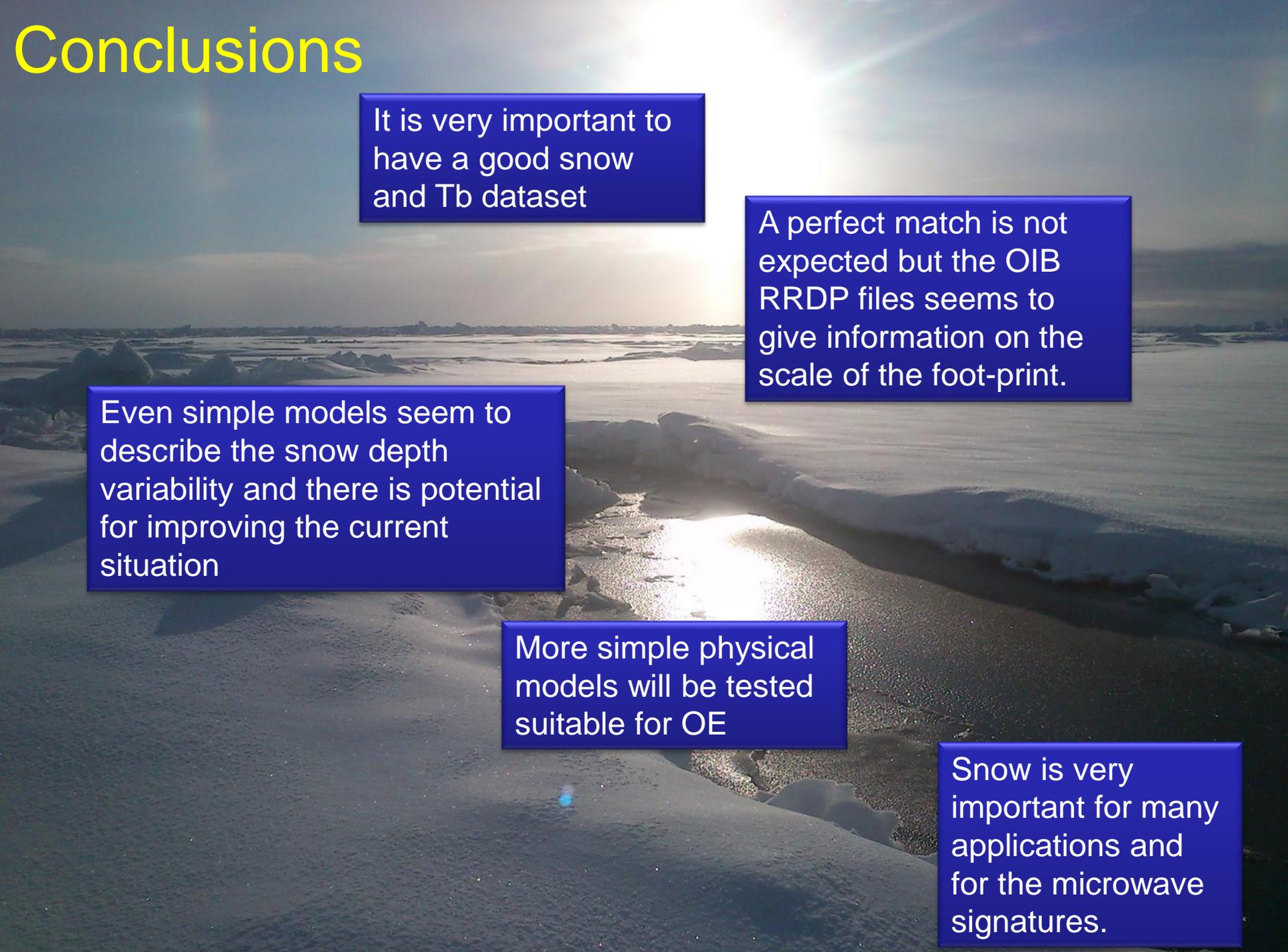
Snow grain size distribution: linear profile [0.07-0.3 mm] in ice 1.5 mm

Salinity distribution: linear distribution in ice, variable at top and fixed at 3 ppt at the bottom.

# Multilayer physical model and OE



# Conclusions



It is very important to have a good snow and Tb dataset

A perfect match is not expected but the OIB RRDP files seems to give information on the scale of the foot-print.

Even simple models seem to describe the snow depth variability and there is potential for improving the current situation

More simple physical models will be tested suitable for OE

Snow is very important for many applications and for the microwave signatures.